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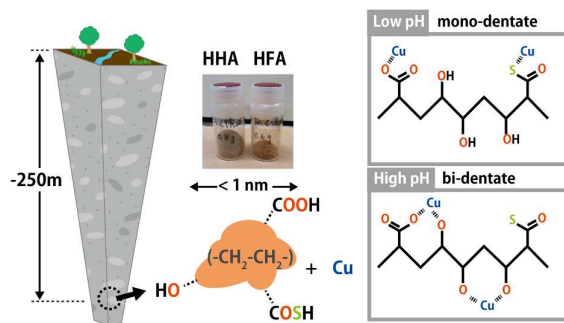
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Table of contents entry



We have revealed distinctive physicochemical and ion-binding properties of humic and fulvic acid from sedimentary deep groundwater.

1 Physicochemical and ion-binding properties of
2 highly aliphatic humic substances extracted from
3 deep sedimentary groundwater

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3 22 Humic substances (HSs) are ubiquitous in various aquatic systems and play important roles in
4
5 23 many geochemical processes. There are increasing evidences on the presence of HSs in deep
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8 24 groundwater; nevertheless, their ion binding properties are largely unknown. In this study we
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10 25 investigated the physicochemical and ion-binding properties of humic and fulvic acids
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12 26 extracted from deep sedimentary groundwater. The binding isotherms of proton (H^+) and
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14 27 copper (Cu^{2+}) were measured by potentiometry and fitted to the NICA-Donnan model, and
15
16 28 the obtained parameters were compared with the generic parameters of the model, which are
17
18 29 the average parameters for HSs from surface environments. The deep groundwater HSs were
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20 30 different from surface HSs, having high aliphaticities, high sulfur contents, and small
21
22 31 molecular sizes. The amounts of their acidic functional groups were comparable to or slightly
23
24 32 larger than those of surface HSs; however, the magnitude of Cu^{2+} binding to the deep
25
26 33 groundwater HSs was smaller. The NICA-Donnan model attributed this to the binding of
27
28 34 Cu^{2+} to chemically homogeneous low affinity sites, which presumably consist of carboxylic
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30 35 groups, via mono-dentate coordination at relatively low pH. The binding mode tended to shift
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32 36 to multi-dentate coordination with carboxylic groups and more heterogeneous
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34 37 alcoholic/phenolic groups at higher pH. X-ray absorption spectroscopy also revealed that
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36 38 Cu^{2+} binds to O/N containing functional groups and to lesser extent S containing functional
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38 39 groups as its divalent form. This study shows the particularity of the deep groundwater HSs
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40 40 in terms of their physicochemical and ion-binding properties, compared with surface HSs.
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42 41

42 **Environmental impact**

43 For future use of deep underground space it is necessary to monitor and protect the quality of
44 deep groundwater. Development of mechanistic models that can describe reactions of
45 pollutants with components in groundwater is mandatory as is the case for surface water
46 systems. Humic substances (HSs) play important roles in the speciation of metal ions;

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4 47 nevertheless, detailed ion binding to deep groundwater HSs is largely unknown. This study
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6 48 reveals the particularity of the physicochemical and ion-binding properties of the HSs
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8 49 extracted from sedimentary groundwater by comparing them to those of surface HSs.
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13 51 **1. Introduction**

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15 52 Humic substance (HS) is a class of natural organic matters, resulting from degradation and
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17 53 condensation of animal, plant and microbial remains, and ubiquitous in various
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20 54 environments: surface and ground water, ocean, soil, and atmosphere.^{1,2} HS is not a
21
22 55 molecular entity with a distinct structure, but should be considered as a group of organic
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24
25 56 molecules with certain physicochemical properties in common.³ Based on the solubility to
26
27 57 water at different pH, they are operationally divided to humic acid (HA), which is soluble at
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29 58 pH > 2, fulvic acid (FA), which is soluble at both acidic and alkaline pH, and insoluble
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31
32 59 humin. HSs play important roles in various environmentally-relevant processes; they
33
34 60 determine structures of micro aggregates in soils,³ stabilize metastable minerals,⁴ catalyze
35
36 61 redox reactions,⁵ and capture inorganic and organic contaminants.⁶⁻⁸ Proton and metal ions
37
38
39 62 readily bind to the functional groups of HSs,^{8,9} mostly carboxylic and phenolic groups and
40
41 63 less significantly amine and sulfur-containing groups, and alter their reactivity,
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43
44 64 bioavailability, and mobility.^{7,10}

45
46 65 Ion binding to HSs has been an active topic of research over decades.¹¹⁻¹³ The
47
48 66 particularity of HSs as ligands lies in their chemical heterogeneity and polyelectrolyte
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50
51 67 nature.^{9,13} The former is manifested in the distribution of the affinity constant of a HS for a
52
53 68 given metal ion due to diversity of the environments surrounding its functional groups. The
54
55 69 latter originates from negative charges locating on its carbon backbone, which creates
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57
58 70 negative electrostatic potential that attracts cations and excludes anions.^{14,15} Recent
59
60 71 mechanistic models for ion binding to HSs such as the NICA-Donnan model^{9,11} and the

1
2
3 72 Model VI and its successor^{12, 16} take these aspects into account and can successfully describe
4
5 73 the binding of various cations over a wide range of conditions. These models have been well
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7
8 74 tested for HSs extracted from different surface environments.^{12, 17, 18} There are certain
9
10 75 similarities in the obtained model parameters, once HAs and FAs are separately discussed.¹⁷
11
12 76 Thus, so-called “generic” parameters have been proposed for these two groups of HSs^{12, 17, 18}
13
14
15 77 and widely used to estimate the level of metal binding to HSs, for which specific model
16
17 78 parameters are unavailable.^{19, 20}

19
20 79 There are increasing evidences showing the presence of HS in deep underground
21
22 80 environments either as dissolved forms in pore water or bound to host rocks.²¹⁻²⁹ Deep
23
24 81 underground environments are rather different from surface aquatic systems, as manifested
25
26 82 by slow groundwater flow, which leads to prolonged interaction between rocks and
27
28 83 dissolved/suspended components, low oxygen concentration, and no direct energy input from
29
30 84 the sun. It is likely that HSs in deep undergrounds are different from their counterparts on
31
32 85 surface environments with respect to their structures and ion binding properties.^{21, 23}
33
34 86 Underground HSs may originate from surface waters transported by downward recharge,
35
36 87 dissolution of sedimentary organic materials, or in-situ production from remains of
37
38 88 microorganisms or algae in connate water, and have experienced long-term diagenesis. Ratios
39
40 89 of dissolved and bound HSs in deep underground environments are different from site to site,
41
42 90 and they may have rather different properties even in a given geological setting.³⁰

43
44 91 Uniqueness of deep underground HSs has been pointed out by several researchers.^{21, 23-26,}
45
46 92 ^{29, 31, 32} Schäfer et al.²³ studied possible sources of FAs in the Gorleben aquifer, based on
47
48 93 isotopic data and C and S X-ray absorption near-edge spectroscopy (XANES). They reported
49
50 94 that FAs derived from the deep brine groundwater at around -216 m below ground level (bgl)
51
52 95 had similar carbon backbone structures to those of FA in the corresponding shallow recharge
53
54 96 groundwater and that HAs and FAs originated from lignite in Miocene sediments were highly

1
2
3 97 aromatic. Alberts et al.³² reported that the properties of HA and FA extracted from
4
5 98 groundwater at -30 to -70 m bgl were different from those of their counterparts from surface
6
7
8 99 water, but copper binding to them was similar. Courdouan et al.³⁰ studied the binding of
9
10 100 trivalent metal ions to extracts of organic matters from pore water and rocks in the Opalinus
11
12 101 clay (OPA) and the Callovo-Oxfordian formations and found stronger binding of curium to
13
14 102 pore-water organic matters from OPA. Although some properties of deep underground HSs
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16
17 103 such as elemental composition, ¹³C NMR carbon distribution, optical properties, and
18
19 104 molecular-size distribution have been reported,^{21,24-26} their ion binding properties over a wide
20
21 105 range of conditions remain largely unknown.³¹⁻³³ This is particularly the case for deep
22
23 106 groundwater HSs, as large-scale extraction of HSs from deep groundwater is limited.

