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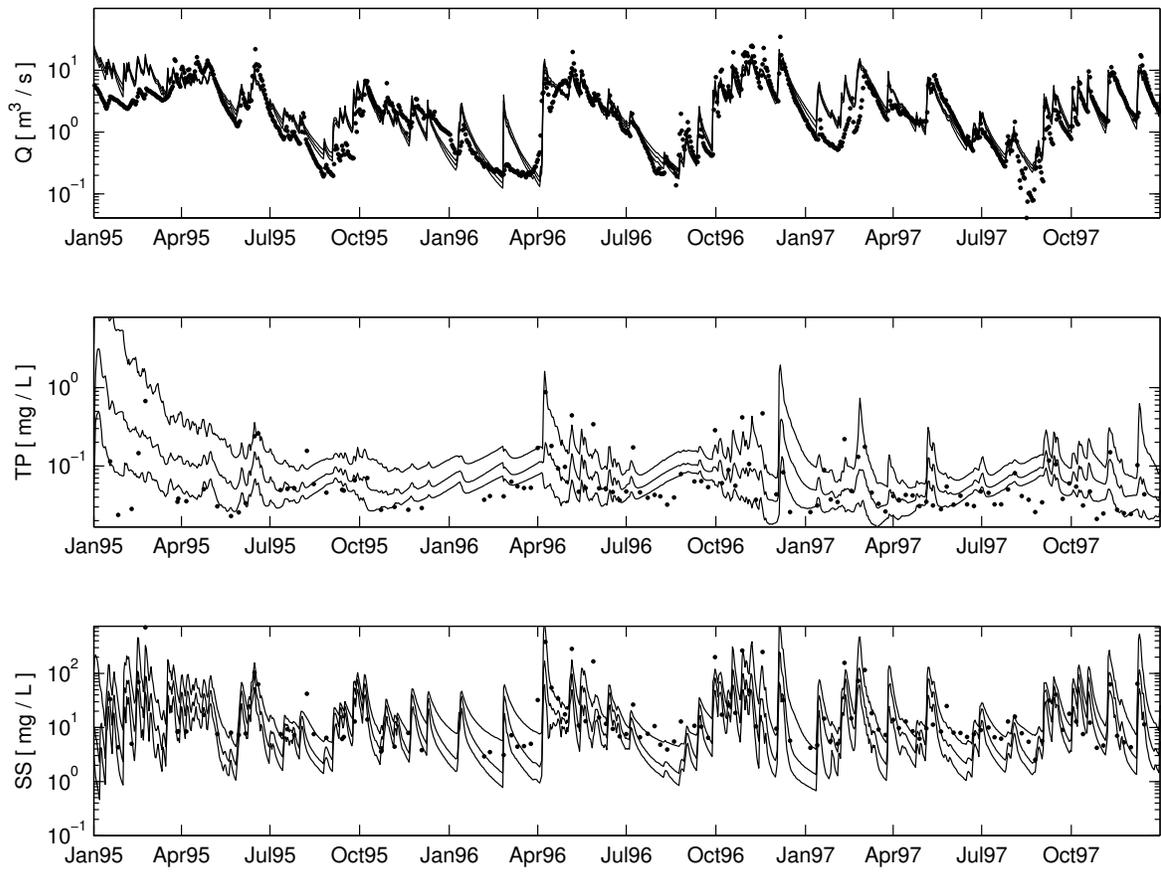


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Bayesian parameter estimation on INCA-P highlights the importance of parameter uncertainty in simulating future scenarios

Environmental impact statement for

Bayesian uncertainty assessment of a semi-distributed integrated catchment model of phosphorus transport.

Authors: Starrfelt, J. and Kaste, Ø.

The article *Bayesian uncertainty assessment of a semi-distributed integrated catchment model of phosphorus transport* details the application of a Bayesian scheme for uncertainty assessment of the Integrated Catchment model of Phosphorus (INCA-P). The scheme includes an autocalibration procedure for arriving at posterior distributions of selected parameters and uses these distributions to generate predictions of phosphorus transport under changed land uses, while including the uncertainty surrounding the parameters. This generates distributions of simulated outputs, i.e. probabilistic statements of predictions and can serve as a more solid foundation for management decisions under uncertainty.

1 **Title:**

2 **Bayesian uncertainty assessment of a semi-distributed integrated catchment model of**
3 **phosphorus transport.**

4

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12

13 **ABSTRACT**

14 Process-based models of nutrient transport are often used as tools for management of eutrophic waters,
15 as decision makers need to judge the potential effects of alternative remediation measures, under
16 current conditions and with future land use and climate change. All modelling exercises entail
17 uncertainty arising from various sources, such as the input data, selection of parameter values and the
18 choice of model itself. Here we perform Bayesian uncertainty assessment of an integrated catchment
19 model of phosphorus (INCA-P). We use an auto-calibration procedure and an algorithm for including
20 parametric uncertainty to simulate phosphorus transport in a Norwegian lowland river basin. Two
21 future scenarios were defined to exemplify the importance of parametric uncertainty in generating
22 predictions. While a worst case scenario yielded a robust prediction of increased loading of
23 phosphorus, a best case scenario only gave rise to a reduction in load with probability 0.78,
24 highlighting the importance of taking parametric uncertainty into account in process-based catchment
25 scale modelling of possible remediation scenarios. Estimates of uncertainty can be included in
26 information provided to decision makers, thus making a stronger scientific basis for sound decisions to
27 manage water resources.

