

JAAS

Accepted Manuscript



This is an *Accepted Manuscript*, which has been through the Royal Society of Chemistry peer review process and has been accepted for publication.

Accepted Manuscripts are published online shortly after acceptance, before technical editing, formatting and proof reading. Using this free service, authors can make their results available to the community, in citable form, before we publish the edited article. We will replace this *Accepted Manuscript* with the edited and formatted *Advance Article* as soon as it is available.

You can find more information about *Accepted Manuscripts* in the [Information for Authors](#).

Please note that technical editing may introduce minor changes to the text and/or graphics, which may alter content. The journal's standard [Terms & Conditions](#) and the [Ethical guidelines](#) still apply. In no event shall the Royal Society of Chemistry be held responsible for any errors or omissions in this *Accepted Manuscript* or any consequences arising from the use of any information it contains.

1
2
3 1 **Optimization of a ^{48}Ca – ^{43}Ca double-spike MC-TIMS method for measuring Ca isotope ratios**
4 2 **($\delta^{44/40}\text{Ca}$ and $\delta^{44/42}\text{Ca}$): limitations from filament reservoir mixing**

5 3
6 4
7 4 **Gregory O. Lehn^{a*}, Andrew D. Jacobson^a**

8 5
9 6 ^a *Department of Earth and Planetary Sciences, Northwestern University, Technological Institute, 2145*
10 7 *Sheridan Road, Evanston, Illinois 60208, USA*

11 8
12 9 *Corresponding author gregorylehn@gmail.com, Tel: 1-847-467-2467.

13
14 10 **Abstract**

15
16 11 We used a Monte Carlo error model to optimize a ^{48}Ca – ^{43}Ca double-spike technique for measuring Ca
17 12 isotope ratios ($\delta^{44/40}\text{Ca}$ and $\delta^{44/42}\text{Ca}$) by Multi-Collector Thermal Ionization Mass Spectrometry (MC-
18 13 TIMS). The model considers errors for counting statistics and Johnson noise, as well as changes in
19 14 collector cup efficiency (drift). For a 20V ^{40}Ca ion-beam implemented in a three-hop, dynamic multi-
20 15 collection routine, the model predicts that a wide range of $^{48}\text{Ca}/^{43}\text{Ca}$ and spike/sample ratios should yield
21 16 internal precisions ($2\sigma_{\text{SEM}}$) of 0.015 – 0.020‰ for $\delta^{44/40}\text{Ca}$ and 0.025 – 0.030‰ for $\delta^{44/42}\text{Ca}$. Using a
22 17 Thermo Fisher MC-TIMS (Triton), we tested $^{48}\text{Ca}/^{43}\text{Ca}= 1.5$ [$^{43}\text{Ca}/(^{48}\text{Ca}+^{43}\text{Ca})= 0.40$ mol/mol] and
23 18 spike/sample= 0.66 [$\text{Ca}_{\text{dsp}}/(\text{Ca}_{\text{dsp}}+\text{Ca}_{\text{smp}})= 0.40$ mol/mol] by repeatedly analyzing OSIL Atlantic
24 19 seawater, NIST SRM 915a, NIST SRM 915b, USGS BHVO-1, and CaF_2 over 4 sessions spanning 1
25 20 month. While the measured internal precision generally agreed with model predictions, external
26 21 reproducibility ($2\sigma_{\text{SD}}$) was much worse than expected. For the 81 measurements made, the average
27 22 external reproducibility was 0.223‰ for $\delta^{44/40}\text{Ca}$ and 0.126‰ for $\delta^{44/42}\text{Ca}$. After processing raw data
28 23 through the double-spike equations, nearly all fractionation-corrected ratios showed remnant fractionation
29 24 patterns. Such patterns reflect deviation from ideal exponential mass-fractionation due to mixing of
30 25 multiple, independently fractionating reservoirs on the filament. Additional model simulations, as well as
31 26 comparison against $\delta^{44/40}\text{Ca}$ values determined with a ^{43}Ca – ^{42}Ca double-spike, support the concept of an
32 27 “average mass rule,” which states that inaccuracies in fractionation-corrected data are greater for isotope
33 28 ratios having an average mass further away from the average mass of the normalizing ratio. Until
34 29 advancements are made to eliminate filament reservoir effects, ^{43}Ca – ^{42}Ca and ^{46}Ca – ^{43}Ca double-spikes
35 30 should yield the most precise $\delta^{44/40}\text{Ca}$ and $\delta^{44/42}\text{Ca}$ values, respectively, when using MC-TIMS. Within the
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

1
2
3 31 limits of the ^{48}Ca - ^{43}Ca double-spike technique, we observed no evidence for ^{40}Ca enrichments among the
4
5 32 standards analyzed. Finally, we found that sample matrix effects do not influence the quality of Ca
6
7 33 isotope measurements by MC-TIMS, and we tentatively propose that the external reproducibility
8
9 34 determined from the repeated analysis of standards can represent the uncertainty of a single sample
10
11 35 analysis.

14 36 **1. Introduction:**

16 37 Calcium (Ca) isotope measurements are relevant to diverse fields, including geochemistry,
17
18 38 cosmochemistry, biology, medicine, archaeology, and anthropology¹⁻⁶. Many studies employ “double-
19
20 39 spike” techniques to correct for instrumental mass-fractionation when analyzing $^{40}\text{Ca}/^{44}\text{Ca}$ ratios by multi
21
22 40 collector-thermal ionization mass spectrometry (MC-TIMS). The details and virtues of double-spiking
23
24 41 have been extensively discussed in several publications⁷⁻¹⁰. The technique involves equilibrating a sample
25
26 42 solution containing the element of interest with another solution containing a well-characterized mixture
27
28 43 of two artificially enriched isotopes of the same element. Both the ratio of the two artificially enriched
29
30 44 isotopes (the “double-spike” ratio) and the ratio of the double-spike to the sample (the “spike/sample”
31
32 45 ratio) must be carefully optimized to minimize propagation of errors and achieve high precision results.
33
34 46 Practically every double-spike combination, such as ^{48}Ca - ^{43}Ca , ^{48}Ca - ^{42}Ca , ^{46}Ca - ^{43}Ca , and ^{43}Ca - ^{42}Ca , has
35
36 47 been implemented to measure $^{40}\text{Ca}/^{44}\text{Ca}$ ratios by MC-TIMS¹¹⁻¹⁵. However, while different methods
37
38 48 appear reasonably accurate, large differences in external reproducibility exist for reasons that are not
39
40 49 entirely understood but appear to extend beyond simple mathematical optimization of the ratio defined
41
42 50 above^{9, 10}. Possible explanations for decreased precision include ion optical aberrations¹³⁻¹⁵, sample matrix
43
44 51 effects^{10, 16}, filament reservoir mixing¹⁶⁻¹⁸, changes in collector cup efficiency (drift)^{19, 20}, and choice of the
45
46 52 double-spike pair relative to the sample isotope ratio of interest²¹. These issues plague analysis of many
47
48 53 isotope systems, but they are especially problematic for resolving the small abundance variations inherent
49
50 54 to the Ca isotope system. To increase statistical confidence, it has become common practice to repeatedly
51
52 55 analyze the same sample at the expense of efficiency and throughput. Moreover, in petrology^{1, 22, 23},
53
54 56 cosmochemistry²⁴, and some areas of biogeochemistry^{15, 25, 26}, it is desirable to determine whether
55
56
57
58
59
60

1
2
3 57 variation of the $^{40}\text{Ca}/^{44}\text{Ca}$ ratio reflects natural mass-fractionation, radiogenic ingrowth of ^{40}Ca from the
4
5 58 decay of ^{40}K ($\tau_{1/2}=1.277 \times 10^9$ yr), or some combination of the two. This question can be addressed by
6
7 59 measuring a second isotope ratio, e.g., $^{42}\text{Ca}/^{44}\text{Ca}$, $^{40}\text{Ca}/^{42}\text{Ca}$, or $^{43}\text{Ca}/^{44}\text{Ca}$ ^{1, 22-25}. It is also possible to
8
9
10 60 measure $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with one double-spike pair but implement two sets of “true” ratios in the data
11
12 61 reduction algorithms: the so-called “normal ratios”¹¹ and unspiked, internally normalized sample ratios²⁴,
13
14 62 ²⁷. This latter approach requires two mass spectrometric runs.

15
16 63 Building on research by Holmden (2005) and Holmden and Bélanger (2010), Lehn et al. (2013) used
17
18 64 a Monte Carlo error model to optimize a ^{43}Ca – ^{42}Ca double-spike method for measuring $^{40}\text{Ca}/^{44}\text{Ca}$ ratios
19
20 65 by MC-TIMS^{14, 15, 20}. The model considers errors for counting statistics and Johnson noise, and it accounts
21
22 66 for collector cup drift. Lehn et al. (2013) verified the model predictions by measuring four commonly
23
24 67 analyzed Ca isotope standards (n=171) and demonstrated that $^{40}\text{Ca}/^{44}\text{Ca}$ ratios can be quantified with an
25
26 68 internal precision ($2\sigma_{\text{SEM}}$) of $\pm 0.024\%$ and a long-term, external reproducibility ($2\sigma_{\text{SD}}$) of $\pm 0.041\%$,
27
28 69 which represents a two-fold or better improvement over previously utilized methods²⁰. The present study
29
30 70 reports an effort where we used the Lehn et al. (2013) error model to optimize a ^{48}Ca – ^{43}Ca double-spike
31
32 71 MC-TIMS technique²⁰. We selected this double-spike combination because it in theory allows
33
34 72 simultaneous analysis of $^{40}\text{Ca}/^{44}\text{Ca}$ and $^{42}\text{Ca}/^{44}\text{Ca}$ ratios, thus offering potential to resolve mass-dependent
35
36 73 fractionation from radiogenic ^{40}Ca enrichments in the context of a single run. (A ^{46}Ca – ^{43}Ca double-spike
37
38 74 offers the same potential, but we were unable to obtain a sufficiently pure ^{46}Ca spike.) In addition, other
39
40 75 theoretical studies have suggested that a ^{48}Ca – ^{43}Ca double-spike should yield more precise data than a
41
42 76 ^{43}Ca – ^{42}Ca double-spike²⁸, but in general, such assertions are challenging to verify because different
43
44 77 studies employ slightly different models, assumptions, and measurement routines. Thus, in this study, we
45
46 78 took care to reproduce as closely as possible the approach taken in Lehn et al. (2013) so that we could
47
48 79 normalize variables and rigorously compare the two techniques²⁰.