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27 107 Deep groundwater HSs can be extracted from pumped groundwater, using boreholes
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29 108 from surface. Some countries are operating underground research laboratories (URLs) for
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31 109 feasibility tests of geological disposal of nuclear wastes, where large amounts of groundwater
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33 110 samples are available with less contamination or alternation.³⁴⁻³⁷ This makes URLs
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35 111 appropriate places for extraction of deep groundwater HSs.

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39 112 Considering the future use of great-depth underground space by mankind such as
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41 113 geological disposal of nuclear wastes and potential deterioration of groundwater quality, ion
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43 114 binding properties of deep groundwater HSs are to be studied and the applicability of the
44
45 115 aforementioned mechanistic models is to be tested, as is the case for surface HSs. Thus, the
46
47 116 objective of this study is to reveal the physicochemical and ion-binding properties of HA and
48
49 117 FA isolated from sedimentary groundwater at the Horonobe URL of the Japan Atomic
50
51 118 Energy Agency (JAEA).³⁵ The physicochemical properties of the Horonobe HSs, which are
52
53 119 denoted as HHSs hereafter, were compared with those of surface HAs and FAs to discuss
54
55 120 their structural differences. Binding isotherms of proton (H⁺) and copper (Cu²⁺) was measured
56
57 121 over a wide range of conditions by potentiometric titration and fitted to the NICA-Donnan

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3 122 model.^{9,11} The results were compared to the model calculations with the generic parameters
4
5 123 proposed by Milne et al.¹⁸, which capture average trends of ion binding to HSs from surface
6
7 124 environments. Oxidation state and local coordination environment of Cu²⁺ bound to HA
8
9 125 fraction of the HHS were also assessed by X-ray absorption spectroscopy (XAS). Copper was
10
11 126 chosen as a representative divalent metal ion in this study to examine general metal binding
12
13 127 properties of the HHSs, as it can be easily quantified by an ion selective electrode (ISE) and
14
15 128 its binding to surface HSs have been well studied^{18,38,39}. Copper is also an essential trace
16
17 129 element for organisms at low concentration and becomes toxic at elevated concentration.^{40,41}
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19 130 It could be introduced to deep groundwater systems by exploitation of underground space, as
20
21 131 it is an important constituent of various materials used in modern industries. Thus, studies on
22
23 132 the speciation Cu²⁺ in the presence of groundwater HSs are relevant for its fate in deep
24
25 133 groundwater environments. Although the HA and FA from the single groundwater source are
26
27 134 examined in this study, the outcomes can be applied or be a good starting point to estimate
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29 135 the degree of metal binding to HSs in sedimentary groundwater similar to this study.
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39 137 **2. Materials and methods**

40 138 *2.1. Materials*

41 139 For the entire experiments, milli-Q grade pure water and analytical-grade chemicals
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43 140 purchased from the Wako Pure Chemical Industries were used, unless otherwise noted.
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48 141 HA and FA were extracted from groundwater collected at the -250-m gallery of the
49
50 142 Horonobe URL located in the northern part of Hokkaido Prefecture, Japan. The geology and
51
52 143 geochemistry of the site are described elsewhere.^{35,42} The -250-m gallery is located at the
53
54 144 boundary of the Pleiocene Koetoi and the Miocene Wakkanai formations, which are
55
56 145 composed of diatomaceous and siliceous mudstones, respectively. Groundwater at the
57
58 146 sampling location is a weakly alkaline Na⁺/HCO₃⁻ type with relatively high total organic
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3 147 carbon (TOC) and Cl^- levels. Groundwater after filtration and acidification was passed
4
5 148 through a column packed with DAX-8 resins (Supelite DAX-8, Sigma-Aldrich). Separation
6
7 149 and extraction of HA and FA fractions from the loaded resins and subsequent purification
8
9 150 was performed in a laboratory on surface, according to the protocol recommended by the
10
11 151 International Humic Substances Society (IHSS).⁴³ In total 6.6 g of HA and 3.5 g of FA were
12
13 152 obtained by treating approximately 6,000 L of the groundwater. These values correspond to
14
15 153 approximately 8.5 and 4.5% of the TOC in the groundwater, respectively.
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22 155 2.2. Characterization of HSs

23
24 156 Elemental compositions and ^{13}C NMR carbon distributions of the HHSs were evaluated in the
25
26 157 same manner in the previous report.³³ Carbon, H and N contents were determined using an
27
28 158 elemental analyzer (Yanagimoto, MT-6), and that of S was analyzed by ion chromatography
29
30 159 after conversion to SO_4^{2-} . Ash contents were also determined by combustion at 550 °C. For
31
32 160 ^{13}C NMR a 50-mg HS sample was dissolved in a mixture of 0.02 mL of 10 M NaOD (99.9%
33
34 161 deuteration, Sigma-Aldrich) and 0.4 mL of D_2O (99.9% deuteration, Sigma-Aldrich) solution.
35
36 162 As a reference material for the chemical shifts, a 0.02-mL of 1.0% solution of sodium 3-
37
38 163 trimethylsilylpropionate-2, 2, 3, 3, D_4 (TMSP; 98% deuteration, French Atomic Energy
39
40 164 Commission, CEA) was added into the mixture. The test solution was then passed through a
41
42 165 glass-fiber filter with a pore size of 0.7 μm , and placed in a 5-mm diameter spin tube. ^{13}C
43
44 166 NMR spectra were recorded by a Bruker AVANCE 500 spectrometer operating at 125.77
45
46 167 MHz. The inverse gated decoupling technique was applied for the measurement with the
47
48 168 pulse width of 45° and the acquisition time of 0.839 s. A total repetition time of 2.5 s was
49
50 169 applied to permit relaxation of all the spins, and 4000 - 20000 scans were accumulated.
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52 170 Chemical shift assignments were made using data reported by Wilson⁴⁴ and Fujitake and
53
54 171 Kawahigashi⁴⁵.

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3 172 The UV/Vis absorption spectra of the HHS as well as those of the standard or reference
4
5 173 HSs from the international humic substances society (IHSS) and the Japanese humic
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8 174 substance society (JHSS) and purified Aldrich HA (PAHA)⁴⁶ were measured in this study by
9
10 175 a UV/Vis spectrometer (UV-3100, Shimadzu). The samples were prepared at 50 mg/L HS
11
12 176 solutions in 0.01 M NaHCO₃ buffer.⁴⁷ The size distributions of HHA and HFA were
13
14
15 177 determined by flow-field flow fractionation (FI-FFF) with 1 kDa polyethersulfone membrane
16
17 178 (AF2000, Postnova), according to Lukman et al.⁴⁸ The electron accepting capacities (EAC)
18
19 179 of the HHS and the standard HSs from the IHSS and JHSS were determined by the mediator
20
21 180 electrochemical reduction (MER) in a similar way to Aeschbacher et al.⁴⁹, using diquat
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23 181 dibromide monohydrate (99.5%, Supelco) as a mediator. The details of the MER
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25 182 measurement are given in the SI.
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32 184 2.3. Potentiometric titration