28

29 Keywords: river basin, modelling, INCA-P, phosphorus, Bayesian inference, uncertainty

30

31 1. INTRODUCTION

32 Eutrophication, often caused by excessive inputs of phosphorus (P) and nitrogen (N) compounds from
33 agriculture, urban areas and scattered dwellings, is one of the main environmental concerns for rivers,
34 lakes and coastal waters around the world¹. Nutrients originate from a variety of sources, and one of
35 the challenging tasks in combatting eutrophication is to identify the main sources and quantify the
36 fluxes of nutrients promoting excessive algal growth. A good understanding of the predominant
37 nutrient sources and pathways is essential to be able to design and effectuate the most cost-effective
38 measures to reduce over-fertilisation of aquatic ecosystems. Process-based, integrated catchment
39 models can provide a tool by which the relative importance of various sources can be quantified, also
40 allowing for manipulative studies with simulation of ecosystem responses to various scenarios of
41 changed policies, land use or climate forcing².

42
43 Catchment-scale nutrient loss models have been developed for different purposes and thus cover a
44 wide range of complexity, level of process representation, input data requirements, and temporal and
45 spatial resolution³. The models simulate water, sediment and P transport from point and non-point
46 sources^{3,4}. Regardless of the span in complexity, all models have inherent uncertainties related to
47 their input data, parameter values and process representation⁵⁻⁸. Parts of this uncertainty can be
48 quantified; structural uncertainty (process representation) can be estimated by using different kinds of
49 models of the same phenomena. Uncertainty related to the input data (e.g. forcing data such as
50 meteorological parameters) can be treated by setting up models of observation errors. To address
51 parameter uncertainty various inverse modelling and automatic calibration techniques can be used.

52
53 Traditional use of catchment-scale process-based models of nutrient transport often does not take full
54 account of such uncertainties. Modellers have long grappled with the challenge to quantify the
55 uncertainty bounds on simulations generated in model applications. The use of manual calibration
56 techniques has been discussed in the hydrological literature for decades see e.g.⁹ and references therein, 10, 11.
57 Manual calibrations are often dependent on the modellers' subjective opinion of when a fit is good

58 enough. Manual calibration techniques often ignore the problem of equifinality see e.g.^{12, 13}; in
59 complex process-based models there are often several parameter sets which will yield the exact same
60 degree of fit, and choosing between them is therefore (in terms of degree of fit) arbitrary. This is
61 particularly problematic in scenario analyses, whereby a parameter can be rather insensitive or
62 insignificant within the calibration or validation period, but have a much larger impact on model
63 results under different scenario conditions. If uncertainty is included in the modelling process, then the
64 outcomes are not as categorical. Addition of uncertainty estimates to the model outcomes means that
65 the scientific basis for decision making becomes stronger for debate on these issues see e.g.^{14, 15} and
66 references therein. Auto-calibration e.g. PEST¹⁶ is a step in the right direction, yet often the goal of such
67 exercises is to arrive at *one* best set of parameters, in which equifinality and parameter uncertainty is
68 not fully addressed.

69

70 In recent years different sensitivity and uncertainty analyses have been applied to distributed, or semi-
71 distributed, hydrological and nutrient leaching models as INCA-N¹⁷, INCA-P¹⁸, and INCA-C / INCA-
72 Hg^{19, 20}. A General sensitivity analysis (GSA) was adapted to the INCA-type of models by Wade et al.
73²¹ and later adopted for INCA-C²². The GSA is performed to identify the model parameters that are
74 most influential in determining system behaviour²³ and is based on a comparison of prior and posterior
75 distributions of model parameter values. Additionally, uncertainty analyses has been carried out within
76 the generalised likelihood uncertainty estimation (GLUE) framework – for INCA-N²⁴ and INCA-P²⁵
77 – and other attempts using Monte Carlo sampling of parameter space²⁶ or parameter optimisation
78 algorithms²⁷.

79

80 One way to quantify the uncertainty surrounding parameters in catchment-scale nutrient loss models
81 and the resulting uncertainty in modelled outcome is to use Bayesian analysis. Using prior
82 distributions of parameter values and a formal likelihood, the procedure simulates a large number of
83 outcomes and the technique arrives at posterior distributions of parameter values in which we are more
84 confident in than our initial distribution. By using these posterior distributions to generate a range of
85 simulated outputs we thereby quantify the effect of parameter uncertainty in our predictions, including

86 scenario analyses. By using Bayesian methods for auto-calibration many of the shortcomings of
87 manual calibration can be circumvented: the method should in principle arrive at the same posterior
88 distributions for the same priors and input data (i.e. it is reproducible); the problem of equifinality is
89 taken into account by allowing for a multitude of parameter values even though they do not affect the
90 simulations in the calibration period; and the modelled outputs are presented as probability
91 distributions making the uncertainty in the parameter estimates in the model explicitly and visually
92 clear.