53 80 **2. Methods**

55 81 *2.1 Theoretical optimization of the ^{48}Ca – ^{43}Ca double-spike using a Monte Carlo error model*

We used the Monte Carlo error model presented in Lehn et al. (2013) to optimize a ^{48}Ca - ^{43}Ca double-spike method allowing the simultaneous analysis of $^{40}\text{Ca}/^{44}\text{Ca}$ and $^{42}\text{Ca}/^{44}\text{Ca}$ ratios²⁰. This section provides a brief overview of the model, and where appropriate, we present equations specific to the ^{48}Ca - ^{43}Ca double-spike method. Following Eugster et al. (1969), we derived the following equations for the determination of $^{40}\text{Ca}/^{44}\text{Ca}$ and $^{42}\text{Ca}/^{44}\text{Ca}$ ratios using a ^{48}Ca - ^{43}Ca double-spike:

$$\left(\frac{^{48}\text{Ca}}{^{43}\text{Ca}}\right)_{dsp} = \left(\frac{^{48}\text{Ca}}{^{43}\text{Ca}}\right)_{mix} + \left[\left(\frac{^{44}\text{Ca}}{^{43}\text{Ca}}\right)_{mix} - \left(\frac{^{44}\text{Ca}}{^{43}\text{Ca}}\right)_{dsp}^{tru} \right] \cdot \left[\frac{\left(\frac{^{48}\text{Ca}}{^{43}\text{Ca}}\right)_{mix} \cdot \left(\frac{^{43}\text{Ca}}{^{44}\text{Ca}}\right)_{smp}^{tru} - \left(\frac{^{48}\text{Ca}}{^{44}\text{Ca}}\right)_{smp}^{tru}}{1 - \left(\frac{^{44}\text{Ca}}{^{43}\text{Ca}}\right)_{mix} \cdot \left(\frac{^{43}\text{Ca}}{^{44}\text{Ca}}\right)_{smp}^{tru}} \right] \quad (1)$$

$$\left(\frac{^{XX}\text{Ca}}{^{44}\text{Ca}}\right)_{smp} = \frac{\left(\frac{^{XX}\text{Ca}}{^{43}\text{Ca}}\right)_{mix} + \left[\frac{\left(\frac{^{XX}\text{Ca}}{^{43}\text{Ca}}\right)_{mix} \cdot \left(\frac{^{44}\text{Ca}}{^{43}\text{Ca}}\right)_{dsp}^{tru} - \left(\frac{^{XX}\text{Ca}}{^{43}\text{Ca}}\right)_{dsp}^{tru}}{\left(\frac{^{44}\text{Ca}}{^{43}\text{Ca}}\right)_{mix} - \left(\frac{^{44}\text{Ca}}{^{43}\text{Ca}}\right)_{dsp}^{tru}} \right]}{\left(\frac{^{44}\text{Ca}}{^{43}\text{Ca}}\right)_{mix}} \cdot \left[1 - \left(\frac{^{44}\text{Ca}}{^{43}\text{Ca}}\right)_{mix} \cdot \left(\frac{^{43}\text{Ca}}{^{44}\text{Ca}}\right)_{smp}^{tru} \right] \quad (2)$$

where *dsp*, *mix*, *tru*, and *smp* stand for double-spike, mixture, true, and sample, respectively⁸. In equation (2), the symbol *XX* refers to either ^{40}Ca or ^{42}Ca . Equations (1) and (2) are solved iteratively, as described in Lehn et al. (2013).

Double-spike compositions were simulated using abundance data for ISOFLEX ^{43}Ca - and ^{48}Ca -enriched CaCO_3 powders (www.isoflex.com; Table 1) according to the equation:

$$^{43}\text{Ca}_{dsp} = \frac{^{43}\text{Ca}}{^{48}\text{Ca} + ^{43}\text{Ca}} \quad (3)$$

where the molar abundance of ^{43}Ca in the double-spike ($^{43}\text{Ca}_{dsp}$) ranges from 0 to 1. To calculate the composition of the spike-sample mixture, the simulated double-spike was combined with a sample having normal ratios¹¹ (Table 1) according to the equation:

$$p_{dsp} = \frac{\text{Ca}_{dsp}}{\text{Ca}_{dsp} + \text{Ca}_{smp}} \quad (4)$$

where the molar proportion of Ca from the double-spike (p_{dsp}) ranges from 0 to 1. It follows from equations (3) and (4) that,

$$\left(\frac{{}^{48}\text{Ca}}{{}^{43}\text{Ca}} \right)_{\text{dsp}} = \frac{1 - {}^{43}\text{Ca}_{\text{dsp}}}{{}^{43}\text{Ca}_{\text{dsp}}} \quad (5)$$

$$(\text{dsp} / \text{smp}) = \frac{p_{\text{dsp}}}{1 - p_{\text{dsp}}}, \quad (6)$$

where $({}^{48}\text{Ca}/{}^{43}\text{Ca})_{\text{dsp}}$ is the ${}^{48}\text{Ca}/{}^{43}\text{Ca}$ ratio of the double-spike and dsp/smp is the spike/sample ratio. An exponential law was used to simulate how mass-dependent fractionation changes the spike-sample mixture during ionization and measurement:

$$\left(\frac{{}^a\text{Ca}}{{}^b\text{Ca}} \right)_{\text{frac}} = \left(\frac{{}^a\text{Ca}}{{}^b\text{Ca}} \right)_{\text{frac}} \cdot \left(\frac{M_a}{M_b} \right)^\beta, \quad (7)$$

where *frac* and *frac* refer to fractionated and unfractionated, M_a and M_b are the exact atomic masses of Ca isotopes a and b, and β is the instrumental mass-fractionation factor. Both ${}^{43}\text{Ca}_{\text{dsp}}$ and p_{dsp} were stepped in 1% increments. We assumed a constant β of 0.5, based on laboratory observations. The model output is insensitive to the magnitude of β .

Because Ca isotopes have very different relative abundances and a wide mass spectrum, it is advantageous to implement a dynamic, multi-collection routine¹⁴. Consistent with Lehn et al. (2013), we adopted a three-hop duty cycle²⁰. A hop refers to a change in the magnet field setting, and the three hops combine to yield one duty cycle. The first hop measures ${}^{40}\text{Ca}$ and ${}^{43}\text{Ca}$ in the L2 and H3 cups, respectively. The second hop measures ${}^{43}\text{Ca}$ and ${}^{48}\text{Ca}$ in the L4 and H4 cups, respectively. The third hop measures ${}^{42}\text{Ca}$, ${}^{43}\text{Ca}$, and ${}^{44}\text{Ca}$ in the L1, center (C), and H1 cups, respectively. For real measurements, ${}^{41}\text{K}$ is collected in L3 during the third hop to monitor the contribution of ${}^{40}\text{K}$ to the ${}^{40}\text{Ca}$ signal. The model does not consider isobaric interferences because real data are not collected unless ${}^{40}\text{K}$ is negligible. The ${}^{40}\text{Ca}/{}^{43}\text{Ca}$ (or ${}^{42}\text{Ca}/{}^{43}\text{Ca}$), ${}^{48}\text{Ca}/{}^{43}\text{Ca}$, and ${}^{44}\text{Ca}/{}^{43}\text{Ca}$ ratios implemented in the data reduction are taken directly from the first, second, and third hops. Unique integration times were considered for each hop, namely 4.194, 8.388, 16.772 and 33.542 s, but the sum of integration times for the three hops was limited to 45 s to avoid significant fractionation differences between the first and last hop during real measurements. We considered 30, 60, 90, and 120 duty cycles per run.

1
2
3 125 The ^{40}Ca ion-beam was set to 20V to be consistent with the approach taken in Lehn et al. (2013).
4
5 126 Normally distributed errors were randomly assigned to the theoretically derived, instrumentally
6
7 127 fractionated ratios²⁰. Errors considered include those from counting statistics and Johnson noise^{20, 29}.
8
9 128 Ratios have variable errors because the errors depend on signal strength. Each pass through the three-hop
10
11 129 duty cycle ultimately generates one $^{40}\text{Ca}/^{44}\text{Ca}$ ratio and one $^{42}\text{Ca}/^{44}\text{Ca}$ ratio. The model reduces
12
13 130 hypothetical data for each duty cycle separately and averages multiple duty cycles to calculate a mean
14
15 131 $^{40}\text{Ca}/^{44}\text{Ca}$ ratio and a mean $^{42}\text{Ca}/^{44}\text{Ca}$ ratio for the entire run. The internal precision is determined by
16
17 132 doubling the standard error of the mean ratio ($2\sigma_{\text{SEM}}$). The model simulates each permutation 100 times
18
19 133 and averages the data to arrive at the representative mean ratio and $2\sigma_{\text{SEM}}$. On a 100×100 grid of $^{43}\text{Ca}_{\text{dsp}}$
20
21 134 and p_{dsp} values, the combination of hops, integration times, and number of duty cycles yields $>25,000,000$
22
23 135 possible permutations for each ratio, $^{40}\text{Ca}/^{44}\text{Ca}$ and $^{42}\text{Ca}/^{44}\text{Ca}$. To limit the rate of collector damage and
24
25 136 associated drift, we only considered methods delivering less than 6.5×10^{11} counts/sample to any
26
27 137 collector^{15, 20}. This number was derived from Holmden and Bélanger (2010), where drift was minimized
28
29 138 for sessions involving 30 or fewer runs. With these constraints, we focused on values for $^{43}\text{Ca}_{\text{dsp}}$ and p_{dsp}
30
31 139 yielding the optimal internal precision for $^{40}\text{Ca}/^{44}\text{Ca}$ and $^{42}\text{Ca}/^{44}\text{Ca}$ ratios, while minimizing analysis time
32
33 140 and collector damage.
34
35
36
37
38