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34 185 Potentiometric titration of HHA and HFA were performed, using the Wallingford titration
35
36 186 system.⁵⁰ HHS solutions were prepared by dissolving the freeze-dried samples in alkaline
37
38 187 solutions around pH 9 and kept stirred overnight.⁵¹ The pH and Cu²⁺ activities were measured
39
40 188 by a glass electrode (Metrohm, 6.0150.100) and a Cu ISE (Metrohm, 6.0502.140), combined
41
42 189 with a 3 M KCl Ag/AgCl reference electrode (Metrohm, 6.0733.100) in an electrolyte bridge
43
44 190 (0.1 M NaClO₄, Schott B511). All titrations were performed in a thermostated vessel under
45
46 191 slight over-pressure of moisturized Ar and continuous stirring. The glass electrode was
47
48 192 calibrated by titrating HClO₄ solutions with 0.1 M NaOH. The Cu ISE was calibrated by
49
50 193 titrating 0.36 mM Cu(ClO₄)₂ solution with 0.02 M ethylenediamine solution (Aldrich). The
51
52 194 electrode calibrations were performed at the same salt levels as those in the subsequent
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54 195 sample titrations.
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3 196 Proton binding isotherms of the HHSs were obtained by acid-base titration, as described
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5
6 197 elsewhere.⁴⁶ Thirty milliliter of a 1 g/L HHA or HFA solution was first titrated to pH 4 and
7
8 198 stirred for 1 hour; then three-sets of forward and backward titrations were performed. The salt
9
10 199 concentration of the solution was increased by adding a 1 M NaClO₄ solution (Merck)
11
12 200 between the different sets of the titration. At every point of the titration the reading of the
13
14 201 glass electrode was recorded when the drift became less than 0.1 mV/min or after 30 min.
15
16 202 The relative positions of the charge ($-q$)-pH curves of the HHSs at the different salt levels
17
18 203 were determined from the amounts of base and acid titrants necessary to back-titrate H⁺
19
20 204 released in pH-stat salt titration. The absolute position of the curves was then determined by
21
22 205 optimizing the initial negative charge, q_0 , at the beginning of the titration in the NICA-
23
24 206 Donnan fitting¹¹, as described in the supporting information (SI). The results of the forward
25
26 207 base titration were used in the subsequent fitting, as the hysteresis between the forward and
27
28 208 backward titrations at a given salt level was small. The uncertainty in the determination of the
29
30 209 HHS charge, q , was estimated to be less than 0.1 meq/g, using a typical standard deviation of
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32 210 the glass electrode calibration (0.05 pH unit).
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39 211 Copper binding isotherms to the HHSs were measured by pH-stat titration at three pH
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41 212 levels (pH 4, 6, and 8) and 0.1 M NaClO₄.³⁸ At pH 4 the additional titration at 0.01 M NaClO₄
42
43 213 was performed. The Cu²⁺ titration to HFA at pH 8 failed most likely because of poor pH
44
45 214 buffering (see the discussion below). Thirty milliliter of a 1 g/L HHA or HFA solution was
46
47 215 first titrated to pH 4 and stirred for 1 hour, and then to desired pH and equilibrated within 0.2
48
49 216 mV (0.003 pH unit) for 30 min. After equilibration a 0.1 M Cu(ClO₄)₂ solution or 10⁻³ M
50
51 217 Cu(ClO₄)₂ solution in 0.1 M NaClO₄ was added. The pH of the sample solution was back-
52
53 218 titrated to the original value and kept within 0.2 mV for 20 min by addition of the acid and
54
55 219 base titrants. Readings of the electrodes were recorded after their drifts became less than 0.1
56
57 220 mV/min or after 20 min. At each titration point the solution was checked for the formation of
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3 221 $\text{Cu}(\text{OH})_2(\text{s})$ ($\log K_{\text{sp}} = -19.32$ ⁵²) and the Cu^{2+} binding amount ($[\text{Cu}^{2+}]_{\text{bound}}$) was calculated by
4
5 222 subtracting the sum of the concentrations of free Cu^{2+} and its hydrolysis products from the
6
7
8 223 total concentration, using the hydrolysis constants of Cu^{2+} .⁵² The magnitude of the uncertainty
9
10 224 in $\log[\text{Cu}^{2+}]_{\text{bound}}$ was estimated to be no more than 0.2, using a typical error of the Cu ISE
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12
13 225 calibration (0.06 as the logarithm of Cu^{2+} activity, $\log a_{\text{Cu}}$).

14
15 226 The obtained H^+ and Cu^{2+} binding isotherms to HHA and HFA were fit to the NICA-
16
17 227 Donnan model,^{9,11} using an in-house MATLAB® program. The details of the model as well
18
19 228 as the fitting procedure are given in the SI. First, the maximum density of H^+ binding sites,
20
21 229 $Q_{\text{max},j,\text{H}}$ of the site j ($j = 1$ and 2 for the low-affinity and high-affinity sites, respectively), the
22
23 230 median values of the affinity constants of the site j for H^+ , $\tilde{K}_{j,\text{H}}$, the apparent heterogeneity
24
25 231 parameters of the site j , m_j , the Donnan parameter, b , and q_0 were optimized, using the
26
27 232 charge/pH curves. Then, the median values of the affinity constants of the site j for Cu^{2+} ,
28
29 233 $\tilde{K}_{j,\text{Cu}}$, the ion-specific non-ideality parameters of the site j for H^+ and Cu^{2+} , $n_{j,\text{H}}$, $n_{j,\text{Cu}}$, and
30
31 234 the heterogeneity parameters of the site j , p_j , which correspond to the reciprocal of the width
32
33 235 of the affinity distribution, were optimized by fitting to the Cu^{2+} binding isotherms, while $n_{j,\text{H}}$
34
35 236 $\times p_j$ was kept equal to m_j .¹¹ The lower and upper boundaries were set to 0 and 1 for the m_j , n_j ,
36
37 237 i and p_j parameters.

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239 2.4. XAS analysis

240 Copper K-edge XANES and extended X-ray absorption fine structure (EXAFS) analyses of
241 Cu^{2+} reacted with HHA and PAHA were carried out at the BL-27B in the Photon Factory,
242 KEK. The HA samples were dissolved in alkaline solutions at 4.0 g/L and stirred overnight
243 under Ar atmosphere. After adjusting pH to 4 or 7 with 0.1 or 0.01 M HCl and NaOH, a 10
244 mM CuCl_2 solution was added to achieve the Cu^{2+} loading of 80 mmol Cu/Kg HA, and pH
245 was re-adjusted to the original values. These values correspond to the Cu^{2+} binding amounts

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4 246 of 13.3 ($\log [\text{Cu}^{2+}]_{\text{bound}} = -1.88$) and 310 ($\log [\text{Cu}^{2+}]_{\text{bound}} = -0.51$) mmol/Kg, respectively,
5
6 247 according to the NICA-Donnan model calculation with the optimized parameters for HHA.
7
8 248 After equilibration for two days, the samples were freeze-dried, mixed with boron nitride
9
10 249 (BN), and pressed into pellets, which were covered by a Kapton® tape. Reference solid
11
12 250 compounds ($\text{Cu}^{\text{II}}\text{O}$, $\text{Cu}^{\text{II}}\text{Cl}_2$, $\text{Cu}^{\text{I}}\text{Cl}$, and $\text{Cu}^{\text{I}}\text{SCN}$) were dispersed in BN and pressed into
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14
15 251 pellets. In addition a 0.02 M Cu^{2+} solution with 0.04 M L(+)-tartrate at pH 7 were also
16
17 252 measured as reference.

18
19
20 253 Copper K-edge X-ray absorption spectra (XAS) were measured in fluorescence mode at
21
22 254 148 K using a liquid N_2 cryostat equipped with Kapton® windows (CoolSpek, UNISOKU)
23
24 255 for the HA samples and in transmission mode at room temperature for the reference materials.
25
26 256 A Si(111) double crystal monochromator was detuned by about 50% to reject higher
27
28 257 harmonic intensity. Reduction and theoretical fitting of the obtained XAS data was performed
29
30 258 by the Athena and Artemis software packages⁵³ and FEFF 6.⁵⁴ The details of the data
31
32 259 reduction and fitting are given in the SI.