93
94 Though the principles of Bayesian analysis are fairly simple, in some cases the posterior distributions
95 are complex and hard to fully explore with traditional techniques. Much effort has been put into
96 improving these techniques e.g. ^{9,28-33}. Improvements of Markov Chain Monte Carlo methods of
97 posterior exploration include the use of several chains to better sample the full posterior distribution
98 and evolutionary algorithms that include some degree of selection among chains e.g. ^{32,33,34}. MCMC-
99 DiffeRential Evolution Adaptive Metropolis (DREAM) ³⁴ is a recently developed algorithm that
100 includes both several chains and evolutionary aspects and has been shown to be successful in
101 analysing complex hydrological models e.g. ^{35,36}.

102
103 Here our goal was to introduce Bayesian parameter uncertainty on one commonly used catchment
104 scale model of nutrient transport, the Integrated Catchment model for Phosphorus INCA-P; ^{18,37} by
105 applying the MCMC-DREAM algorithm. INCA-P was set up for simulating water flow, suspended
106 sediment and phosphorus (P) loads in Hobøl River, the main tributary to Lake Vansjø, SE Norway.
107 The catchment is characterised by high nutrient loads and recurrent blooms of toxin-producing
108 cyanobacteria, and an improved understanding of the sources of P and the uncertainties surrounding
109 them is instrumental in the improvement of the conditions of the lake. Following auto-calibration of
110 the model, two scenarios of land use changes were simulated and the uncertainty associated with
111 model parameters were taken into account through sampling of the posterior distributions.

112

113 **2. MATERIAL AND METHODS**

114 2.1 Study catchment

115 The Vansjø-Hobøl catchment comprises several small rivers and lakes, and one large lake (Vansjø)
116 (Figure 1). The catchment area is 690 km², and land use is dominated by agriculture (16%) and
117 forestry (80%)^{38, 39}. The agricultural production in the area consists mostly of grain production, with a
118 smaller fraction grass production (Table 1). Most of the catchment lies below 200 m elevation and is
119 covered by marine sediments deposited during the last glacial period. Mean annual rainfall is 810 mm
120 and the specific runoff is 14.4 L s⁻¹ km². Lake Vansjø is 35 km² in surface area and consists of two
121 major basins, Storefjorden and Vanemfjorden, with mean depths of 9 and 4 m, respectively. The main
122 inlet river, Hobøl, has a catchment area of 337 km² and discharges into the Storefjorden basin.

123

124 <<Figure 1 here>>

125

126 <<Table 1 here>>

127

128 2.2 Description of the INCA-P model

129 INCA-P is a process-based mass balance model designed for simulating the P dynamics in
130 catchments^{18, 40}, and was developed based on an integrated catchment model of nitrogen^{41, 42}. INCA-P
131 simulates the flow of water and addition/removal of P in the plant/soil system in different land use
132 types. The water containing both P and suspended particles is then routed downstream in the
133 catchment after accounting for direct effluent discharges and in-stream biological and sediment
134 processes. Effluent discharges, inorganic-P fertilisers and farmyard manure, slurry applications,
135 livestock wastes, and atmospheric deposition can be applied as input fluxes. The input fluxes and P
136 addition/removal processes are differentiated by land use type and varied according to environmental
137 conditions (e.g. soil moisture and temperature). The model also accounts for accumulated pools of
138 inorganic and organic P in the soil (in readily available and firmly bound forms), in groundwater and
139 in the stream reaches.

140

141 Since INCA-P is semi-distributed rather than fully-distributed, the catchment is decomposed into three
142 spatial levels: At level 1, the catchment is decomposed into sub-catchments. At level 2, each sub-
143 catchment is further decomposed into a maximum of six land use classes. At the third level, a generic
144 cell is then applied to each land-use type within each sub-catchment. Generalised equations define the
145 P transformations and stores within the cell, and six user-defined parameter sets derived through
146 calibration are used to simulate the differences between the land-use types. The numerical method for
147 solving the equations is based on the fourth-order Runge-Kutta technique, which allows a
148 simultaneous solution of the model equations, thereby ensuring that no single process represented by
149 the equations takes precedence over another¹⁸.