39 141 *2.1 Ca Isotope Measurements*

40
41 142 We verified the optimization by repeatedly measuring common standards, such as OSIL Atlantic
42
43 143 Seawater (ASW), NIST SRM 915a, NIST SRM 915b, and USGS BHVO-1. The standards 915a and 915b
44
45 144 are calcium carbonate powders, and BHVO-1 is a basaltic rock. We also analyzed Saskatchewan Isotope
46
47 145 Lab gravimetric calcium fluoride (CaF_2). Measurements were made in the Radiogenic Isotope Clean
48
49 146 Laboratory at Northwestern University using a Thermo Fisher Triton MC-TIMS. The optimal ^{48}Ca - ^{43}Ca
50
51 147 double-spike was prepared using ISOFLEX $^{48}\text{CaCO}_3$ and $^{43}\text{CaCO}_3$ powders (Table 1) and calibrated
52
53 148 against the Saskatchewan Isotope Lab CaF_2 standard¹⁴. Table 1 provides the measured composition of the
54
55 149 Northwestern University (NU) double-spike.
56
57
58
59
60

1
2
3 150 Spike-sample equilibration, purifications, filament loading, and the drift correction follow the
4
5 151 protocols provided in Lehn et al. (2013)²⁰. Table 2 summarizes instrumental parameters specific to the
6
7 152 ⁴⁸Ca-⁴³Ca method. In brief, aliquots containing 50 µg of Ca were spiked and refluxed in acid-cleaned
8
9
10 153 Teflon vials for several hours to ensure complete equilibration. Spiked aliquots of ASW, CaF₂, and
11
12 154 BHVO-1 were purified by ion-exchange chromatography, using Teflon columns packed with Bio-Rad
13
14 155 AG MP-50 cation exchange resin and ultrapure HCl eluents. The purified Ca solutions were dried and
15
16 156 treated with 30% H₂O₂ to oxidize any organic compounds that might have leached from the resin. Spiked
17
18 157 aliquots of 915a and 915b were not purified. All solutions were dried and treated with two drops of
19
20 158 ultrapure HNO₃, to convert Ca to nitrate form. Single Ta filaments were outgassed under vacuum for 1.5
21
22 159 hours at 4800 mA, which is over 1500 mA higher than running conditions in the TIMS. Approximately
23
24 160 10-16 µg of Ca was loaded onto the outgassed filaments with 0.5 mg of 10% H₃PO₄ (Sigma-Aldrich
25
26 161 *TraceSELECT*® Ultra H₃PO₄) and dried. The Ca was loaded between parafilm “dams” to prevent
27
28 162 spreading on the surface of the filament. We also measured the ⁴⁰Ca/⁴⁴Ca ratio of BHVO-1 using the Lehn
29
30 163 et al. (2013) ⁴³Ca-⁴²Ca double-spike method²⁰. Unlike the ⁴⁸Ca-⁴³Ca double-spike method, the third hop of
31
32 164 the ⁴³Ca-⁴²Ca double-spike method requires zoom optics. For all measurements, peak shapes were
33
34 165 examined to confirm the absence of potential backscatter and reflections. All runs were checked for
35
36 166 isobaric interferences from K, and occasionally, Ti and doubly-charged Sr¹⁰. No interferences were
37
38 167 observed. We quantified our total procedural blank several times (n = 6) using isotope-dilution. The 500:1
39
40 168 sample-to-blank ratio is negligible. All ratios and their uncertainties are reported in delta notation, given
41
42 169 in per mil, relative to OSIL Atlantic seawater (sw):

$$\delta^{44/XX}\text{Ca}_{\text{sw}} \text{ (in } \text{‰}) = \left[\frac{\left(\frac{{}^{44}\text{Ca}}{\text{XX}\text{Ca}} \right)_{\text{sample}}}{\left(\frac{{}^{44}\text{Ca}}{\text{XX}\text{Ca}} \right)_{\text{sw}}} - 1 \right] \cdot 1000, \quad (8)$$

46
47
48
49 170
50
51
52
53
54 171 where *XX* refers to either ⁴⁰Ca or ⁴²Ca.

55
56
57 172 **3. Results**
58
59
60

1
2
3 173 Similar to optimization of the ^{43}Ca - ^{42}Ca double-spike, the Monte Carlo simulations for the ^{48}Ca -
4
5 174 ^{43}Ca double-spike yielded a wide range of non-unique solutions²⁰. To minimize analysis time and
6
7
8 175 collector damage at a marginal expense to precision, we chose a method having 90 duty cycles with
9
10 176 integration times of 4.194, 8.388, and 16.776 s for the first, second, and third hops, respectively. This
11
12 177 method requires 2.5 hours/sample, including filament warm-up. Figures 1a and 1b show contour maps of
13
14 178 the predicted internal precisions for $\delta^{44/40}\text{Ca}$ and $\delta^{44/42}\text{Ca}$. Statistically identical internal precisions are
15
16 179 possible for a wide range of double-spike and spike/sample ratios. Internal precision is more sensitive to
17
18 180 p_{dsp} than $^{43}\text{Ca}_{dsp}$. We ultimately selected $^{43}\text{Ca}_{dsp} = 0.40$ and $p_{dsp} = 0.40$, as these values offer the best
19
20 181 precision for both $\delta^{44/40}\text{Ca}$ and $\delta^{44/42}\text{Ca}$ during a single mass spectrometric run. For $\delta^{44/40}\text{Ca}$, the predicted
21
22 182 internal precision ($2\sigma_{SEM}$) is $\pm 0.015 - 0.020\text{‰}$, and for $\delta^{44/42}\text{Ca}$, the predicted internal precision is $0.025 -$
23
24 183 0.030‰ .

25
26
27 184 Table 3 and Figure 2 summarize the standard measurements. For the 81 measurements made with
28
29 185 the ^{48}Ca - ^{43}Ca double-spike, the internal precision was more variable and on average, slightly higher than
30
31 186 the model predictions. For $\delta^{44/40}\text{Ca}$, the measured internal precision ranged from 0.013 to 0.081‰, with an
32
33 187 average of 0.025‰. For $\delta^{44/42}\text{Ca}$, the measured internal precision ranged from 0.020 to 0.093‰, with an
34
35 188 average of 0.034‰. The external reproducibility ($2\sigma_{SD}$) for $\delta^{44/40}\text{Ca}$ ranged from 0.201 to 0.253‰, with an
36
37 189 average of 0.223‰, which is ~11 times higher than the average internal precision. The external
38
39 190 reproducibility for $\delta^{44/42}\text{Ca}$ ranged from 0.091 to 0.153‰, with an average of 0.126‰, which is ~4 times
40
41 191 higher than internal precision. By comparison, the 6 measurements of BHVO-1 made with the ^{43}Ca - ^{42}Ca
42
43 192 double-spike yielded results consistent with findings reported in Lehn et al. (2013). The external
44
45 193 reproducibility was 0.033‰, which is only ~1.5 times higher than the average internal precision of
46
47 194 0.023‰.

51 195 4. Discussion

52 196 4.1 Filament Reservoir Mixing

53 197 Poor external reproducibility for the ^{48}Ca - ^{43}Ca method points to an additional source of error
54
55 198 beyond counting statistics and Johnson noise. When examining uncorrected and fractionation-corrected
56
57
58
59
60

1
2
3 199 ratios within a given run, remnant fractionation patterns were observed for nearly all analyses. Some
4
5 200 trends in the corrected ratios were linear, while others were curved or had multiple inflections. Figure 3
6
7 201 shows examples for one “good” run and two “bad” runs. For the good run, the uncorrected ratios show a
8
9 202 linear fractionation trend that is eliminated after the fractionation correction and spike un-mixing. The bad
10
11 203 runs show irregular fractionation trends that persist after the fractionation correction and spike un-mixing.
12
13 204 Such patterns were not observed, or were minor, during development of the Lehn et al. (2013) method,
14
15 205 and nor did we observe any remnant fractionation patterns during application of the ^{43}Ca - ^{42}Ca double-
16
17 206 spike in the present study. Because the ^{48}Ca - ^{43}Ca and ^{43}Ca - ^{42}Ca double-spike methods employed the same
18
19 207 chemical separation, filament loading, filament heating, and drift correction protocols, we can eliminate
20
21 208 these steps as possible sources of error. The patterns are consistent with deviation from ideal exponential
22
23 209 mass-fractionation, caused by mixing of multiple, independently fractionating reservoirs on the filament¹⁷,
24
25 210 ^{18, 30}. Both double-spike methods experience this problem, but the effects are more evident and
26
27 211 detrimental for the ^{48}Ca - ^{43}Ca method according to the “average mass rule,” which states that inaccuracies
28
29 212 in fractionation-corrected data are greater for isotope ratios having an average mass further away from the
30
31 213 average mass of the normalizing ratio^{21, 31-33}.