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38 39 261 **3. Results and discussion**

40 41 262 *3.1. Physicochemical properties of the HHSs*

42
43 263 The elemental compositions and ^{13}C NMR distributions of HHA and HFA are given in Table
44
45 264 1 and the more detailed physicochemical properties of HHA and HFA are summarized in
46
47 265 Table S1 in the SI together with those of the IHSS and JHSS HSs and PAHA. These standard
48
49 266 or reference HSs are isolated from various surface environments, ranging from soils (EHA
50
51 267 from the Elliot soil, IHA and IFA from the Inogashira volcanic ash soil, DHA and DFA from
52
53 268 the Dando forest soil), peats (PHA from the Pahokee peat), oxidation products of lignite
54
55 269 (Leonardite HA, LHA), rivers (SRHA and SRFA from the Suwannee River), and lakes (NHA
56
57 270 and NFA from the Nordic lake and BFA from the Biwako lake). The elemental compositions

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3 271 of the HHSs are characterized by their low oxygen and ash contents and high hydrogen and
4
5 272 sulfur contents, compared to the HAs and FAs from surface environments. The relatively
6
7
8 273 small O/C and large H/C ratios of the HHSs can be manifested in the van Krevelen plot (Fig.
9
10 274 1).⁵⁵ In Fig. S1 of the SI a similar plot with broader data in literatures is also shown. The
11
12 275 locations of the standard and commercial surface HAs and FAs depend on their types and
13
14 276 origins (Fig. 1). The former is characterized by relatively high O/C ratios; the latter exhibits a
15
16 277 wide range of H/C ratios with constant and relatively small O/C ratios. The aquatic HA and
17
18 278 FA are close to each other except for BFA. Groundwater HSs tend to have small O/C ratios,
19
20 279 as is shown in Fig. S1. Some of groundwater HA as well as marine HS show relatively large
21
22 280 H/C ratios. Compared with HSs from various environments, HHA and HFAs are different in
23
24 281 terms of the O/C and H/C ratios as is shown in Figs. 1. The large H/C values of the HHSs
25
26 282 indicate the abundance of saturated carbons, which is in line with their small aromaticities
27
28 283 estimated from ¹³C NMR C distribution (Table 1 and Table S1 in the SI). Highly aliphatic
29
30 284 nature of organic matters in deep underground environments has been pointed out by several
31
32 285 researchers.^{26,29,31} Pettersson et al.²⁶ reported an even higher H/C value (H/C = 1.7) for HA
33
34 286 extracted from the granitic groundwater at -280 m bgl. Claret et al.³¹ also reported the highly
35
36 287 aliphatic nature of FAs extracted from argillite of Meuse and Opalinus shale and discussed
37
38 288 their origin as oceanic sediments at high burial temperature.

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40 289 The densities of oxygen-containing carboxylic and phenolic functional groups of the
41
42 290 HHSs determined by conventional end-point acid-base titration in Table S1 are comparable
43
44 291 to those of the surface HSs. This means that the oxygen depletion indicated by the small O/C
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46 292 ratios of the HHSs occurs in functional groups other than carboxylic and phenolic-type
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48 293 groups, as described by Thurman.⁵⁶ The UV/Vis optical properties of the HHSs are
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50 294 characterized by relatively small $A_{250/210}$ and $A_{350/280}$ ratios (Table S1). This points to the
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3 295 presence of small conjugated systems in the HHSs⁴⁸ and again in accordance with their low
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5 296 aromaticities.

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8 297 The size distributions of the HHSs measured by FI-FFF are compared to those of the
9
10 298 IHSS and JHSS standard HSs and PAHA in Fig. 2. For the surface HSs the sizes of the HAs
11
12 299 are larger than those of the FAs. The size distributions of the HAs are largely overlapped,
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14 300 although their shapes are somewhat different from each other; EHA, IHA, and LHA possess
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16 301 multiple peaks. The size distributions of the FAs are all mono-modal, and the peak locations
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18 302 are different, depending on their sources. It seems that the JHSS FAs (BFA, DFA, and IFA)
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20 303 are somewhat smaller than the IHSS FAs (SRFA and NFA). The size distributions of HHA
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22 304 and HFA are mono-modal with the modal sizes of 0.6 and 0.3 nm, respectively, which are
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24 305 appreciably smaller than those of the surface HSs. Relatively small sizes of deep
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26 306 underground organic matters have reported in the literatures.^{57,58} Bouby et al.⁵⁸ reported that
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28 307 organic matters in the Gorleben groundwater had a modal size of 1 nm. Saito et al.⁵⁷
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30 308 compared the size distribution of organic matters in granitic and sedimentary groundwater by
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32 309 FI-FFF. The sedimentary groundwater was taken from a borehole at the same depth in the
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34 310 Horonobe URL as in this study and exhibited a mono-modal size distribution with a peak
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36 311 around 2 nm. This may indicate that HA and FA fractions may account for only a part of
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38 312 dissolved organic matters in this groundwater.

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41 313 The EAC of a HS corresponds to the number of electrons transferred to the HS from the
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43 314 mediator, normalized by the mass of the HS, and represent its redox capacity. For the HHSs
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45 315 EAC values are relatively small, compared with those of the IHSS and JHSS HSs (Fig. S2 in
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47 316 the SI). As in Aeschbacher et al.⁴⁹, we found linear relationship between the EAC and the
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49 317 aromaticities of the HSs investigated (Fig. S2). This is because the concentration of quinone
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51 318 moieties, that are predominantly responsible for redox reactions in HSs, tends to be
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3 319 proportional to the amount of aromatic carbon.⁴² Thus, the HHSs with low aromaticities have
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5 320 small redox capacities, compared with the surface HSs.
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8 321 In summary the HHSs can be viewed as relatively small organic matters with abundant
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10 322 aliphatic carbons and sulfurs. The densities of acidic functional groups are comparable to
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12 323 those of the surface HSs. The differences between HHA and HFA are small. The cluster
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14 324 analysis (Fig. S3 in the SI) performed for the physicochemical properties compiled in Table
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16 325 S1 clearly indicates that they are different from the IHSS and JHSS HSs. BFA is an exception,
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18 326 being clustered into the same group as the HHSs. This may indicate the presence of similar
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20 327 formation processes among them.
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26 27 329 *3.2. H⁺ and Cu²⁺ binding isotherms to the HHSs*

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29 330 The charge/pH curves of HHA and HFA at the different salt levels are presented in Fig. 3.
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31 331 The negative charge ($-q$) of the HHSs increases with pH and salt concentration, as is usually
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33 332 seen for surface HSs.¹⁷ The maximum negative charge of HFA is larger than that of HHA,
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35 333 which predominantly arises from larger deprotonation at acidic pH ($\text{pH} < 6$). This further
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37 334 suggests that HFA possesses more acidic functional groups, mostly carboxylic groups, than
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39 335 HHA. In Fig. 3 the charge/pH curves of the HHSs are compared to those calculated by the
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41 336 NICA-Donnan model with the generic parameters derived for surface HSs (Table 2). At
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43 337 acidic pH the slopes of the curves are larger for the HHSs. At neutral and alkaline pH this
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45 338 trend is reversed, although the differences are small. The slope of a charge/pH curve of a HS
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47 339 reflects the width of the corresponding affinity distribution of its functional groups. A smaller
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49 340 slope means a wider distribution and larger chemical heterogeneity.⁹ Note that at $6 < \text{pH} < 9$
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51 341 the negative charges of the HHSs hardly change, indicating that the number of the acidic
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53 342 functional groups having corresponding $\text{p}K_a$ is small. Thus, these comparisons imply that the
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55 343 negative charges of the HHSs largely originate from H⁺ dissociation from chemically
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3 344 homogeneous low-affinity sites, which should mainly consist of carboxylic groups located on
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5 345 the aliphatic chains of the HHSs, considering their elemental compositions and ^{13}C -NMR
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8 346 carbon distributions (Table 1). The salt dependence of the charge/pH curves is also different
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10 347 between the HHSs and the model calculation. Relatively large salt effects are observed at
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12 348 neutral to alkaline pH for the HHSs, but at acidic pH in the model calculation with the
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14 349 generic parameters.