150

151 Being a model for river transport of nutrients, INCA-P requires hydrological forcing on daily time
152 steps. The HBV model⁴³ was used to produce the hydrological input time series. HBV is a semi-
153 distributed conceptual model with subdivision in altitude zones and distributed snow and soil moisture
154 descriptions. For Norwegian conditions, a version of the model developed by Killingtveit and Sælthun
155⁴⁴ and Sælthun⁴⁵ is most suitable. The general model structure consists of four main components: a
156 snow module, a soil moisture zone module, a dynamic module comprising the upper and lower soil
157 zone, and a routing module. The HBV model parameters can be grouped into two main categories,
158 free and confined parameters. The confined parameters are based on physical measurements and not
159 subject to calibration, for instance catchment area, area elevation curve and lake percentage. The free
160 parameters must be determined by calibration. The external forcings for HBV are time series of
161 precipitation and air temperature. The areal precipitation is based on point correction for rainfall and
162 snowfall measurement errors, fixed station weights and linear altitude increase of precipitation. HBV
163 produces daily hydrological input data for INCA-P -- hydrological effective rainfall (HER, the part of
164 the precipitation/snowmelt that contributes directly to runoff) and the soil moisture deficit (SMD). The
165 HBV model was not subject to uncertainty estimation, and was calibrated using the PEST procedure
166⁴⁶.

167

168 Among its outputs, INCA-P produces daily estimates of discharge Q (i.e., water flow), concentrations
169 of suspended solids (SS) and P at discrete points along a river's main channel. The different P fractions
170 simulated in the model are total dissolved P (TDP), particulate P (PP), and soluble reactive P (SRP).
171 The TDP and PP sum up to total P (TP). Because the model is semi-distributed, the hydrological and
172 nutrient fluxes from different land use classes and sub-catchment boundaries are modelled
173 simultaneously, and information is fed sequentially into a multi-reach river model. Therefore, spatial
174 variations in land use and farming practises can be taken into account, although the hydrological
175 connectivity of different land use patches is not modelled in the same manner as in a fully distributed
176 modelling approach.

177

178 **2.3 Reach structure and model parameters.**

179 The Hobøl River was divided into 5 sub-catchments (reaches, Figure 1) and five land use types were
180 defined (see Table 1). For some of the parameters, individual values can be given for each simulated
181 land use type or reach/subcatchment. The model parameter set can be roughly divided into four main
182 categories of parameters:

- 183 • land phase parameters and initial values (53 parameters per each land use class);
- 184 • in-stream parameters and initial values (8 parameters);
- 185 • reach parameters and initial values (46 parameters per each reach);
- 186 • subcatchment parameters and initial values (31 parameters per each subcatchment)

187 The main analysis was performed with the first 4 reaches, as the observations are from the end of
188 reach 4 (Kure). Parameters for reach 5 and its subcatchment were assumed to be identical to those of
189 reach 4 (except for land use proportions, length and area) for the posterior predictive simulations.

190

191 A total of 94 parameters involved in all phases of P transport were estimated. In addition, the effluent
192 inputs to the model were deemed to be uncertain and one parameter scaling these inputs was also
193 estimated. Some parameters of the model were set to be identical for specific land use classes,
194 subcatchments or reaches; this reduced the effective number of parameters to 49. The supplementary
195 material lists all parameters estimated, their priors and posteriors. The parameters not varied in this

196 exercise were based on an application of INCA-P in a smaller but similar catchment in the same area

197 ⁴⁷

198

199 **2.4 Formal likelihood.**

200 For application of Bayesian methods a formal likelihood relating the simulated variables and

201 observations needs to be defined. Following ^{34, 48, 49} we here develop a formal likelihood used for

202 analysing INCA-P in a Bayesian framework. Note that in a Bayesian framework the prior and

203 posterior are linked through a formal likelihood, as opposed to the informal metrics used in a GSA

204 approach²³. The output from INCA-P for which relevant observations are available are flow (Q , [m³s⁻

205 ¹]), suspended sediments in the water column (SS [mg L⁻¹]) and total P in the water column (TP [mg L⁻

206 ¹]).

207 Treating INCA-P as a model (h) yielding a set of outputs $\mathbf{Y}_o = \{y_{o,1}, \dots, y_{o,n_o}\}$ given a set of forcing

208 data (\mathbf{X}) and a set of parameters ($\boldsymbol{\theta}$)

$$\mathbf{Y}_i = h(\mathbf{X}, \boldsymbol{\theta}),$$

209 we get a set of residuals for each type of observations ($\hat{y}_{o,t}$)

$$\varepsilon_{o,t} = y_{o,t}(\mathbf{X}, \boldsymbol{\theta}) - \hat{y}_{o,t}, o = \{Q, SS, TP\}, t = 1, \dots, n_o.$$

210 We perform logarithmic transformations of our observed variables. The errors of the transformed

211 variables (i.e. the error model) are assumed normally distributed, and we use Gibbs sampling of the

212 error variances during the MCMC simulations. We thus assume that these residuals are mutually

213 independent (uncorrelated) and normally distributed with a variance associated with each type of

214 observation ($\sigma_o^2, o = \{Q, SS, TP\}$) and get an expression for the posterior probability density function