32
33
34
35
36 214 To examine this idea in more detail, we used the filament reservoir-mixing model presented in
37
38 215 Fantle and Bullen (2009)¹⁷ to generate theoretical, instrumentally fractionated Ca isotope ratios that were
39
40 216 then reduced with the double-spike equations (Equations 1 and 2). The model simulates two pools of a
41
42 217 spike-sample mixture that have the same starting composition but fractionate independently. Calcium
43
44 218 isotope ratios in each pool are fractionated with the exponential law (Equation 7). However, β is now
45
46 219 replaced with a time dependent value, β_t :

$$\beta_t = \beta_0 + \Delta\beta \cdot x, \quad (9)$$

47
48
49 220 where β_0 represents the starting fractionation factor based on modification during filament warm up and
50
51 221 focusing, $\Delta\beta$ is rate of change for the fractionation factor as a function of duty cycle, and x is the duty
52
53 222 cycle number. For each duty cycle, the percentage contribution of each pool to the total ion beam varies.
54
55
56
57
58
59
60

1
2
3 224 We considered two scenarios and a control summarized in Table 4. The scenarios were designed to
4
5 225 simulate as closely as possible the filament mixing patterns shown in Figure 3. In the first scenario, one
6
7 226 pool is the sole source of Ca ions for the first several duty cycles. Later, the other pool begins
8
9 227 fractionating and delivers increasingly more Ca ions until the first pool completely decays. In the second
10
11 228 scenario, the second pool begins fractionating during the first duty cycle and starts supplying Ca ions
12
13 229 during the following duty cycle. The transition between the two pools occurs more slowly, such that they
14
15 230 contribute equally by the end of the run. The control scenario simulates ideal fractionation from a single
16
17 231 pool. The scope of our analysis includes the measurement of $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with ^{48}Ca - ^{43}Ca , ^{43}Ca - ^{42}Ca ,
18
19 232 ^{48}Ca - ^{42}Ca , and ^{46}Ca - ^{43}Ca double-spikes and the measurement of $^{42}\text{Ca}/^{44}\text{Ca}$ ratios with ^{48}Ca - ^{43}Ca and ^{46}Ca -
20
21 233 ^{43}Ca double-spikes.

22
23
24 234 Figure 4 illustrates the model output by plotting uncorrected ratios and simulated measurement
25
26 235 errors (deviations) as a function of duty cycle number. Deviations, given in ppm/amu, are calculated as
27
28 236 the difference between the true isotope ratio and the final output ratio after processing through the double-
29
30 237 spike equations. Figures 4A and C demonstrate that simple reservoir mixing can reasonably simulate the
31
32 238 actual uncorrected data provided in Figures 3C and E, while Figures 4B and D show that error magnitudes
33
34 239 depend on the combination of double-spike and sample ratios. Analysis of $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{43}Ca -
35
36 240 ^{42}Ca double-spike yields the lowest errors, followed by analysis of $^{42}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{46}Ca - ^{43}Ca
37
38 241 double-spike. Analysis of $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{46}Ca - ^{43}Ca double-spike, $^{42}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{48}Ca - ^{43}Ca
39
40 242 double-spike, and $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{48}Ca - ^{42}Ca double-spike all yield higher errors having a similar
41
42 243 magnitude. Finally, analysis of $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{48}Ca - ^{43}Ca double-spike produces the highest errors.

43
44
45 244 To further examine these relationships, Figure 5 plots the average error for the various double-
46
47 245 spike methods as a function of the average error for the measurement of $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{43}Ca - ^{42}Ca
48
49 246 double-spike. The model scenarios define remarkably linear correlations. However, deeper consideration
50
51 247 reveals that the correlations are insensitive to the details of the mixing model, such as the parameters in
52
53 248 equation (9) or the percentage contribution of the two pools to the total ion beam. Other mixing
54
55 249 simulations produce identical correlations. Rather, the results are consistent with predictions from the
56
57
58
59
60

1
2
3 250 average mass rule. For example, the average mass difference between the ^{43}Ca - ^{42}Ca pair and the ^{40}Ca - ^{44}Ca
4 pair is 0.5 amu. The average mass difference between the ^{48}Ca - ^{43}Ca pair and the ^{40}Ca - ^{44}Ca pair is 3.5 amu.
5 251
6 The ratio of 3.5 amu to 0.5 amu equals 7, which is effectively identical to the slope of 6.9 shown in Figure
7 252
8
9 253 5. In other words, the average mass rule predicts that errors attributable to filament reservoir mixing will
10
11 254 always be roughly 7 times larger when analyzing $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{48}Ca - ^{43}Ca double-spike compared
12
13 255 to a ^{43}Ca - ^{42}Ca double-spike. Real measurements are consistent with this assertion. As reported in Section
14
15 256 3, analysis of $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with the ^{48}Ca - ^{43}Ca double-spike yields an external reproducibility of
16
17 257 0.22‰, whereas analysis of $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with the ^{43}Ca - ^{42}Ca double-spike yields an external
18
19 258 reproducibility of 0.04‰²⁰. The ratio of 0.22‰ to 0.04‰ equals 5.5. Most likely, the theoretical ratio (7)
20
21 259 differs from the actual ratio (5.5) because the former only considers the error from filament reservoir
22
23 260 mixing, whereas the latter also includes errors for counting statistics and Johnson noise, which differ
24
25 261 between the two methods.
26
27

28
29 262 According to the average mass rule, errors for the analysis of $^{42}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{46}Ca - ^{43}Ca
30
31 263 double-spike, $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{46}Ca - ^{43}Ca double-spike, $^{42}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{48}Ca - ^{43}Ca double-
32
33 264 spike, and $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{48}Ca - ^{42}Ca double-spike will be approximately 3, 5, 5, and 6 times worse,
34
35 265 respectively, than errors for the analysis of $^{40}\text{Ca}/^{44}\text{Ca}$ ratios with a ^{43}Ca - ^{42}Ca double-spike. As shown in
36
37 266 Figure 5, the corresponding slopes are 4.0, 5.4, 5.4, and 5.5. We note that the slopes provided in Figure 5
38
39 267 slightly differ from the predictions offered by the average mass rule (e.g., 6.9 versus 7, 4 versus 3, etc.).
40
41 268 This is a second-order effect reflecting the double-spike and spike/sample ratios applied in the data
42
43 269 reduction algorithm (see Table 4). Modestly adjusting the optimal ratios can entirely eliminate the
44
45 270 discrepancies such that the average mass law predictions and mixing model slopes are identical.
46
47

48
49 271 Our findings surrounding the average mass rule lead to important conclusions and may offer new
50
51 272 directions for improving the precision of Ca isotope measurements by MC-TIMS. It appears that efforts
52
53 273 to optimize Ca isotope ratios by double-spike MC-TIMS should consider errors associated with filament
54
55 274 reservoir mixing in addition to those for counting statistics and Johnson noise^{20, 29, 34}. In fact, we suggest
56
57 275 that the choice of the double-spike pair relative to the sample ratio of interest should be considered before
58
59
60

1
2
3 276 any attempt to optimize the double-spike and spike/sample ratios. While selecting a double-spike pair
4
5 277 with a large mass difference may enable better quantification of the instrumental fractional factor⁹, this
6
7 278 seems secondary to constraints imposed by the average mass rule. As demonstrated here, filament
8
9 279 reservoir mixing explains why the ⁴⁸Ca-⁴³Ca double-spike method consistently underperforms predictions
10
11 280 from the optimization model. The present study and Lehn et al. (2013) provide compelling evidence that a
12
13 281 ⁴³Ca-⁴²Ca double-spike offers the most precise analysis of ⁴⁰Ca/⁴⁴Ca ratios²⁰. We go on to propose that a
14
15 282 carefully optimized ⁴⁶Ca-⁴³Ca double-spike method should offer the most precise analysis of ⁴²Ca/⁴⁴Ca
16
17 283 ratios, as the errors due to reservoir mixing will be smaller than those for the ⁴⁸Ca-⁴³Ca double-spike.
18
19 284 However, these recommendations could change if techniques can be developed to limit reservoir mixing.
20
21 285 Building on methods presented in Holmden and Bélanger (2010), Lehn et al. (2013) hypothesized that
22
23 286 loading samples between parafilm dams would mitigate reservoir mixing on the filament²⁰, but the current
24
25 287 study indicates that parafilm dams do not eliminate the problem. It is uncertain if reservoir mixing would
26
27 288 be more prevalent without the dams, but we note that another study employing a ⁴⁶Ca-⁴³Ca double-spike
28
29 289 found little difference with and without dams³⁵. We load relatively large amounts of Ca on the filament
30
31 290 (10 – 16 µg) to optimize counting statistics and because the ionization efficiency of Ca is notoriously
32
33 291 poor. However, reservoir effects are likely more prevalent for large Ca pools. Using double and triple
34
35 292 filament assemblies^{25, 36-38} can enable smaller loads, improve ionization efficiency, slow the fractionation
36
37 293 rate, and smooth fractionation patterns, all of which have potential to reduce reservoir mixing errors.
38
39 294 Finally, selecting a double-spike pair with the same average mass as the sample ratio pair could in theory
40
41 295 eliminate the effect entirely. Examples include using ⁴⁶Ca-⁴⁰Ca, ⁴⁶Ca-⁴²Ca, or ⁴⁶Ca-⁴⁴Ca double-spikes to
42
43 296 measure ⁴²Ca/⁴⁴Ca, ⁴⁰Ca/⁴⁸Ca, and ⁴²Ca/⁴⁸Ca ratios, respectively. However, the natural abundances of Ca
44
45 297 are highly unfavorable for such applications and introduce other limitations to high-precision
46
47 298 measurements.

53 299 *4.2 ⁴⁰Ca enrichments*

55 300 Recent studies have highlighted potential problems with common Ca isotope standards.

