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**Diel Mercury Concentration Variations in a Mercury-
Impacted Stream**

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Environmental Significance Statement

Each component of the mercury (Hg) cycle in stream ecosystems may respond differently to the daily photocycle, across seasons, and influenced by watershed characteristics resulting in complex patterns in observed concentrations. We measured Hg and monomethylmercury (MMHg) concentrations and ancillary water quality parameters over multiple diel cycles in contrasting seasons at several locations along a Hg contaminated creek. Portions of the Hg cycle underwent rapid oscillations correlated with the daily photocycle (sometimes negatively) and were consistent within seasons and opposite in contrasting seasons. Macro-biotic (bioturbation), microbiological (methylation), and abiotic factors exerted control on the observed patterns. The results provide new insights into Hg cycling in contaminated stream ecosystems with potential implications for biotic receptors, monitoring, and remediation.

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11 **Diel Mercury Concentration Variations in a Mercury-Impacted Stream**
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ABSTRACT

Filtered and particulate mercury (Hg) and methylmercury (MMHg), and associated water chemistry parameters, were evaluated bi-hourly for several 30-hr periods during the summer and winter seasons at several distinct locations (downstream forested, midstream urban/suburban, upstream industrial) along a creek contaminated with high levels of inorganic Hg to determine if biogeochemical Hg and MMHg cycles respond to the daily photocycle. In summer particulate Hg and MMHg concentrations doubled overnight (excluding the upstream industrial site) concurrent with increases in turbidity and total suspended sediment; no such pattern was evident in winter. Seasonal and diel changes in the activity of macrobiota affecting the suspension of contaminated sediments are likely responsible for these patterns as other potential explanatory variables (e.g., instrument drift, pH, discharge) could not account for the range and timing of our observations. Diel patterns in filtered Hg (Hg_D) were significant only at locations and times of the year when channel shading was not present and daytime concentrations increased 22-89% above nighttime minima likely caused by direct and indirect photochemical reactions. Relationships between Hg_D and dissolved organic carbon (DOC) concentration or character were inconsistent between sites. Unlike Hg_D , there were significant diel patterns in filtered MMHg ($MMHg_D$) at all sites and times of year, with summer concentrations peaking in mid to late afternoon while the timing differed in winter, with concentrations peaking after sunset. Daily variability in $MMHg_D$ concentration ranged between 25-75%. The results imply key controls on net methylation occur within the stream or on the stream bed and include factors such as small-scale temperature changes in the water column and photosynthetic activity of stream biofilm. With respect to stream monitoring, results from this study indicate 1) consistent timing in stream Hg and MMHg sampling is required for accurate assessment of long-term trends, 2) *in situ* measurements of turbidity can be used to quantify diel dynamics of both particulate Hg and MMHg concentrations, and 3) *in situ* fluorescing dissolved organic matter (FDOM), a potential proxy for DOC, was not capable of resolving diel dynamics of filtered Hg or MMHg.

1. INTRODUCTION

Diel biogeochemical processes, including photosynthesis and respiration, flow variation, photochemical reactions, adsorption, desorption, and mineral precipitation and resultant chemical patterns in aquatic ecosystems, including those of dominant and trace anions and contaminants, have been reported in the literature for some time¹⁻⁸. Many of these reports have focused on lentic ecosystems with fewer studies conducted in lotic ecosystems. The physical setting and characteristics of streams and rivers (e.g., unidirectional turbulent flow, varying degrees of canopy cover, longitudinal variation in width to depth ratio) differ significantly from wetlands and lakes making it unclear which lessons from the latter will apply to the former particularly with respect to monomethylmercury (MMHg) concentration dynamics.