215 (pdf) ⁴⁸

$$p(\boldsymbol{\theta} | \hat{\mathbf{Y}}, \mathbf{X}, \boldsymbol{\sigma}^2) = c \cdot p(\boldsymbol{\theta}) \prod_o \prod_{t=1}^{n_o} \frac{1}{\sqrt{2\pi\sigma_o^2}} \times \exp\left(-\frac{(y_{o,t}(\mathbf{X}, \boldsymbol{\theta}) - \hat{y}_{o,t})^2}{2\sigma_o^2}\right),$$

216 where c is a normalizing constant, $p(\boldsymbol{\theta})$ is the prior probability of a set of parameters, combining the

217 data likelihood (the multiplicative part) and with a prior distribution using Bayes theorem. The

218 posterior ($p(\boldsymbol{\theta} | \hat{\mathbf{Y}}, \mathbf{X}, \boldsymbol{\sigma}^2)$) is thus the distribution of parameters given the model, input data and

219 observations. Working with the logarithm of likelihoods (\mathcal{L}) is often preferred both for simplicity and
 220 stability of calculations;

$$\mathcal{L}(\theta|\hat{\mathbf{Y}}, \mathbf{X}, \sigma^2) = \sum_o \left(-\frac{n_o}{2} \ln(2\pi) - \frac{n_o}{2} \ln(\sigma_o^2) - \frac{1}{2\sigma_o^2} \sum_{t=1}^{n_o} (y_{o,t}(\mathbf{X}, \boldsymbol{\theta}) - \hat{y}_{o,t})^2 \right)$$

221

222 2.5 MCMC-DREAM algorithm

223 Estimating posterior probability distributions (as $(\theta|\hat{\mathbf{Y}}, \mathbf{X}, \sigma^2)$) can be a fairly complicated exercise.

224 We proceed by using Markov Chain Monte Carlo techniques; instead of trying to get a function

225 describing the parameter distributions, an algorithm samples from it. MCMC algorithms start with a

226 given parameter set, runs the model and calculates a likelihood for this specific parameter

227 combination. A new (and usually fairly similar) *proposed* parameter set is then put into the model and

228 a new likelihood is calculated using this proposed parameter set. If the new parameter set yields a

229 better fit in terms of the likelihood (and the prior) it is then kept. If the new parameter set yields a

230 slightly worse fit it is sometimes kept (according to the Metropolis Hastings method), in all other cases

231 the *old* parameter set is kept. When run iteratively for a long period of time, this yields a *chain* of

232 parameter sets that have been kept (a Markov Chain). When a histogram is generated from these kept

233 parameter values, an estimate of the posterior distribution is achieved. Usually the first portion of the

234 chain is discarded, to reduce the impact of badly chosen initial parameter guesses.

235

236 To estimate the posterior probability density ($p(\theta|\hat{\mathbf{Y}}, \mathbf{X}, \sigma^2)$) we utilize the MCMC-DREAM

237 algorithm³⁴. Essentially this algorithm works by simulating several Markov Chains at the same time,

238 which sample parameter proposals from distributions that are automatically tuned in both magnitude

239 and direction during the evolution of the chains. The likelihood of these parameter proposals are then

240 evaluated in a traditional Metropolis Hastings algorithm. The algorithm is succinctly described in³⁴.

241 Delayed rejection of parameter proposals was originally included in the DREAM algorithm, but we

242 have not included this feature in our application.

243

244 Included in the DREAM algorithm are checks for convergence of chains through the calculation of
245 Gelman-Rubin statistics^{34, 50} after which the chains are runs for a given number of iterations to sample
246 the posterior distributions. These chains were then stored and resampled to simulate all 5 reaches for a
247 longer period (1995-2005), a form of posterior predictive modelling⁵⁰.

248

249 **2.6 Routines for model calibration and evaluation of outputs**

250 Model code for MCMC-DREAM evaluation of INCA-P was coded in Matlab⁵¹ and utilized a
251 command-line version of INCA-P. Matlab code was used to generate the proposal values and store the
252 parameter and input files for each chain, after which INCA-P was called and output stored. The INCA-
253 P output was then read by the code and evaluated according to the algorithm described above. Due to
254 the computational cost we simulated INCA-P for years 1995-1997 and used observations from 1996
255 and 1997 to calculate the likelihoods used in parameter estimation. After convergence of the algorithm
256 we then re-ran INCA-P using the estimated parameters sampled from the chains for a baseline run to
257 estimate yearly loads from the river (including all 5 reaches) as well as for the scenarios for the period
258 1995-2005.