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17 350 The Cu^{2+} binding isotherms to HHA and HFA are presented in Fig. 4. The isotherms are
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19 351 similar between them. The binding amounts of Cu^{2+} to the HHSs increases with the Cu^{2+}
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21 352 concentration and pH, as is expected for metal binding to surface HSs.^{18,38} The observed pH
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23 353 dependency of Cu^{2+} binding is the result of diminished H^+ competition to the functional
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25 354 groups of the HHSs with an increase of pH. The Cu^{2+} binding amounts also slightly decrease
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27 355 with an increase of salt concentration due to screening of the negative electrostatic potential
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29 356 of the HHSs. In Figs. 4 the Cu^{2+} binding isotherms calculated by the NICA-Donnan model
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31 357 with the generic parameters given in Table 2 are presented for comparison. The binding
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33 358 amounts of Cu^{2+} to the HHSs are smaller than those of the model calculations regardless of
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35 359 pH and salt concentration. Weak binding of europium is also reported for HA and FA
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37 360 extracted from groundwater collected at - 495-550 m bgl through a surface borehole in the
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39 361 Horonobe URL.³³ At pH 4 the slopes of the isotherms to the HHSs are close to 1 in the log-
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41 362 log plot, which are larger than those of the model calculations with the generic parameters at
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43 363 the same pH. Interestingly, the differences in the slopes became smaller at higher pH, and at
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45 364 pH 8 for HHA it becomes similar to the slope of the calculated Cu^{2+} isotherms. The slope of a
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47 365 metal-binding isotherm of HS in log-log plot is determined by the combination of the
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49 366 chemical heterogeneity of sites and the ion-specific non-ideality such as stoichiometry of the
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51 367 binding.⁹ The slope close to 1 in the isotherms of the HHSs at pH 4 together with the
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53 368 relatively large slopes of their charge-pH curves at acidic pH (Fig. 3) indicate Cu^{2+} binding to
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3 369 relatively homogeneous sites via mono-dentate coordination. At higher pH it seems that the
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5 370 binding mode tends to shift to coordination with greater denticity such as bi-dentate
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8 371 coordination.

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13 373 *3.3. NICA-Donnan modeling*

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15 374 The results of the NICA-Donnan fitting to the H^+ and Cu^{2+} isotherms to HHA and HFA are
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17 375 presented in Figs. 3 and 4 and the optimized parameters are given in Table 2 together with
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19 376 the Milne's generic parameters.¹⁷ The 95% confidence intervals and the correlation matrices
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21 377 of the optimized parameters are given in Tables S2, S3 and S4 in the SI. Note that some
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23 378 parameters associated with the high-affinity sites, namely $Q_{\max,2,H}$, $\log \tilde{K}_{2,H}$, and $\log \tilde{K}_{2,Cu}$
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25 379 suffered from large errors. This is because that these parameters were not well fitted to the
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27 380 model due to the limited experimental conditions for H^+ and Cu^{2+} binding to the HHSs in
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29 381 alkaline pH ($pH \leq 10$ for the charge-pH curves and $pH \leq 8$ or 6 for the Cu^{2+} isotherms). In
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31 382 order to unequivocally determine these parameters, potentiometric titration in non-aqueous
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33 383 media would be necessary.

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35 384 For H^+ binding the model successfully reproduces the charge-pH curves especially at pH
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37 385 < 6 . At neutral to alkaline pH the model somewhat underestimated the magnitude of the salt
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39 386 effect. The electrostatic part of the NICA-Donnan model (eps (S4) and (S5) in the SI)
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41 387 assumes a relatively simple functional form for the so-called Donnan volume, which depends
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43 388 only on the salt concentration.¹⁴ Although the Donnan model is relatively simple with only
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45 389 one adjustable parameter and advantageous over other more sophisticated but complex
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47 390 models, its potential flaw for small FAs has been recognized.⁵⁹ Considering the sizes of HHA
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49 391 and HFA, which are smaller than the Debye length of the solutions (1 and 3 nm for 0.1 and
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51 392 0.01 M $NaClO_4$), electrostatic potential calculation by the rigid-sphere or ion-permeable
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53 393 sphere model would be more realistic.¹⁴ The discrepancy observed at $pH > 6$ may also

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3 394 indicate the presence of a pH-dependent conformational change in the HHSs. Such
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5 395 conformational changes are common for linear aliphatic polyelectrolytes such as
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8 396 polymethacrylic acid.⁶⁰ For HHA Cu²⁺ binding is well reproduced by the model over a wide
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10 397 range of conditions, using the single set of the parameters. For HFA the model overestimated
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12 398 the salt effect and failed describing the Cu²⁺ binding at 0.01 M NaClO₄ and pH 4. It is likely
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14 399 that the model mishandles the electrostatic potential of HFA.
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18 400 The optimized NICA-Donnan parameters for HHA and HFA are more or less similar to
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20 401 each other except for $Q_{\max 1, H}$, which is larger for HFA. The maximum densities of H⁺ binding
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22 402 sites are similar between the high ($j = 1$) and low ($j = 2$) affinity sites in HHA. For HFA the
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24 403 density of the latter group was smaller by 1.5 meq/g. The obtained parameters can be
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26 404 compared to those of surface HSs with various origins and the generic parameters in Table
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28 405 2.¹⁸ The values of $Q_{\max 1, H}$ of the HHSs are in the ranges reported for the surface HSs, while
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30 406 those of $Q_{\max 2, H}$ are larger.¹⁷ The median affinity constants of H⁺ of the HHS are larger than
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32 407 those of most of the surface HSs and the Milne's generic parameters. The heterogeneity
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34 408 parameter, p_j , and ion-specific non-ideality parameter, $n_{j, H}$ are also relatively large for the
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36 409 HHSs. This is especially the case for the low affinity carboxylic-type sites, reflecting the
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38 410 large slopes of their charge/pH curves (Fig. 3) at pH < 6. The Donnan parameters, b , are also
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40 411 larger for the HHSs. Considering the large aliphaticity of the HHSs as discussed in the
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42 412 physicochemical characterization, the $\log \tilde{K}_{2, H}$ values of HHA (10.62) and HFA (10.48),
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44 413 which are larger than the corresponding values of the generic HA (8.60) and FA (8.00), may
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46 414 indicate larger contribution of alcoholic hydroxyl groups to the sites of the HHSs than those
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48 415 of surface HSs, although the presence of phenolic groups with large pK_a can not be entirely
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50 416 neglected as the HHSs still contain a certain amount of aromatic carbons (Table 1). This can
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52 417 also explain weak H⁺ buffering by the HHSs at neutral to alkaline pH (Fig. 3).
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3 418 The NICA-Donnan parameters of Cu^{2+} binding to HHSs are rather different from those
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5 419 of the generic parameters derived by Milne et al. (Table 1).¹⁸ For the low-affinity sites
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8 420 $\log \tilde{K}_{1,\text{Cu}}$ is larger for HHA and smaller for HFA than the corresponding generic parameters;
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10 421 whereas $n_{1,\text{H}}$ for both HHA and HFA are 1 and larger than the corresponding generic
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12 422 parameters (0.56 for GHA and 0.53 for GFA). For the high-affinity sites $\log \tilde{K}_{2,\text{Cu}}$ of HHA
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14 423 and HFA are larger than those of the generic parameters, and $n_{2,\text{H}}$ are smaller. Thus, the Cu^{2+}
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16 424 binding to the HHSs are characterized by $n_{1,\text{Cu}} = 1$ for the low-affinity sites and large
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18 425 $\log \tilde{K}_{2,\text{Cu}}$ and small $n_{2,\text{Cu}}$ values for the high-affinity sites. A ratio of the parameter $n_{j,i}$ of a
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20 426 metal ion and proton is a good indicator of underlying denticity of complexation reaction
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22 427 ($n_{i,j}/n_{j,\text{H}}$ close to 1 for mono dentate binding and 0.5 for bi-dentate binding). Thus, the
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24 428 optimized NICA-Donnan parameters for the HHSs suggest the mono-dentate nature of Cu^{2+}
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26 429 binding to the chemically homogeneous low-affinity sites at acidic pH. Considering the
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28 430 comparable or larger densities of the sites, $Q_{\text{max}1,\text{H}}$, of the HHS to/than those of surface HSs,
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30 431 this is most likely caused by sparsely distributing carboxylic groups on aliphatic backbones
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32 432 of the HHSs, which are hard to form bi-dentate coordination with Cu^{2+} . This can also explain
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34 433 the weak Cu^{2+} binding to the HHSs (Fig. 4). The relatively small $n_{2,\text{Cu}}/n_{2,\text{H}}$ values for the
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36 434 HHSs together with the relatively small p_2 values indicate that the more heterogeneous
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38 435 phenolic/alcoholic-type groups involve in the binding of Cu^{2+} at neutral to alkaline pH via
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40 436 multi-dentate coordination. It is of interest to compare the NICA-Donnan parameters for the
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42 437 HHSs to those optimized for other groundwater HSs. Marang et al.⁶¹ reported the NICA-
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44 438 Donnan parameters for the binding of divalent and trivalent metal ions including Cu^{2+} to HA
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46 439 extracted from deep groundwater (-139 m bgl) of the Gorleben aquifer in Germany. The
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48 440 NICA-Donnan parameters of Cu^{2+} for this HA are more like those of GHA than those of
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50 441 HHA, suggesting diversity of ion binding properties of groundwater HSs. Further research is
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60 442 needed to relate it to origin and genesis of groundwater HSs.