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5 36 Mercury (Hg) is a pollutant of global, regional, and local concern. In aquatic systems Hg can be
6 37 methylated by microorganisms creating MMHg which is much more toxic and bioaccumulative than its
7 38 non-methylated form. Hg has a complex cycle in aquatic environments where it is subject to a suite of
8 39 abiotic (e.g., sorption-desorption, complexation, precipitation-dissolution), photochemical (photo-
9 40 reduction and -oxidation), and microbially mediated reactions (methylation-demethylation, reduction, and
10 41 oxidation). Each of these reaction paths may have a unique sensitivity (in direction and extent) and
11 42 temporal response to the daily photocycle and co-occurring physical and chemical diel patterns. The
12 43 cumulative impact of these reactions could result from individual components of Hg (dissolved and
13 44 particulate fractions and different species) exhibiting unique diel patterns that could change throughout
14 45 the year or between locations with different chemical, physical, and land use characteristics.
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16 47 A few studies have specifically addressed daily patterns in total and filtered MMHg (MMHg_T and
17 48 MMHg_D, respectively) concentrations in aquatic systems, with the majority located in lakes or wetlands.
18 49 Observed patterns have been inconsistent for both types of systems, indicating a host of variables may
19 50 have a role in determining the resultant diel patterns. In a border wetland to the Great Salt Lake, MMHg_D
20 51 concentration declined during the day and increased at night ⁴. The authors attributed the decline in
21 52 MMHg during the day to photo-demethylation while the increase at night was believed to be due to
22 53 turnover of thermally stratified water that was below the zone of photodemethylation. In contrast, MMHg
23 54 concentrations increased during the day, correlated with solar radiation, in two lakes evaluated by
24 55 Siciliano et al. ⁹. The field observations and the results of *in situ* bottle incubations led the authors to
25 56 conclude that photoproduction of MMHg was occurring and was dependent on the structure and
26 57 concentration of DOM in the lakes. Differing diel patterns have been observed in wetlands, MMHg
27 58 concentrations doubled at night relative to daytime in an agricultural wetland planted with wild rice but
28 59 did not change over a daily cycle in an adjacent plot planted with white rice suggesting other differences
29 60 in ambient aquatic biogeochemical properties (e.g., DOM structure, microbial community differences
30 61 between crops) can moderate the influence of photodegradation on the net MMHg budget ⁵. In contrast to
31 62 the prior wetland study, Krabbenhoft et al. ¹⁰, reported diel trends in MMHg in the Everglades but
32 63 determined they were not correlated to the photocycle; no hypothesis was provided for the patterns. In
33 64 addition to field studies, consistent findings in bottle incubation studies that have illustrated a link
34 65 between MMHg photodegradation and light intensity ^{11, 12} providing support for the hypothesis that
35 66 photo-demethylation processes are likely to influence field concentrations but the variety of diel
36 67 concentration trends illustrate other processes may be more dominant in some systems.
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3 69 Nimick et al. ¹³, the only study to evaluate MMHg dynamics in a stream, reported increasing MMHg_D
4 70 concentrations during the day, peaking in early afternoon with overnight concentration minima, and
5 71 suggested patterns were likely driven by sunlight and temperature driven methylation. Due to a lack of
6 72 measurements in any other stream systems, the prevalence and/or consistency of this pattern and
7 73 dominating processes are unknown.

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11 75 Evaluations of diel patterns of unfiltered and dissolved Hg are rare in any aquatic system. Krabbenhoft et
12 76 al. ¹⁰, working in the Florida Everglades, reported reproducible diel cycles in unfiltered Hg concentrations
13 77 and its components, dissolved gaseous Hg (DGM; Hg⁰) and reactive Hg (i.e., SnCl₂ reducible). All Hg
14 78 fractions were linked to the daily photocycle, increasing during the day, and the pattern was attributed to
15 79 photolysis of chromophores on the surface of a solid substrate (e.g., the periphyton mat) giving rise to
16 80 destabilization of sorbed mercury and net desorption during daylight. Nimick et al. ¹³, is the first and only
17 81 other study to evaluate diel unfiltered Hg and Hg_D dynamics in stream systems. In a study of two streams
18 82 affected by mining and geothermal discharge, they found distinct dynamics with one site having no diel
19 83 patterns and one having increased daytime concentrations of unfiltered Hg and decreasing daytime
20 84 concentrations of Hg_D. Hg increases tracked suspended sediment, and the pattern of Hg_D concentrations
21 85 were attributed to either adsorption-desorption of Hg²⁺ or reduction of Hg(II) to Hg⁰ and subsequent
22 86 evasion of Hg⁰. The prevalence and/or consistency of these patterns and dominating processes in stream
23 87 systems are unknown.

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27 89 The inconsistent findings from lake and wetland studies and lack of studies in streams, with respect to
28 90 diel variability in Hg and MMHg, illustrate a knowledge deficit of the dominant controls and resultant
29 91 patterns tied to the daily photocycle. Furthermore, the diel evaluations that have been conducted have
30 92 focused on a single location or time of year and were not evaluated within the context of Hg dynamics
31 93 that occur over longer time periods (monthly/annual).

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35 95 Our objectives for this study conducted in a well characterized Hg-impacted stream were to (i) quantify
36 96 diel patterns of particulate and dissolved Hg and MMHg and associated water quality parameters at one
37 97 site in distinct seasons (winter and summer), at multiple sites in the same season (summer), and over
38 98 several years at the same site in the same season, (ii) determine if diel patterns in Hg and/or MMHg
39 99 fractions are present at distinct locations and times of year, and if they are related to the daily photocycle,
40 100 (iii) evaluate the relationship between any observed diel patterns in particulate and dissolved Hg and/or
41 101 MMHg to water quality parameters, specifically those associated with co-transport (e.g., total suspended
42 102 solids (TSS), DOC). We also evaluate the potential for *in situ* measurements of turbidity and Fluorescing
43 103 Dissolved Organic Matter (FDOM) to be used as surrogates for TSS and particulate Hg and MMHg and