259

260 **2.7 Posterior predictive simulations and scenario definitions.**

261 After distributions of parameters were estimated posterior predictive simulations were performed, i.e.
262 parameter values were sampled from the converged chains and INCA-P was rerun with these sets of
263 parameters. This was performed for an extended period (1995-2005) for both a baseline (i.e. with
264 current land use and management) and two scenarios. These future scenarios were developed together
265 with stakeholders at a workshop discussing the future of Vansjø catchment. Two possible futures were
266 envisioned and further limited to changes easily implemented in INCA-P. In the worst case scenario,
267 greater demand for agricultural products was expected and we parameterized this as 25 % of all
268 grassland was put into vegetable production (with a corresponding fertilizer regime), 10 % of the
269 forest was made into grass production and the amount of fertilizer application was overall increased by
270 25% for all relevant land uses. Best case scenario included a 90 % reduction in effluent inputs (from
271 scattered dwellings and waste water treatment plants), 25% of land allocated to vegetables and crops

272 were changed to grassland production and 50 % direct decreases in fertilizer amounts were
273 implemented. The reduction in effluent inputs was implemented as a second scaling factor in addition
274 to the scaling factor estimated in the parameter estimation, thus still including the uncertainty of these
275 inputs.

276

277 **2.8 Data sources**

278 **Meteorology:** Daily data on temperature, precipitation, and snow cover are obtained from
279 three stations (1715 Rygge; 1750 Fløter; 378 Igsi) operated by the Norwegian Meteorological Institute
280 (met.no).

281 **Hydrology:** Daily flow at the gauging station 3.22.0.1000.1 Høgfoss are obtained from the
282 Norwegian Water Resources and Energy Directorate (NVE). The HBV model was calibrated using the
283 PEST procedure⁴⁶ for the period 1.9.1990 to 31.8.2000, achieving a Nash Sutcliffe value of 0.78 and
284 0% volume error. The HBV was then run to simulate SMD and HER up to 31.12.2005.

285 **Water chemistry:** Water chemistry data come from the MORSA monitoring programme,
286 conducted by Norwegian Institute for Agricultural and Environmental Research (Bioforsk) and
287 Norwegian Institute for Water Research (NIVA). Data from the monitoring station Kure (reach 4)
288 were used for calibration of INCA-P.

289 **Land-cover:** General land cover data are obtained from the Norwegian Forest and Landscape
290 Research Institute, whereas more detailed information about land use fertilisation regimes on
291 agricultural fields are provided by Bioforsk.

292 **Municipal wastewater:** Nutrient outputs from sewage treatment plants are obtained from
293 Statistics Norway and the database KOSTRA. Outputs from scattered dwellings are provided by
294 Bioforsk and their information system “GISavløp”.

295

296 **3. RESULTS**

297 The MCMC-DREAM algorithm successfully managed to quantify uncertainty associated with the
298 selected parameters in the INCA-P model. The 40 chains used in the analysis achieved a Gelman

299 Rubin diagnostic of < 1.2 for all parameters after about 3350 iterations, indicating that the algorithm
300 managed to converge and sample the posteriors⁵⁰. The chains were run for 10 000 iterations and the
301 last 2500 iterations were considered to satisfactorily sample the posterior distributions, with an
302 acceptance rate of 0.288, indicating well-mixed sampling. INCA-P is a computation-time demanding
303 model with each simulation in the parameter estimation phase requiring about 8 minutes to run,
304 leading to a CPU demand of approximately 54 000 hrs. This was drastically reduced by using a
305 multicore computer and the 10 000 iterations were completed in about two weeks.

306 Several of the marginal posterior distributions were dramatically smaller than the priors, but for others
307 a wide range was still evident in the posteriors, indicating that the parameters did not have clear non-
308 interactive effects on the likelihood. Several of the parameters, however, exhibit strong covariances,
309 which are not evident in marginal distributions. Marginal distributions of all parameters and an
310 example of the importance of covariance of parameters are presented in the supplementary material.

311 Figure 2 presents the simulated flow, total P and suspended sediment for the fourth reach in the
312 calibration period 1995-1997 as well as observed values for 1996 and 1997, clearly showing the
313 parametric uncertainty has low impacts on the flow (top panel), while giving rise to considerable
314 uncertainty in P and particle matter predictions. For P the simulations span up to a factor of
315 magnitude, particularly so in cases of heavy rain, a pattern opposite to the one exhibited for sediments
316 in which the variation in predictions seem to narrow in the case of rain events but increase in periods
317 of lower flow. While the confidence bands of both P and sediments are spanning up to an order of
318 magnitude, several of the observations fall outside the prediction interval. This is partly due to the
319 exclusion of the impact of the error variance estimated (see supplementary material), which would
320 increase the uncertainty in the predictions considerably. It also underlines the importance of additional
321 data potentially narrowing the priors for the parameters to get at both more constrained and more
322 accurate predictions.

323

324 <<Figure 2 here>>

325

326 In the posterior predictive simulations for the period 1995-2005 (Figure 3) the flood in the end of 2000
327 and early 2001 is clearly evident, resulting in a dramatic increase in the uncertainty of the predictions.
328 The winter 2000/2001 was mild with heavy rain, increasing the transport of sediments and particles to
329 the lake (unpublished data). In our simulations, the effect of the flooding on P transport is noticed as
330 an increase in the uncertainty in predictions of daily concentrations (Figure 3) and yearly loads (Figure
331 4) of P.