57 301 Controversy has developed around possible ⁴⁰Ca enrichments in 915a and/or seawater. Some studies have

1
2
3 302 reported mass-independent isotope enrichments in 915a^{24, 27, 39}; others have interpreted small
4
5 303 discrepancies in ⁴⁰Ca/⁴⁴Ca and ⁴²Ca/⁴⁴Ca measurements as a ⁴⁰Ca enrichment in seawater^{16, 23}; and yet
6
7 304 others have found no enrichment in either standard^{1, 25, 26, 40}. While significant differences in external
8
9 305 reproducibility exist, we find that the ⁴³Ca-⁴²Ca and ⁴⁸Ca-⁴³Ca double-spike methods yield the same
10
11 306 $\delta^{44/40}\text{Ca}$ values for 915a and ASW²⁰, as well as the three other standards analyzed. As Figure 6
12
13 307 demonstrates, a plot of $\delta^{44/40}\text{Ca}$ versus $\delta^{44/42}\text{Ca}$ does not reveal significant ⁴⁰Ca enrichments.
14
15

16 308 *4.3 Ca isotope error reporting*

17
18 309 Consistent with the convention in radiogenic isotope geochemistry⁴¹, we report external
19
20 310 reproducibility as twice the standard deviation ($2\sigma_{\text{SD}}$) of repeated standard measurements. A recent study
21
22 311 (Fantle and Tipper, 2014)¹⁶ has argued that twice the standard error of the mean ($2\sigma_{\text{SEM}}$), based on
23
24 312 multiple analyses of samples, better represents the external reproducibility of Ca isotope measurements
25
26 313 because standards often have simple matrices that obviate the need for column chemistry and real samples
27
28 314 have more complex matrices that can translate into outlier measurements, despite purification by ion-
29
30 315 exchange chromatography. However, our data tentatively suggest that filament reservoir effects can fully
31
32 316 explain outlier measurements. As shown in Table 3, matrix-rich standards purified by ion-exchange
33
34 317 chromatography (ASW and BHVO-1), as well as matrix-poor standards purified by ion-exchange
35
36 318 chromatography (CaF₂), can be measured within the same external reproducibility as matrix-poor
37
38 319 standards loaded directly onto filaments (915a and 915b). This holds true for both the ⁴³Ca-⁴²Ca and ⁴⁸Ca-
39
40 320 ⁴³Ca double-spike techniques. These results indicate that the external reproducibility determined from the
41
42 321 repeated analysis of standards can be assigned as the uncertainty for a single sample analysis. However,
43
44 322 additional experiments specifically targeting matrix effects are required to confirm this hypothesis.
45
46
47

48 323 **5. Conclusions**

49
50 324 Building on previous research by Lehn et al. (2013), this study used a Monte Carlo error model to
51
52 325 optimize a ⁴⁸Ca-⁴³Ca double-spike for measuring $\delta^{44/40}\text{Ca}$ and $\delta^{44/42}\text{Ca}$ values by MC-TIMS. We tested the
53
54 326 optimization by repeatedly analyzing common Ca isotopes standards, such as OSIL Atlantic seawater,
55
56 327 NIST SRM 915a, NIST SRM 915b, USGS BHVO-1, and CaF₂. For $\delta^{44/40}\text{Ca}$, the average external
57
58
59
60

1
2
3 328 reproducibility ($2\sigma_{SD}$) of $\pm 0.223\%$ was ~ 11 times higher than internal precision ($2\sigma_{SEM}$). For $\delta^{44/42}\text{Ca}$, the
4
5 329 average external reproducibility of $\pm 0.126\%$ was ~ 4 times higher than internal precision. By comparison,
6
7 330 $\delta^{44/40}\text{Ca}$ values measured with a ^{43}Ca - ^{42}Ca double-spike had an external reproducibility of $\pm 0.04\%$, which
8
9 331 is only ~ 2 times higher than internal precision (Lehn et al, 2013). We attribute the poor performance of
10
11 332 the ^{48}Ca - ^{43}Ca double-spike to filament reservoir effects, which cause deviations from ideal exponential
12
13 333 mass-fractionation during ionization. We employed a mixing model to examine why errors are either
14
15 334 amplified or diminished for methods involving different combinations of double-spike and sample ratios.
16
17 335 We found that the difference between the average mass of the double-spike pair and the sample ratio pair
18
19 336 can be used to predict the extent of error magnification. These findings suggest that the “average mass
20
21 337 rule” should be taken into consideration when selecting a Ca isotope double-spike method. Because the
22
23 338 average mass difference between the ^{43}Ca - ^{42}Ca pair and the ^{40}Ca - ^{44}Ca pair is only 0.5 amu, the ^{43}Ca - ^{42}Ca
24
25 339 double-spike technique should provide the most precise analysis of $\delta^{44/40}\text{Ca}$ values. Limiting the effects of
26
27 340 filament reservoir mixing may further improve precision. Finally, we note that many of the problems that
28
29 341 hamper the quality of Ca isotope measurements by MC-TIMS, such as collector cup drift^{20, 35} and filament
30
31 342 reservoir mixing¹⁷, are not particularly unusual^{18, 32, 33}. Carefully monitoring and accounting for these
32
33 343 factors largely eliminates the need to repeatedly analyze the same sample, although it remains good
34
35 344 practice to occasionally measure duplicates. With the ^{43}Ca - ^{42}Ca double-spike method, our data suggest
36
37 345 that the external reproducibility determined from the repeated measurement of standards could accurately
38
39 346 represent the uncertainty assigned to a single sample measurement, but we note that additional
40
41 347 experiments are required to fully test this idea.

46 348 6. Acknowledgements

47
48 349 The authors thank Chris Holmden for providing Saskatchewan Isotope Lab CaF_2 and for many thoughtful
49
50 350 discussions that aided and improved this study. Insightful comments from four anonymous reviewers
51
52 351 improved the manuscript. This work was supported by an Environmental Protection Agency STAR
53
54 352 Fellowship awarded to G.O.L., as well as a Major Research Instrumentation Grant (EAR-0723151) and a
55
56
57
58
59
60

353 David and Lucile Packard Foundation Fellowship awarded to A.D.J. The study also benefited from NSF
354 EAR-0643317 and NSF-0617585 awarded to A.D.J.

355 7. References

- 356 1. S. Huang, J. Farkaš and S. B. Jacobsen, *Earth and Planetary Science Letters*, 2010, **292**, 337-344.
- 357 2. L. M. Reynard, G. M. Henderson and R. E. M. Hedges, *Journal of Archaeological Science*, 2011,
358 **38**, 657-664.
- 359 3. A. D. Melin, B. E. Crowley, S. T. Brown, P. V. Wheatley, G. L. Moritz, F. T. Yit Yu, H. Bernard,
360 D. J. DePaolo, A. D. Jacobson and N. J. Dominy, *American journal of physical anthropology*,
361 2014, DOI: 10.1002/ajpa.22530.
- 362 4. N.-C. Chu, G. M. Henderson, N. S. Belshaw and R. E. M. Hedges, *Applied Geochemistry*, 2006,
363 **21**, 1656-1667.
- 364 5. C. Holmden, K. Panchuk and S. C. Finney, *Geochimica Et Cosmochimica Acta*, 2012, **98**, 94-
365 106.
- 366 6. C. Holmden, D. A. Papanastassiou, P. Blanchon and S. Evans, *Geochimica Et Cosmochimica*
367 *Acta*, 2012, **83**, 179-194.
- 368 7. M. Dodson, *Geochimica et Cosmochimica Acta*, 1970, **34**, 1241-1244.
- 369 8. O. Eugster, F. Tera and G. J. Wasserburg, *J. Geophys. Res.*, 1969, **74**, 3897-3908.
- 370 9. D. J. DePaolo, *Reviews in mineralogy and geochemistry*, 2004, **55**, 255-288.
- 371 10. A. Heuser, A. Eisenhauer, N. Gussone, B. Bock, B. T. Hansen and T. F. Nagler, *International*
372 *Journal of Mass Spectrometry*, 2002, **220**, 385-397.
- 373 11. W. A. Russell, D. A. Papanastassiou and T. A. Tombrello, *Geochimica et Cosmochimica Acta*,
374 1978, **42**, 1075-1090.
- 375 12. T. F. Nagler, A. Eisenhauer, A. Muller, C. Hemleben and J. Kramers, *Geochemistry Geophysics*
376 *Geosystems*, 2000, **1**.
- 377 13. I. R. Fletcher, A. L. Maggi, K. J. R. Rosman and N. J. McNaughton, *International Journal of*
378 *Mass Spectrometry and Ion Processes*, 1997, **163**, 1-17.
- 379 14. C. Holmden, *Measurement of $\delta^{44}\text{Ca}$ Using a ^{43}Ca - ^{42}Ca Double-spike TIMS*, Saskatchewan
380 Geological Survey, 2005.
- 381 15. C. Holmden and N. Bélangier, *Geochimica et Cosmochimica Acta*, 2010, **74**, 995-1015.
- 382 16. M. S. Fantle and E. T. Tipper, *Earth-Science Reviews*, 2014, **129**, 148-177.
- 383 17. M. Fantle and T. Bullen, *Chemical Geology*, 2009, **258**, 50-64.
- 384 18. S. R. Hart and A. Zindler, *International Journal of Mass Spectrometry and Ion Processes*, 1989,
385 **89**, 287-301.
- 386 19. R. S. Hindshaw, B. C. Reynolds, J. G. Wiederhold, R. Kretzschmar and B. Bourdon, *Geochimica*
387 *et Cosmochimica Acta*, 2011, **75**, 106-118.
- 388 20. G. O. Lehn, A. D. Jacobson and C. Holmden, *International Journal of Mass Spectrometry*, 2013,
389 **351**, 69-75.
- 390 21. K. Gopalan, D. Macdougall and C. Macisaac, *International Journal of Mass Spectrometry*, 2006,
391 **248**, 9-16.
- 392 22. S. Huang, J. Farkaš and S. B. Jacobsen, *Geochimica et Cosmochimica Acta*, 2011, **75**, 4987-4997.
- 393 23. J.-S. Ryu, A. D. Jacobson, C. Holmden, C. Lundstrom and Z. Zhang, *Geochimica et*
394 *Cosmochimica Acta*, 2011, **75**, 6004-6026.
- 395 24. J. I. Simon and D. J. DePaolo, *Earth and Planetary Science Letters*, 2010, **289**, 457-466.
- 396 25. J. Farkaš, A. Déjeant, M. Novák and S. B. Jacobsen, *Geochimica et Cosmochimica Acta*, 2011,
397 **75**, 7031-7046.
- 398 26. G. Caro, D. A. Papanastassiou and G. J. Wasserburg, *Earth and Planetary Science Letters*, 2010,
399 **296**, 124-132.
- 400 27. J. I. Simon, D. J. DePaolo and F. Moynier, *The Astrophysical Journal*, 2009, **702**, 707.