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444 *3.4. Oxidation state and chemical environment of Cu²⁺ bound to HHA*

445 The oxidation state of Cu and its coordination environment bound to HHA was assessed by
446 XAS. The XANES spectra of Cu²⁺ bound to HHA and PAHA are compared to those of the
447 reference compounds in Fig. 5. The Cu^I reference materials exhibit characteristic pre-edge
448 features (8.985 keV for CuSCN and 8.988 keV for CuCl) originating from 1s → 4p
449 transitions.⁶² The Cu^{II} compounds, on the other hand, are characterized by intense white lines
450 around 8.997 to 9.000 keV due to 1s → continuum transitions.⁶² The XANES spectra of Cu²⁺
451 bound to HHA and PAHA at pH 4 and 7 resemble those of Cu^{II} reference materials,
452 especially Cu-tartrate. This suggests that reduction of Cu²⁺ bound to HHA and PAHA is
453 negligible, although the magnified plot of the edge regions reveals a slight increase at 8.985
454 keV and a decrease at 8.998 keV for HHA (Fig. S4 in SI). Fulda et al.⁶² reported reduction of
455 Cu²⁺ upon binding to a soil HA under anoxic condition by XAS. It has been also reported that
456 SRFA rapidly reduces Cu²⁺ even under oxic conditions.⁶³ Virtually no reduction of Cu²⁺ by
457 HHA may agree with its low EAC obtained by MER (Fig. S2).⁴⁹

458 The k^3 -weighted EXAFS spectra of Cu²⁺ bound to HHA and PAHA at pH 4 and 7 (Fig. 6
459 (a)) exhibit systematic differences between the two HAs at 5.7, 6.5, and 7.4 Å⁻¹. In Fig. S4 of
460 SI the same spectra are shown as an overlapped plot. The magnitudes of the corresponding
461 Fourier transforms (Fig. 6 (b) and Fig. S5 (b) in SI) are also different between them. For both
462 HAs the Fourier transformed magnitudes are dominated by the intense peaks around 1.5 Å,
463 which corresponds to the scattering from the nearest O (and, to lesser extent, N). For HHA
464 additional peaks are noticeable around 2 Å, although their magnitudes are small. These peaks
465 most likely originate from Cu-S path as suggested by others.^{49,64} This assignment agrees with
466 the high S content of HHA (Tables 1). Quantitative modeling of the first coordination sphere
467 of Cu²⁺ (Fig. 5 (b) and Table 3) shows that approximately four O exist at 1.93 Å for PAHA,

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3 468 which corresponds to the Jahn-Teller distorted coordination geometry around Cu^{2+} . For HHA
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5 469 at pH 4 the coordination number (CN) of O in the first shell is decreased and a small but non-
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8 470 negligible number of S (CN = 0.4 ± 0.2) is found at 2.35 Å. At pH 7 the presence of S in the
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10 471 second shell is inconclusive. Note that this does not necessarily mean bi-dentate coordination
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12 472 of Cu^{2+} with O/N and S in HHA. Considering the results of the NICA-Donnan fitting (Fig. 4
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14 473 and Table 2), it seems more likely that the two independent mono-dentate sites exist in HHA,
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16 474 as an EXAFS spectrum is one-dimensional representation of coexisting coordination
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18 475 environments of a target element. The S containing functional group could be the thioacetic
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20 476 group, R-COSH, which then should be a part of the low affinity sites in the NICA-Donnan
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22 477 modeling, as the $\text{p}K_a$ value of thioacetic acid (CH_3COSH , $\text{p}K_a = 3.33$) is close to that of acetic
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24 478 acid ($\text{p}K_a = 4.76$) and smaller than that of methanethiol (CH_3SH , $\text{p}K_a = 10.33$).⁶⁵ It is also
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26 479 noteworthy to mention that the contribution of S to Cu^{2+} binding to HHA might be
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28 480 underestimated, compared with that in the original groundwater, as the functional groups
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30 481 containing reduced S may be oxidized during the extraction and storage. For more detailed
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32 482 discussion S speciation in the HHSs should be determined by soft X-ray absorption
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34 483 spectroscopy.²³
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43 485 **4. Conclusion**

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45 486 The present study aims to reveal physicochemical and ion binding properties of the HHSs
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47 487 isolated from deep groundwater in a sedimentary rock formation. It is found that the HHSs
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49 488 are rather different from surface HSs, characterized by high aliphaticity and S contents and
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51 489 relatively small sizes. Proton and Cu^{2+} binding was studied by potentiometric titration and fit
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53 490 to the NICA-Donnan model for comparison to their binding to surface HSs. The results
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55 491 clearly indicate distinctive binding behaviors of H^+ and Cu^{2+} to the HHSs, likely caused by
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57 492 unique chemical nature of their functional groups, compared with those of surface HSs.
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3 493 These ions predominantly bind to chemically homogeneous carboxylic groups of the HHSs
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5 494 and to a lesser extent S containing group by mono-dentate coordination at low pH, which can
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8 495 explain the larger slopes of their isotherms and the smaller magnitude of Cu^{2+} binding than
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10 496 those of surface HSs. At neutral to alkaline pH H^+ dissociation from the HHSs is small, likely
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12 497 because the majority of the high affinity sites consist of alcoholic OH groups with larger pK_a .
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14 498 At the same pH range the slope of the Cu^{2+} isotherm to HHA becomes smaller and is close to
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16 499 the model calculations with the generic parameters, suggesting that binding mode could
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18 500 change to bi-dentate coordination with both carboxylic and alcoholic/phenolic groups. The
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20 501 XAS analyses further indicate that Cu^{2+} binds to HHA largely as Cu^{II} via O/N containing
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22 502 functional groups and to a lesser extent S containing functional groups. The generality of the
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24 503 results obtained in this study should be examined in the future by performing similar
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26 504 investigation for HSs isolated from various groundwater and the generic parameter sets
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28 505 applicable for groundwater HSs should be developed. Effects of co-existing metal ions,
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30 506 which are present in deep underground water at relatively high concentration, on ion binding
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32 507 to the HHSs are of great interest as they can alter electrostatic properties and secondary
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34 508 structures of the HHSs.
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3 626 **List of the table captions**
4