332

333 <<Figure 3 here>>

334

335 <<Figure 4 here>>

336

337 The scenarios were developed with the aim of spanning the potential futures of the catchment by
338 positing best case and worst case developments, and the majority of the simulations predict a
339 corresponding decrease or increase in the yearly loads (Table 2). The median reduction in P loads
340 under the best case scenario was about 10%. Quite a few simulations (22%) actually showed an
341 increase in annual loads, and conversely, taking parameter uncertainty into account yields a 78%
342 probability of a decrease in loading. The parameter sets yielding an increase of loading under the best
343 case scenario were the initial values for P storage in the changes land use classes and the parameter
344 scaling the effluent inputs. Essentially, with the degree of uncertainty surrounding the parameters for
345 the grass and crop areas from our calibration, there is a corresponding degree of uncertainty with the
346 effect of the scenario.

347

348 <<Table 2 here>>

349

350 **4. DISCUSSION**

351 Our analysis of the parametric uncertainty in the INCA-P model is the first example of the use of a
352 Bayesian framework on any model in the INCA-family. Earlier attempts at analysing uncertainty on

353 INCA models are limited to one application of the generalized likelihood uncertainty estimation
354 GLUE, see ^{13, 52} technique on INCA-P ²⁵ and other attempts using Monte Carlo sampling of parameter
355 space of INCA-N, the nitrogen cousin of INCA-P ^{24, 26}. Even though GLUE and Bayesian approaches
356 formally have different fundamental philosophies, predicted outcomes can be fairly similar, see e.g. ⁴⁸.
357
358 Though our analysis shows that it is possible to both estimate parameters and perform scenario
359 simulations using INCA-P while taking the parameter uncertainty into account, several assumptions
360 were needed. Firstly, as opposed to the GLUE framework, a proper Bayesian analysis requires a
361 formal definition of the error model (i.e. how predictions relate to observations). As a first step we
362 assumed that these errors were uncorrelated, an assumption which is clearly invalid for most
363 hydrological data. Though we did not include them here, there are more sophisticated ways to define
364 the error model which could take into account the auto-correlated nature of hydrological data e.g. ⁵³.
365 Secondly, the INCA-P model can only simulate nutrient transport using hydrological input which itself
366 is output from an external model (HBV), and by using a single parameterization of HBV much of the
367 uncertainty associated with parameters in the hydrological model was ignored. Ideally, the whole
368 model chain HBV → INCA-P should be treated as one model, and analysed in a Bayesian framework,
369 but this has to be left for a future exercise. Lastly, as with any exercise where model parameters are
370 varied, several parameters of the model are still left fixed, even though they entail their own
371 uncertainty. A final improvement to the analysis would be to better select which parameters to include
372 in the Bayesian estimation, and also gather more information for defining the priors.
373
374 The outcome of the posterior predictive simulations (Figure 3, 4 and Table 2) of the scenarios
375 underlines the importance of taking parameter uncertainty into account particularly when models such
376 as INCA-P are used for policy purposes. When the parameter values that gave the highest degree of fit
377 (in terms of the Nash Sutcliffe criterion) are used alone, the model predicts a 19% decrease in the
378 yearly loads of P to Lake Vansjø under the best case scenario. Though the scenarios here are not
379 meant as real policy alternatives, a simulated decrease in loading of 19% could be seen as an argument
380 for management of the catchment in the direction of this scenario (e.g. implementing incentives for

381 farmers to switch from cereal crops to grass production). When the full uncertainty of the parameter
382 values are taken into account, however, the median prediction is a much lower reduction in load (11
383 %) and about 22% of the simulations actually showed an increase in loads given the best case
384 scenario. In summary, with the current knowledge and quantification of our parametric uncertainty
385 there's only about 80 % chance that such mitigation measures would actually decrease the loading
386 from the river. An inspection of which parameter sets that yield an increase in loading under the best
387 case scenario (not shown) reveals that these predictions arise from parameter combinations of high and
388 low initial P values for the grass and crop land use classes respectively. In addition this increased
389 loading under a proposed better future are predicted when the parameter scaling the effluent inputs are
390 in the lower range, reducing the impact of effluent reduction (see supplementary information for all
391 parameters estimated). As the observations against which the model is calibrated are from the river
392 itself, calibration can not distinguish between exports from different land use classes (unless they
393 change over time, which they do not in our case). This is a good example of equifinality; parameters
394 detailing the mobilization and initial storage of P in different land use classes will covary negatively
395 and can result in the same degree of fit for a wide range of parameter settings. Instead of choosing one
396 single parameterization for this process, our approach captures the uncertainty surrounding this
397 process through our scenario simulations.