- 1
2
3 401 28. J. F. Rudge, B. C. Reynolds and B. Bourdon, *Chemical Geology*, 2009, **265**, 420-431.
4 402 29. S. G. John and J. F. Adkins, *Marine Chemistry*, 2010, **119**, 65-76.
5 403 30. D. Upadhyay, E. E. Scherer and K. Mezger, *Journal of Analytical Atomic Spectrometry*, 2008, **23**,
6 404 561.
7 405 31. G. J. Wasserburg, S. B. Jacobsen, D. J. DePaolo, M. T. McCulloch and T. Wen, *Geochimica et*
8 406 *Cosmochimica Acta*, 1981, **45**, 2311-2323.
9 407 32. D. Vance and M. Thirlwall, *Chemical Geology*, 2002, **185**, 227-240.
10 408 33. R. Andreasen and M. Sharma, *International Journal of Mass Spectrometry*, 2009, **285**, 49-57.
11 409 34. S. G. John, *Journal of Analytical Atomic Spectrometry*, 2012, **27**, 2123-2131.
12 410 35. R. S. Hindshaw, PhD, ETH, 2011.
13 411 36. L. C. Nielsen and D. J. DePaolo, *Geochimica et Cosmochimica Acta*, 2013, **118**, 276-294.
14 412 37. E. J. Catanzaro, *Journal of Geophysical Research*, 1967, **72**, 1325-1327.
15 413 38. L. Moore, E. Heald and J. Filliben, 1978.
16 414 39. M. Schiller, C. Paton and M. Bizzarro, *Journal of Analytical Atomic Spectrometry*, 2012, **27**, 38.
17 415 40. M. Amini, A. Eisenhauer, F. Böhm, C. Holmden, K. Kreissig, F. Hauff and K. P. Jochum,
18 416 *Geostandards and Geoanalytical Research*, 2009, **33**, 231-247.
19 417 41. P. Deines, S. L. Goldstein, E. H. Oelkers, R. L. Rudnick and L. M. Walter, *Chemical Geology*,
20 418 2003, **202**, 1-4.
21
22
23
24
25
26
27
28

423 Table Captions

29 424 **Table 1:** Isotopic abundances of Ca normal, Ca single spikes, and Northwestern University (NU) double-
30
31 425 spikes

32
33 426 **Table 2:** Method settings for Triton MC-TIMS

34
35 427 **Table 3:** Ca isotope results for 5 standards using the ^{48}Ca - ^{43}Ca and ^{43}Ca - ^{42}Ca double-spike techniques

36
37 428 **Table 4:** Input parameters for filament reservoir mixing model
38
39
40
41

429 430 Figure Captions

431 **Figure 1:** Contour plots of theoretical (A) $\delta^{44/40}\text{Ca}$ and (B) $\delta^{44/42}\text{Ca}$ internal precisions ($2\sigma_{\text{SEM}}$) using a
432 ^{48}Ca - ^{43}Ca double-spike. The model output reflects a 20 V ^{40}Ca beam and 90 duty cycles with integration
433 times of 4.194, 8.388, and 16.766 s for hops 1, 2, and 3, respectively. The curves represent theoretical
434 thresholds for collector damage (6.5×10^{11} counts/sample) as a function of $^{43}\text{Ca}_{\text{dsp}}$ and p_{dsp} for collectors
435 L4, L1, C, H1, H3, and H4. The filled circle marks the optimal ^{48}Ca - ^{43}Ca double-spike composition tested
436 in this study ($^{43}\text{Ca}_{\text{dsp}} = 40\%$ and $p_{\text{dsp}} = 40\%$).
437

1
2
3 438 **Figure 2:** $\delta^{44/40}\text{Ca}$ and $\delta^{44/42}\text{Ca}$ values for ASW, 915a, 915b, CaF_2 and BHVO-1. Circles show data
4
5 439 measured with the ^{48}Ca - ^{43}Ca double-spike method. Diamonds show $\delta^{44/40}\text{Ca}$ values for BHVO-1
6
7 440 measured with the ^{43}Ca - ^{42}Ca double-spike method. The dotted lines denote average values. The gray areas
8
9 441 display external reproducibility ($2\sigma_{\text{SD}}$).
10

11 442
12
13 443 **Figure 3:** Fractionation patterns observed in uncorrected and corrected data for 915b obtained with the
14
15 444 ^{48}Ca - ^{43}Ca double-spike method. Panels A and B show data with minimal fractionation patterns. Panels C
16
17 445 – D and E – F provide different examples of filament reservoir effects. The dashed lines in Panels B, D,
18
19 446 and F display the mean 915b value for this study (Table 3).
20
21 447
22

23 448 **Figure 4:** Uncorrected and corrected data from the filament reservoir mixing model. Panels A and B
24
25 449 show data for scenario 1, and Panels C and D show data for scenario 2 (see Table 4). Mixing scenarios 1
26
27 450 and 2 were designed to simulate real data from Figure 3. Note that results for $\delta^{44/40}\text{Ca}$ measured with a
28
29 451 ^{48}Ca - ^{42}Ca double-spike, $\delta^{44/40}\text{Ca}$ measured with a ^{46}Ca - ^{43}Ca double-spike, and $\delta^{44/42}\text{Ca}$ measured with a
30
31 452 ^{48}Ca - ^{43}Ca double-spike overlap.
32
33
34
35

36 453
37
38 454 **Figure 5:** Summary of filament reservoir mixing model results. Errors for $\delta^{44/4X}\text{Ca}$ measured with
39
40 455 different double-spikes are plotted against errors for $\delta^{44/40}\text{Ca}$ measured with a ^{43}Ca - ^{42}Ca double-spike.
41
42 456 Note that three sets of results overlap (see caption for Figure 4). The slopes demonstrate how errors
43
44 457 propagate according to predictions from the average mass rule. Other parameters, such as the double-
45
46 458 spike and spike/sample ratios, also influence the magnitude of errors, but to a lesser degree than the
47
48 459 choice of the double-spike pair relative to the sample ratio of interest. See text.
49
50
51

52 460
53 461 **Figure 6:** $\delta^{44/42}\text{Ca}$ versus $\delta^{44/40}\text{Ca}$ for five standards analyzed with a ^{48}Ca - ^{43}Ca double-spike. Error bars
54
55 462 represent $2\sigma_{\text{SEM}}$ reported in Table 3. The solid line shows theoretical kinetic mass fractionation (slope =
56
57 463 0.488). The dashed line shows theoretical equilibrium mass fractionation (slope = 0.476). Both lines are
58
59
60

1
2
3 464 anchored to zero on the seawater scale, although this does not convey less uncertainty for ASW. Both
4

5 465 lines provide an acceptable fit to the data within the uncertainties.
6

7
8 466
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

Table 1

Isotope	Ca normal ^a	⁴⁸ CaCO ₃ ^b	⁴³ CaCO ₃ ^b	⁴⁸ Ca- ⁴³ Ca double-spike ^c	⁴³ Ca- ⁴² Ca double-spike ^d
⁴⁰ Ca	0.9698	0.0272	0.0540	0.0396	0.0435
⁴² Ca	0.0064	0.0002	0.0018	0.0009	0.4678
⁴³ Ca	0.0013	0.0001	0.90 ± 0.006	0.3626	0.4630
⁴⁴ Ca	0.0206	0.0015	0.0444	0.0191	0.0256
⁴⁶ Ca	0.0000	<0.0001	<0.0001	0.0002	0.0000
⁴⁸ Ca	0.0018	0.971 ± 0.002	0.0008	0.5775	0.0002

a: Russell et al. (1978)

b: Isoflex.com

c: Measured in this study, see text

d: Lehn et al. (2013)

Table 2

Method Parameter	Setting
⁴³ Ca abundance in ⁴⁸ Ca- ⁴³ Ca double-spike	$^{43}\text{Ca}_{\text{dsp}} = ^{43}\text{Ca} / (^{48}\text{Ca} + ^{43}\text{Ca}) = 0.40 \text{ mol/mol}$
Double-spike ratio	$(^{48}\text{Ca}/^{43}\text{Ca})_{\text{dsp}} = (1 - ^{43}\text{Ca}_{\text{dsp}}) / ^{43}\text{Ca}_{\text{dsp}} = 1.5 \text{ mol/mol}$
Proportion of spike Ca in spike-sample mixture	$p_{\text{dsp}} = \text{Ca}_{\text{dsp}} / (\text{Ca}_{\text{dsp}} + \text{Ca}_{\text{smp}}) = 0.40 \text{ mol/mol}$
Spike-sample ratio	$\text{dsp/smp} = p_{\text{dsp}} / (1 - p_{\text{dsp}}) = 0.66 \text{ mol/mol}$
Filament assembly	Single Ta (0.002" h x 0.030" w)
Ca load	10 - 16 µg
⁴⁰ Ca ion beam intensity	20 V
Accelerating voltage	10 kV
Resistors	10 ¹¹ Ω
Source Vacuum	< 1 x 10 ⁻⁷ mbar
Analyzer Vacuum	< 6 x 10 ⁻⁹ mbar
Hop 1	⁴⁰ Ca (L2), ⁴³ Ca (H3), 4.194 s
Hop 2	⁴³ Ca (L4), ⁴⁸ Ca (H4), 8.338 s
Hop 3	⁴¹ K (L3), ⁴² Ca (L1), ⁴³ Ca (axial), ⁴⁴ Ca (H1), 16.776 s
Idle time	4 s after each hop
Duty cycles	90
Blocks	9 (1 block= 10 duty cycles)
Gain Calibration	Every 3 samples
Focus	Warm-up and every 4 blocks
Peak center	Warm-up and every 4 blocks
Baseline	Before each block, defocused, 30 cycles at 1.05 s/cycle
Amplifier rotation	After each block
Total analysis time/ sample	2.5 h