5 627 **Table 1** Elemental compositions and ^{13}C NMR carbon distributions of HHA and HFA.
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10 629 **Table 2** Optimized NICA-Donnan parameters for H^+ and Cu^{2+} binding to HHA and HFA. For
11
12 comparison the generic parameters for HA and FA, denoted as GHA and GFA, are given.¹⁸
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15 631
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17 632 **Table 3** EXAFS parameters optimized for the first coordination sphere of Cu^{2+} bound to
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19 HHA and PAHA at pH 4 and 7.
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3 634 **List of the figure captions**
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5 635 **Fig. 1** van Krevelen plot of the HHSs, the IHSS and JHSS reference and standard HSs, and
6 PAHA. The types of HS and origins are designated by different symbols and colors.
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12 638 **Fig. 2** Fractograms of the HHSs, the IHSS and JHSS reference or standard HSs, and PAHA
13 by FI-FFF, using 5 mM Tris buffer as effluent. Fractionated HS was measured by a UV/Vis
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17 640 detector at 255 nm. The fractograms of HAs are shown in (a) and those of FAs in (b).
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22 642 **Fig. 3.** Charge/pH curves (symbols) of HHA (a) and HFA (a) at the three salt concentrations.
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24 643 The solid curves represent the results of fitting to the NICA-Donnan model for the HHSs and
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26 644 the dotted curves correspond to the calculation of the NICA-Donnan model with the generic
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28 645 parameters in Table 2.¹⁸ The negative charges ($-q$) are plotted in the ordinates.
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34 647 **Fig. 4** Copper binding isotherms to HHA (a) at pH 4, 6, and 8 and to HFA (b) at pH 4 and 6,
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36 648 measured in the presence of 0.1 M NaClO₄. For pH 4, the results at 0.01 M NaClO₄ are also
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38 649 presented. The solid and dashed curves correspond to the results of the NICA-Donnan fitting
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40 650 at 0.1 and 0.01 M NaClO₄, respectively. The dotted and chained curves correspond to the
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42 651 calculation of the NICA-Donnan model with the generic parameters in Table 2 at 0.1 and
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44 652 0.01 M salt concentrations.¹⁸
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50 654 **Fig. 5** XANES spectra of Cu²⁺ bound to HHA and PAHA at pH 4 and 7 together with those
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52 655 of the Cu^{II} and Cu^I reference materials
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57 657 **Fig. 6** k^3 -weighted Cu K-edge EXAFS spectra (a) and the corresponding Fourier transform
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59 658 magnitude (b) of Cu²⁺ bound to HHA and PAHA at pH 4 and 7. The Cu²⁺ loading was 80
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3 659 mmol Cu/Kg HA. Solid curves correspond to the experimental results and dashed curves in
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6 660 the Fourier transform magnitude plots to the results of theoretical fitting of the first
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8 661 coordination sphere of Cu²⁺ (see the text for the details).
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663 **Table 1** Elemental compositions and ¹³C NMR carbon distributions of HHA and HFA.

	Elemental composition (%) ^a						¹³ C NMR (%) ^b				
	C	H	N	O	S	Ash	C _I	C _{II}	C _{III}	C _{IV}	C _V
HHA	62.29	6.40	3.36	25.44	2.51	N.D. ^c	3.0	12.7	26.4	17.9	40.6
HFA	60.23	6.84	2.06	29.00	1.87	N.D. ^c	4.2	13.9	21.4	15.5	45.6

664 a. Ash free basis.

665 b. C_I: carbonyl C (190 - 220 ppm), C_{II}: carboxyl C (165 - 190 ppm), C_{III}: aromatic C (110 -
666 165 ppm), C_{IV}: Methoxyl and carbohydrate C (48-110 ppm), C_V: aliphatic C (5 - 48 ppm).

667 c. Not detected.

668

669 **Table 2** Optimized NICA-Donnan parameters for H⁺ and Cu²⁺ binding to HHA and HFA.
 670 For comparison the generic parameters for HA and FA, denoted as GHA and GFA, are
 671 given.¹⁸

	HHA	HFA	GHA	GFA
q_0 (eq/Kg)	-0.64	-0.53	-	-
b	0.81	0.87	0.49	0.57
$Q_{\max 1, H}$ (eq/Kg)	4.38	5.64	3.15	5.88
p_1	<i>I^a</i>	<i>I^a</i>	0.62	0.59
$Q_{\max 2, H}$ (eq/Kg)	4.44	4.09	2.55	1.86
p_2	0.36	0.27	0.41	0.70
$\log \tilde{K}_{1, H}$	3.74	3.63	2.93	2.34
$n_{1, H}$	0.82	<i>I^a</i>	0.81	0.66
$\log \tilde{K}_{2, H}$	10.62	10.48	8.00	8.60
$n_{2, H}$	<i>I^a</i>	<i>I^a</i>	0.63	0.76
$\log \tilde{K}_{1, Cu}$	1.32	1.16	2.23	0.26
$n_{1, Cu}$	<i>I^a</i>	<i>I^a</i>	0.56	0.53
$\log \tilde{K}_{2, Cu}$	14.43	15.05	6.85	8.26
$n_{2, Cu}$	0.28	0.29	0.34	0.36
R^2 (H ⁺) ^b	0.9970	0.9975	-	-
R^2 (Cu ²⁺) ^b	0.9951	0.9876	-	-

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 673 a. The values in italic are constrained in physically meaningful ranges of the corresponding
 674 parameters (see the text of the SI for the detail).
 675 b. The correlation coefficients of the fitting for H⁺ and Cu²⁺.
 676

677 **Table 3** EXAFS parameters optimized for the first coordination sphere of Cu²⁺ bound to
 678 HHA and PAHA at pH 4 and 7.

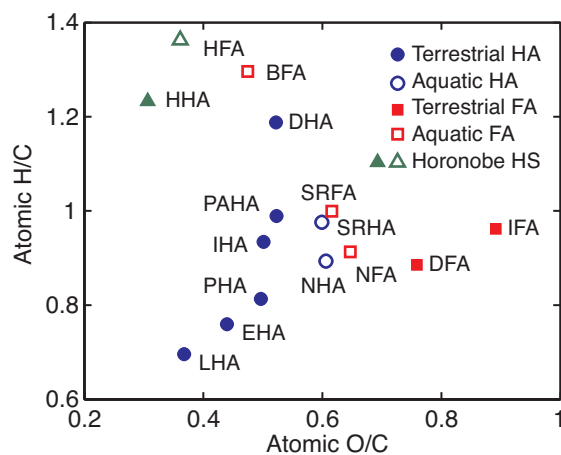
	First shell (Cu-O/N)			Second shell (Cu-S)			ΔE_0 (eV)
	<i>R</i>	CN ^a	σ^2 ^b	<i>R</i>	CN	σ^2 ^b	
	(Å)		(10 ³ Å ²)	(Å)		(10 ³ Å ²)	
HHA							
pH 4	1.95 ± 0.01	2.7 ± 0.6	3 ± 1	2.35 ± 0.04	0.4 ± 0.2	3 ± 1	1.53 ± 2.36
pH 7	1.95 ± 0.02	3.1 ± 0.9	4 ± 3	2.35 ± 0.16	0.1 ± 0.3	4 ± 3	1.52 ± 3.51
PAHA							
pH 4	1.94 ± 0.01	3.7 ± 0.8	5 ± 2	-	-	-	0.71 ± 2.80
pH 7	1.93 ± 0.01	3.6 ± 0.5	4 ± 1	-	-	-	1.19 ± 2.07

679 a. Coordination number.