398

399 Information regarding the probability of success of mitigation measures is highly useful for
400 policymakers, and knowing that a particular management decision has a 1 in 5 chance of having the
401 opposite effect would be beneficial. This underlines the importance of communicating uncertainty in
402 the interaction between science and policy, and not to obscure possible risks of non-optimal effects or
403 even failure when implementing a mitigation measure for a perspective see ¹⁴. It also underlines the
404 importance of equifinality issues in models such as INCA-P; even though the best calibration (based
405 on classical NS values) predicts a certain reduction in load, incorporating a wider range of both
406 plausible and probable parameter ranges may have a substantial effect on the probability of modelled
407 outcomes under scenario analysis.

408

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416

417 **TABLES**

418

419 Table 1. Reaches/subcatchments defined within the Vansjø-Hobøl catchment, including land use
 420 information.

Reach	Name	Size	Wetland	Forest	Grass	Crops	Vegetables
no.		<i>km²</i>	<i>km²</i>	<i>km²</i>	<i>km²</i>	<i>km²</i>	<i>km²</i>
1	Mjær	146.32	13.2	125.8	1.5	5.9	0.0
2	Tomter	88.16	0.9	66.1	1.8	19.4	0.0
3	Kråkstad	50.6	0.5	33.9	1.0	14.7	0.5
4	Kure	19.32	0.4	12.6	0.6	5.8	0.0
5	Våler	32.41	1.0	25.0	0.6	5.8	0.0

421

422

423 Table 2. Summary results of scenario runs including parameter uncertainty. Scenarios are run by sampling parameter sets from the Markov Chains, and for
 424 each sampled parameter set we run all scenarios. The ratio of change is the ratio of individual yearly loads for each specific parameter set under all scenarios
 425 (i.e. the only difference is the scenario parameters), i.e. $r = Load_{year, scenario} / Load_{year, baseline}$. First year of the simulation (1995) is excluded. The single INCA-P
 426 parameterization with the highest Nash-Sutcliffe value for total P in the calibration period 1996-1997 ($NS_{TP} = 0.4616$) gives the following median loads in the
 427 period 1996-2005; 15.00 T/y (Baseline), 18.21 T/y (Worst Case) and 11.96 T/y (Best Cse), with median ratios of change at 1.22 and 0.81 respectively. Please
 428 note that the ratio of change to the right are calculated on the distribution of yearly values and not on the median values reported to the left.

Percentiles and mean yearly loads

1996-2005 (ton / year)

Percentiles and mean ratio of change.

	1996-2005 (ton / year)			Percentiles and mean ratio of change.		
	2.5%	50% (mean)	97.5%	2.5%	50% (mean)	97.5%
Baseline	5.18	11.96 (24.24)	108.13	1	1	1
Worst Case	7.18	16.73 (30.48)	131.61	1.02	1.30 (1.41)	2.39
Best Case	4.07	11.10 (23.03)	103.41	0.69	0.89 (0.92)	1.39

429 **FIGURE CAPTIONS**

430

431 Figure 1. The Vansjø-Hobøl catchment. The 5 subcatchments that constitute Vansjø's main tributary
432 river, Hobøl flowing from North to south and a map of Norway showing location of the catchment.
433 The observations against which the model was calibrated was collected at the southern end of the Kure
434 subcatchment. Lake Vansjø, with its basins Vanemfjorden to the west and Storefjorden to the east, can
435 be seen at the lower part of the catchment.

436

437 Figure 2. Confidence intervals for simulated flow, P and sediment 1995-1997. Simulated flow, total P
438 and suspended sediment (log scale) for the burn-in (1995) and calibration period (1996-1997) for the
439 4th reach (Kure). The lines represent the daily 2.5, 50 and 97.5 percentiles of the output from the
440 posterior predictive simulations performed with the converged chains sampled 100 times (i.e. from
441 4000 simulations). Observations are represented by dots. Error variance is not included in these plots,
442 so the uncertainty in predictions arises solely from parameter uncertainty.

443

444 Figure 3. Confidence intervals for simulated flow, P and sediment 1995-2005. Simulated flow, total P
445 concentrations and suspended sediment (log scale) for the period 1995-2005, with lines representing
446 the 2.5, 50 and 97.5 percentiles of the posterior predictive simulations for the 4th reach. Observations
447 are represented by dots. When running the posterior predictive modelling for the whole period 196 of
448 the parameter sets resulted in crashes of INCA-P and confidence intervals are calculated from the
449 3804 successful runs. Note the flood occurring Oct-Nov 2000, and how it affects the uncertainty in the
450 predicted total P concentrations. Error variance is not included in these plots, so the uncertainty in
451 predictions arises solely from parameter uncertainty.

452

453 Figure 4. Yearly loading from Hobøl River. Simulated yearly loading of P (log scale) from Hobøl
454 River into Lake Vansjø in metric ton per year, i.e. from the 5th reach. The lines represent the 2.5, 50
455 and 97.5 percentiles. The high degree of uncertainty in 1995 is due to initial conditions that have
456 reduced impact outside the burn-in year. Also note how the flooding in Autumn 2000 affects both the

457 median and the surrounding uncertainty. The median loading for the whole period is approximately 12
 458 tons/year (see Table 2).

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