Table 3

Standard Name	⁴⁸ Ca- ⁴³ Ca double-spike							⁴³ Ca- ⁴² Ca double-spike			
	$\delta^{44/40}\text{Ca}_{\text{sw}}$	$2\sigma_{\text{SD}}$	$2\sigma_{\text{SEM}}$	$\delta^{44/42}\text{Ca}_{\text{sw}}$	$2\sigma_{\text{SD}}$	$2\sigma_{\text{SEM}}$	n	$\delta^{44/40}\text{Ca}_{\text{sw}}$	$2\sigma_{\text{SD}}$	$2\sigma_{\text{SEM}}$	n
OSIL Atlantic SW	0.000	0.224	0.047	0.000	0.135	0.028	23	-	-	-	-
915b	-1.144	0.234	0.054	-0.585	0.136	0.031	19	-	-	-	-
915a	-1.839	0.264	0.054	-0.930	0.153	0.033	22	-	-	-	-
CaF ₂	-1.398	0.201	0.061	-0.666	0.091	0.028	11	-	-	-	-
BHVO-1	-1.061	0.203	0.083	-0.494	0.115	0.047	6	-1.085	0.033	0.013	6

Table 4

Double-spike methods		
^{48}Ca - ^{43}Ca	$^{43}\text{Ca}_{dsp} = 0.40$ and $p_{dsp} = 0.40$ (This Study)	
^{48}Ca - ^{42}Ca	$^{42}\text{Ca}_{dsp} = 0.3964$ and $p_{dsp} = 0.1424$ (Rudge et al., 2009)	
^{46}Ca - ^{43}Ca	$^{43}\text{Ca}_{dsp} = 0.4898$ and $p_{dsp} = 0.2239$ (Rudge et al., 2009)	
^{43}Ca - ^{42}Ca	$^{42}\text{Ca}_{dsp} = 0.50$ and $p_{dsp} = 0.25$ (Lehn et al., 2013)	
Scenario #1	Pool 1	Pool 2
Starting Parameters		
Fractionation begins	$x=1$	$x=11$
β_0	0.235	0
$\Delta\beta$	0.0025	0.0025
Percentage Contribution to ion beam (%)		
Duty Cycle		
1 to 10	100	0
11 to 76	=Previous value - 1.5%	=100 - Pool 1
77 to 90	0	100
Scenario #2	Pool 1	Pool 2
Starting Parameters		
Fractionation begins	$x=1$	$x=1$
β_0	0.300	0
$\Delta\beta$	0.0025	0.0025
Percentage Contribution to ion beam (%)		
Duty Cycle		
1	100	0
2 to 90	=Previous value - 0.6%	=100 - Pool 1
Control	Pool 1	Pool 2
Starting Parameters		
Fractionation begins	$x=1$	n/a
β_0	0.500	n/a
$\Delta\beta$	0.0025	n/a
Percentage Contribution to ion beam (%)		
Duty Cycle		
1 to 90	100	0