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3 104 DOC and dissolved Hg and MMHg, respectively. These analyses will be used to gain insight into the
4 105 likely mechanisms responsible for observed diel variability of Hg and MMHg in this stream ecosystem
5 106 and quantify the importance of processes that operate on a short-term timeframe relative to long-term
6 107 annual variability.
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11 109 2. METHODS

12 13 110 2.1 SITE DESCRIPTION

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15 111 East Fork Poplar Creek (EFPC) is a fourth-order stream in the Valley and Ridge Province of the southern
16 112 Appalachians in eastern Tennessee, USA. EFPC has been heavily studied and many detailed summaries
17 113 of its history and characteristics can be found in the literature¹⁴⁻¹⁷. The creek and surrounding environs
18 114 were contaminated with Hg due to activities at the Y-12 National Security Complex located at the creek
19 115 headwaters. The climate is humid subtropical with mean annual temperature 14.5°C and precipitation
20 116 evenly distributed through the year (mean annual ~137 cm)¹⁸. Streamflow can exhibit strong seasonality
21 117 due to high rates of evapotranspiration in summer^{18,19}. During winter there is no canopy cover although
22 118 there are localized areas of heavy shading due to high, steep creek banks and as the creek flows at base of
23 119 the north facing slopes of prominent ridges. During summer heavy canopy cover shades much of the
24 120 creek channel. From its headwaters, the creek meanders 26 km through commercial, residential, open-
25 121 land and woodland sections of the city of Oak Ridge, TN to its confluence with Poplar Creek (Figure 1).
26 122 Our study sites are located 5.4, 16.2 and 23.4 km upstream of the mouth of the creek (hereinafter
27 123 referred to as EFK 5.4, EFK 16.2, and EFK 23.4; Table S.1). Comparison of aerial photographs taken in
28 124 February versus June indicates that in summer 18%, >95% and 90% of the channel is covered by canopy
29 125 at EFK 23.4, EFK 16.2, and EFK 5.4, respectively (Supporting Information). The EFK 5.4 site is 3 km
30 126 upstream of the confluence with Bear Creek, the major tributary to EFPC making EFPC a third-order
31 127 stream at our farthest downstream sampling site. The Oak Ridge wastewater treatment facility (ORWTF)
32 128 discharges treated effluent at EFK 13.5. This effluent is a source of nutrients (nitrate, phosphate), organic
33 129 carbon, and chloride to EFPC but makes insignificant contributions of Hg and MMHg to the creek.
34 130 Between 2011-2018 average (s.d.) DOC and SUVA₂₅₄ values of the effluent were 5.5 mg·L⁻¹ (1.6, n = 51)
35 131 and 2.4 L·mgC⁻¹·m⁻¹ (0.7, n = 44), respectively, compared to samples collected 300 m upstream of the
36 132 outfall DOC = 1.9 mg·L⁻¹ (1.6, n = 32), SUVA₂₅₄ = 2.6 L·mgC⁻¹·m⁻¹ (0.8, n = 29) and 80 m downstream
37 133 of the outfall, approximately 8-10 creek widths, DOC = 3.0 mg·L⁻¹ (1.3, n = 26), SUVA₂₅₄ = 2.5 L·mgC⁻¹·m⁻¹
38 134 (0.7, n = 26). Scheduled backwash operations and variable influent result in lower overnight
39 135 effluent discharge rates but only mean daily discharge data are available. Based on rating curves
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3 137 be \geq 28 hours, from EFK 16.2 to EFK 5.4 to be \sim 24 hours, and from the ORWTF outfall to EFK 5.4 to be
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5 138 \sim 9 hours.