680 b. Debye-Waller factor. The Debye-Waller factor of the Cu-S shell was set to be equal to
 681 that of the Cu-O/H shell.

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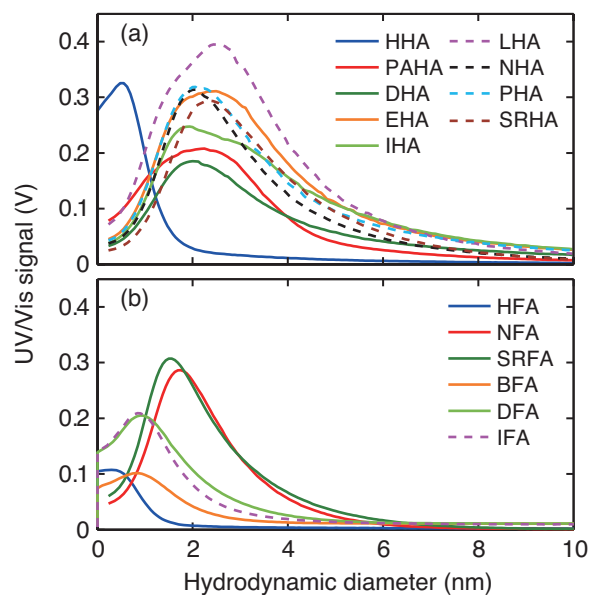
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686 PAHA. The types of HS and origins are designated by different symbols and colors.

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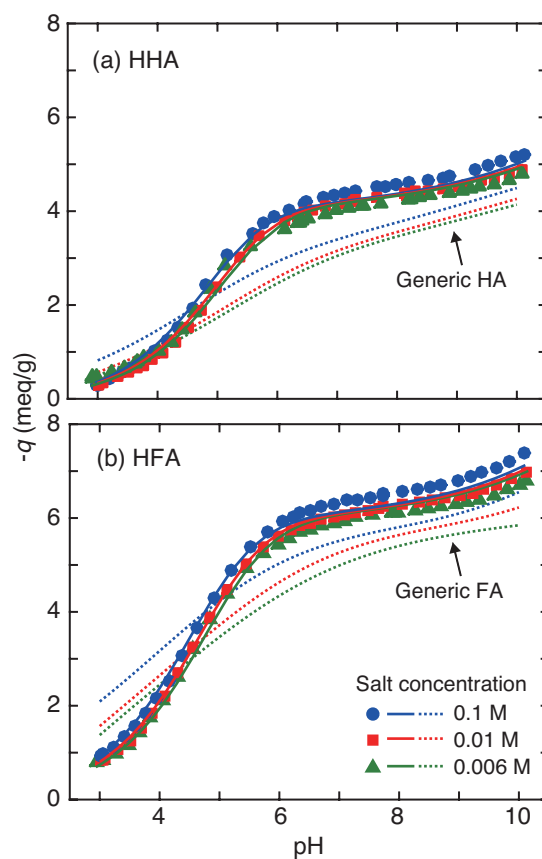


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689 **Fig. 2** Fractograms of the HHSs, the IHSS and JHSS reference and standard HSs, and
690 PAHA by FI-FFF, using 5 mM Tris buffer as effluent. Fractionated HS was measured by a
691 UV/Vis detector at 255 nm. The fractograms of HAs are shown in (a) and those of FAs in (b).

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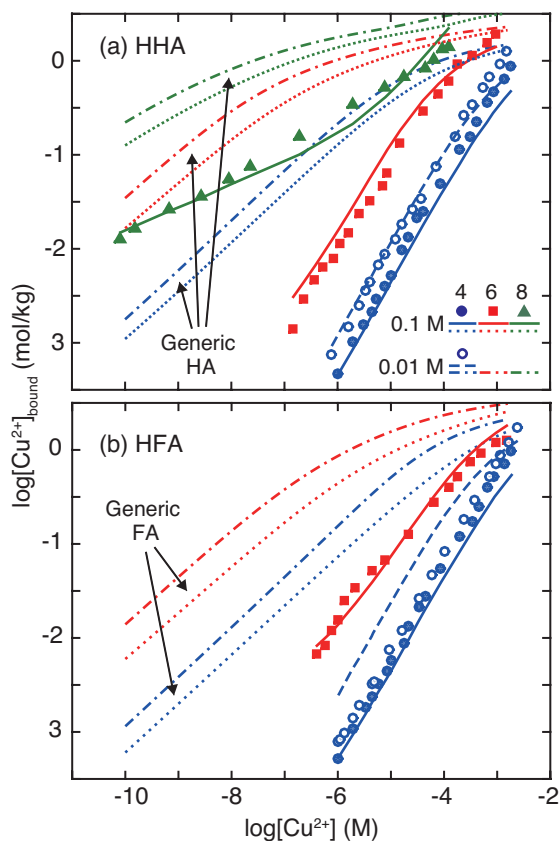


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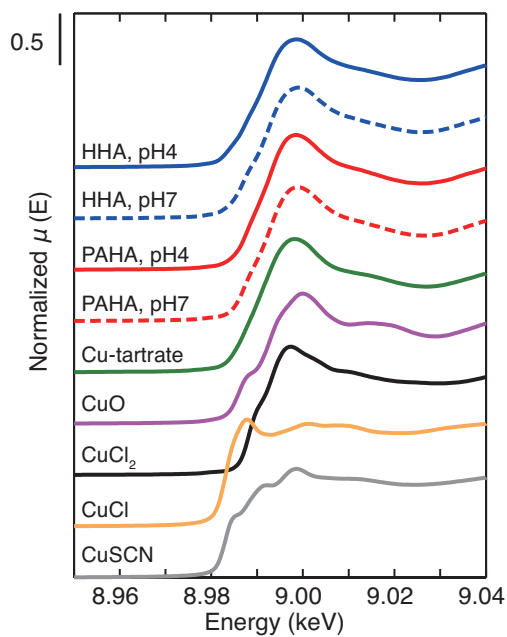
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701 **Fig. 4** Copper binding isotherms to HHA (a) at pH 4, 6, and 8 and to HFA (b) at pH 4 and 6,
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 704 at 0.1 and 0.01 M NaClO_4 , respectively. The dotted and chained curves correspond to the
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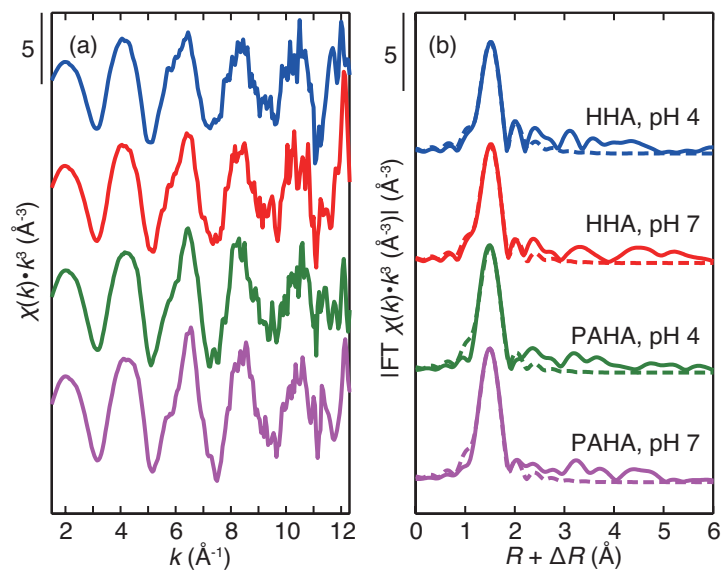


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