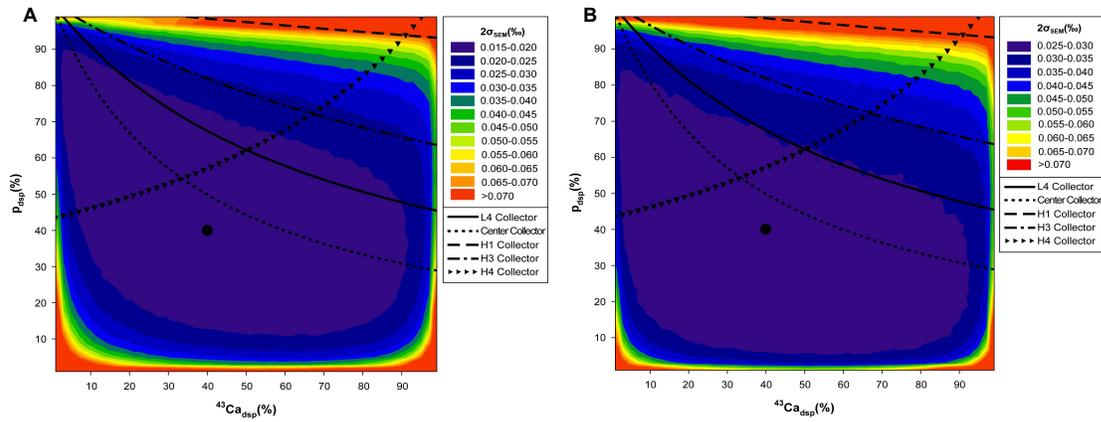


Figure 1

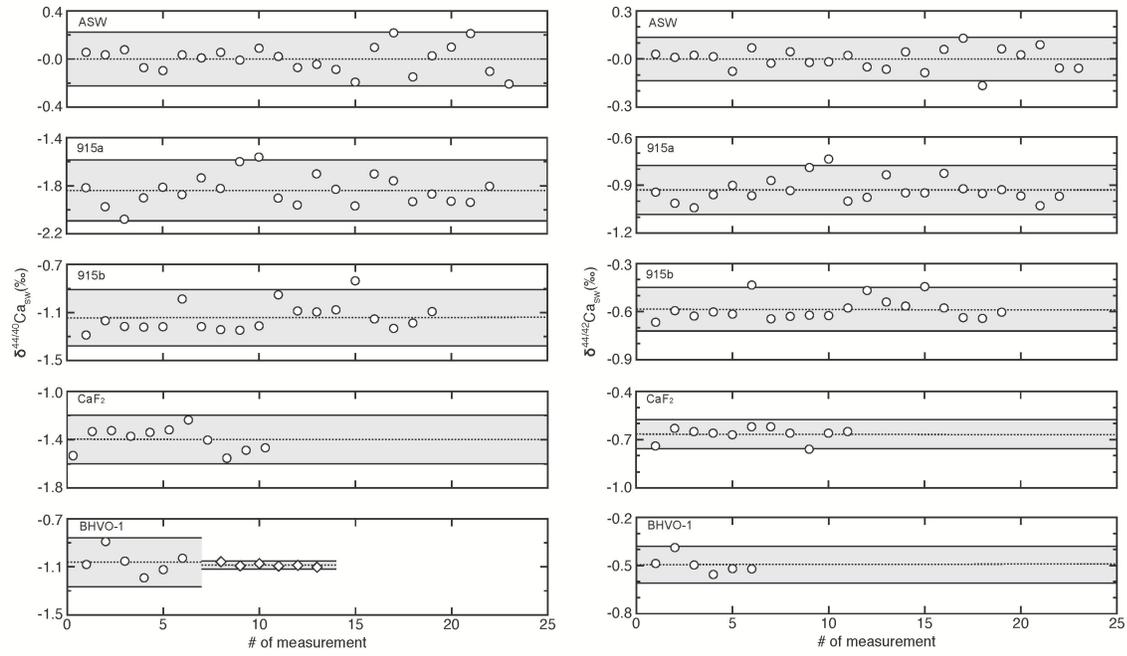


Figure 2

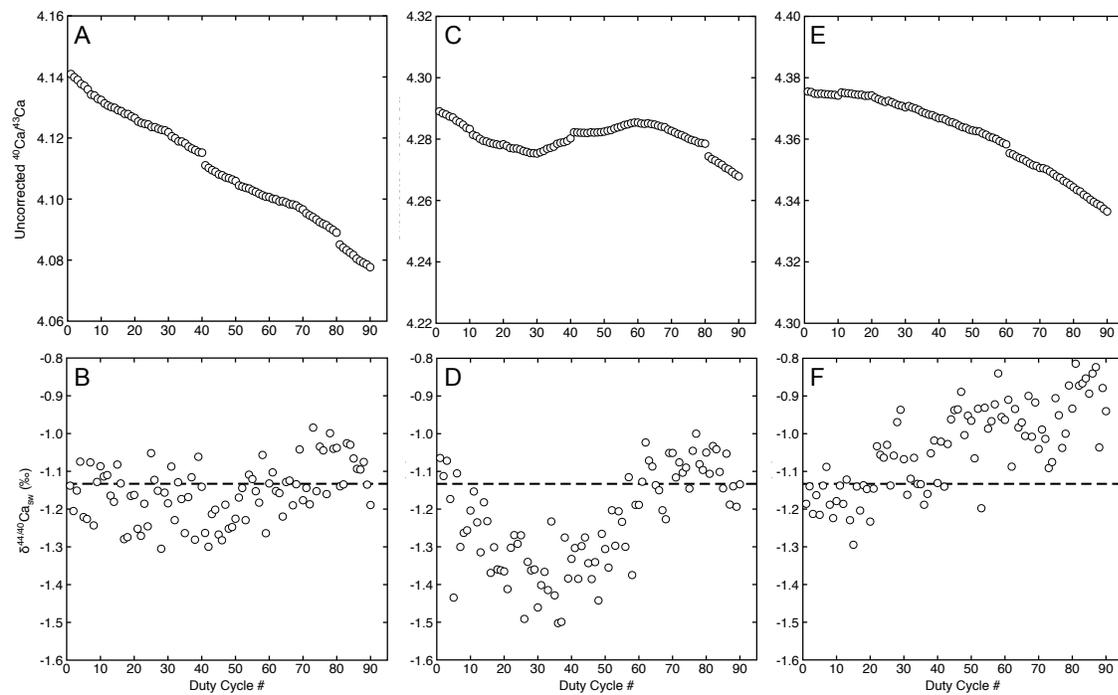


Figure 3

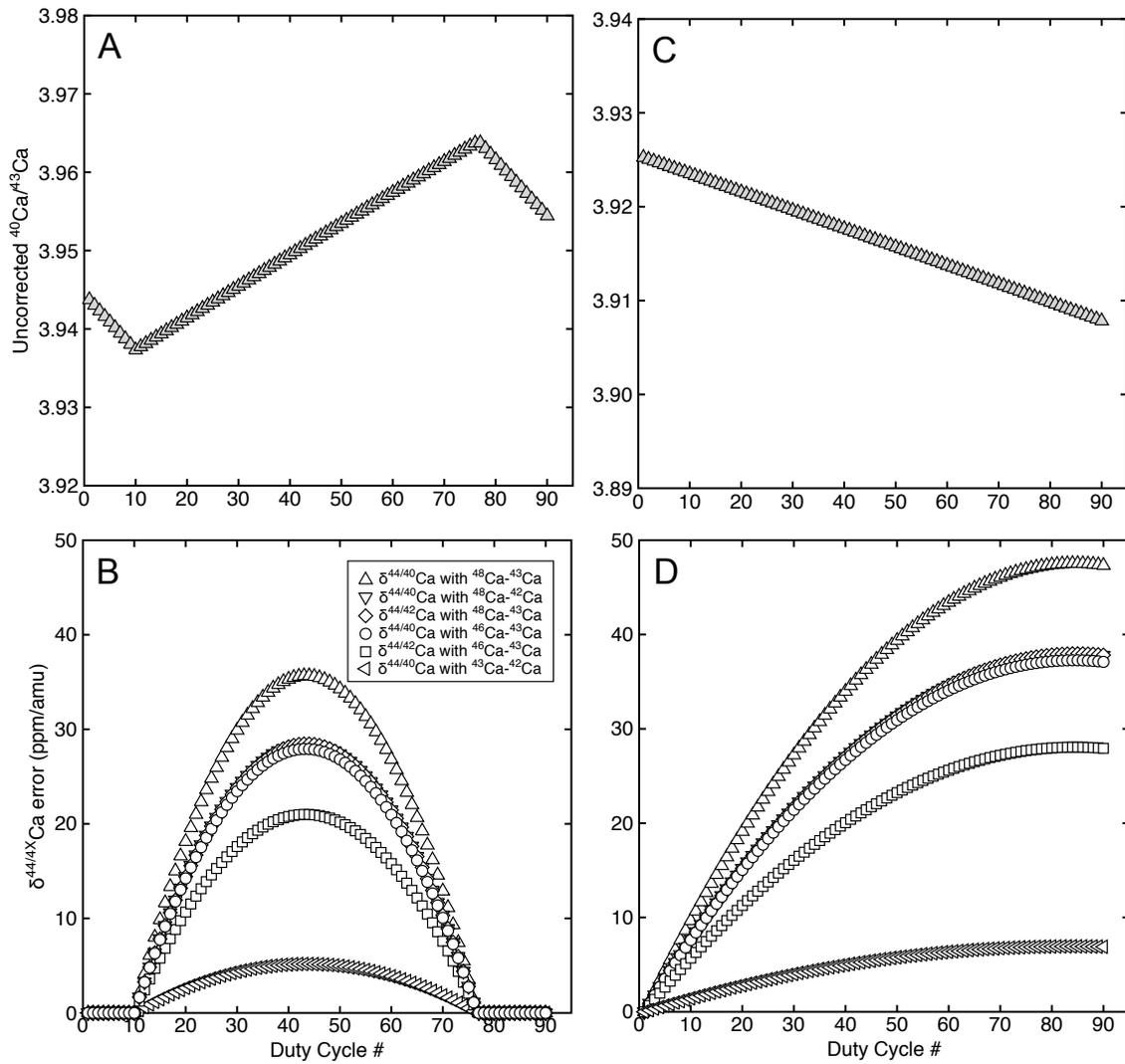


Figure 4

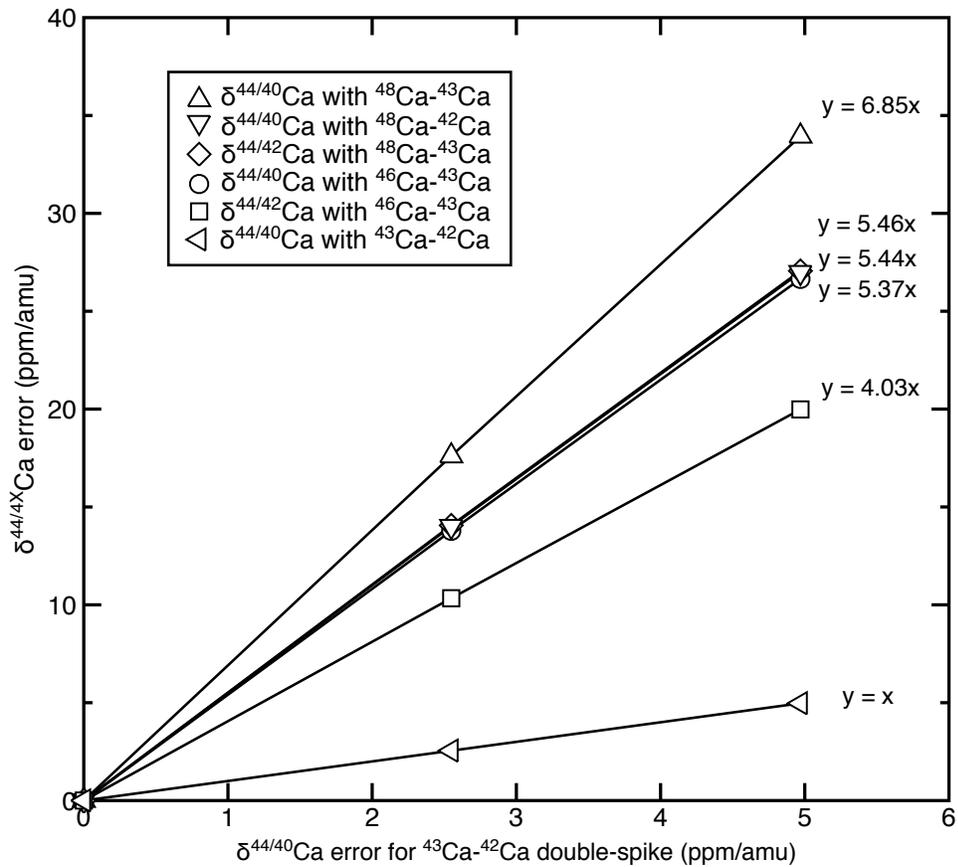


Figure 5

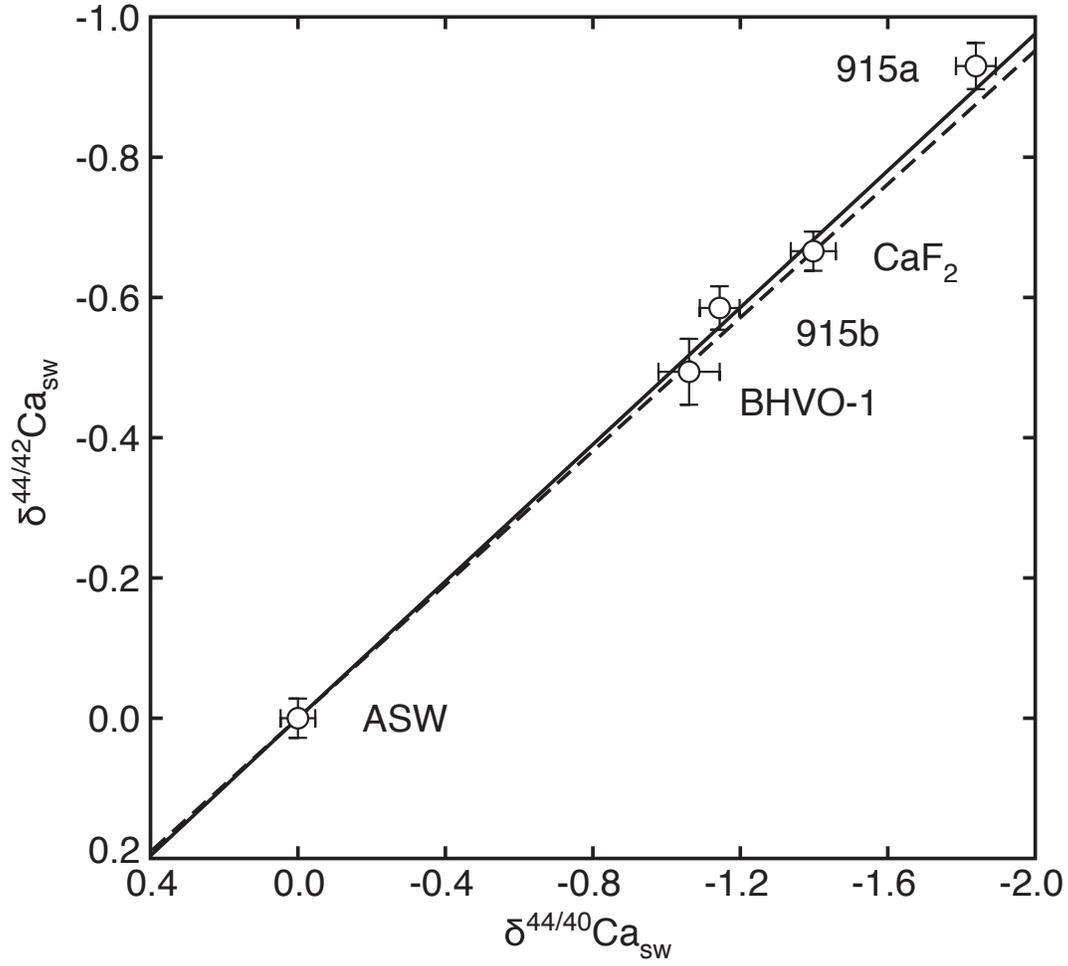


Figure 6