6 7 139 **2.2 FIELD METHODS**

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9 140 Most of the field work described in this paper occurred between 28 August 2013 and 18 September 2015.
10 141 All times given are eastern standard time (EST = UTC – 5:00) using a 24-hour clock. Three diel
11 142 sampling cycles were conducted under baseflow conditions. Two were conducted in summer (Su2013 =
12 143 28-29 August 2013; Su2015 = 17-18 September 2015) and one in winter (W2014 = 10-11 February
13 144 2014). During the Su2013 and W2014 campaigns samples were collected at EFK 5.4 to evaluate
14 145 distinctions between seasons. For the Su2015 campaign samples were collected at EFK 5.4 (forested),
15 146 EFK 16.2 (urban/suburban), and EFK 23.4 (industrial) to evaluate commonalities and differences at
16 147 multiple locations within the same channel. During each campaign, samples were collected every 2 hours
17 148 for 30 hours beginning at 0900 h on the first day and continuing until 1500 h on the following day.
18 149 Samples were collected using a combination of manually sampling into a new 250 ml glycol-modified
19 150 polyethylene terephthalate (PETG) bottle or automated sampling with a Teledyne ISCO® sampler
20 151 (typically overnight periods) retrofitted with a strainer and sample tubing made from Teflon® and
21 152 collected into Isco ProPak™ disposable sample bags made of low-density polyethylene (LDPE). We
22 153 previously demonstrated that this autosampling method does not compromise sample integrity for our
23 154 creek water²⁰. Grab samples were collected by wading into the middle of the stream, triple rinsing the
24 155 bottle with creek water, then filling and sealing the bottle for transport to the lab.
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36 157 In addition to the diel campaigns described, grab samples were collected at EFK 5.4 each month, twice a
37 158 day, between the Su2013 and W2014 campaigns to track the timing of the transition between any
38 159 differences observed in the summer and winter dynamics. During these monthly samplings, one sample
39 160 was collected in the morning (0800 h -1000 h) and another in the afternoon (1245 h – 1500 h). A final
40 161 diel campaign of reduced complexity was conducted over a 48-hour period from 14-16 August 2018
41 162 (Su2018) at EFK 5.4 in which grab samples were collected from the surface water and from a dedicated
42 163 push-point probe with a screened interval 30 cm below the sediment-water contact, to evaluate the
43 164 potential influence of pore water dynamics on stream concentrations. Samples were collected every 3
44 165 hours beginning at 0820 h on the first day. During these latter two efforts (monthly sampling from
45 166 Su2013 to W2014 and Su2018) samples were analyzed for dissolved and total Hg and dissolved and total
46 167 MMHg only.
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55 169 Local meteorological data (air temperature, solar radiation) were obtained from a tower located 4.5 km to
56 170 the southwest of EFK 5.4 (Tower K on the Oak Ridge Reservation; Table S.2). Total diffuse light

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3 171 intensity at sampling sites was measured every 5 minutes using an Onset Inc. HOBO Pendant
4 172 Temperature/Light Data Logger (model UA-002-64) mounted 1.2 meters above the surface of the water.
5 173 These instruments have a response range from 0 to 320,000 lux over a wavelength range from 150 to
6 174 1200 nm. Photosynthetically active radiation (PAR; $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ in the wavelength range 400-700 nm)
7 175 was measured every 15 minutes with a LI-COR Quantum sensor and LI-COR Model LI-1000 data logger.
8 176 The sensor was mounted 30 centimeters above the water surface near the center of the creek channel.
9 177 Multiparameter sondes were installed adjacent to the pressure transducer at each site and measured and
10 178 recorded pH, dissolved oxygen, specific conductivity, turbidity (except at EFK 23.4 for Su2015), water
11 179 temperature, and fluorescent dissolved organic matter (FDOM; Su2013 and W2014 only) at 15-min
12 180 intervals. Discharge was measured at 15-minute intervals (i) throughout the entire study at EFK 5.4, and
13 181 (ii) from spring 2015 throughout the study at EFK 16.2. Discharge for EFK 23.4 at 6-minute intervals
14 182 was provided courtesy of Y-12 environmental compliance personnel. Details on discharge measurements
15 183 and calculations are provided in the Supplemental Information.

184 **2.3 SAMPLE HANDLING AND ANALYTICAL METHODS**

185 A list of measured parameters is provided in Table S.2. Immediately after collection, samples were
186 transported from the field site to Oak Ridge National Laboratory (15 min drive). An aliquot of the sample
187 was retained for analysis of unfiltered Hg and MMHg (Hg_T and MMHg_T , respectively) and the remaining
188 sample was filtered through either a 0.2 μm analytical filter unit for Hg and MMHg analysis or a 0.2 μm
189 syringe filter (both polyethersulfone (PES) membranes) for dissolved organic carbon (DOC), UV-vis,
190 anion, and cation analysis. Previous work has shown that in EFPC MMHg passing a 0.2 μm filter also
191 passes a 3 kD ultrafilter²¹. Samples collected for total suspended solids (TSS) and particulate organic
192 carbon (POC) analysis were filtered through a tared and pre-baked Whatman GFF glass fiber filter with a
193 0.7 μm particle retention and dried in an oven at 100°C to a constant weight. Unfiltered and filtered Hg
194 and MMHg samples were preserved to 0.5% (v/v) HCl and DOC samples were preserved to 0.1% (v/v)
195 HCl. Trace metal grade HCl was used for sample preservation. Samples for cation analysis were
196 preserved to 0.5% (v/v) HNO_3 (trace metal grade). Filtered samples for anion and nutrient analysis were
197 held in new PETG or polyethylene containers. All samples were refrigerated at 4°C or frozen
198 (ammonium and soluble reactive phosphorous (SRP)) in the dark until subsequent analysis. Filtered Hg
199 and MMHg samples are herein referred to as dissolved Hg (Hg_D) and MMHg (MMHg_D).

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201 Analysis of Hg samples was conducted using a Hg purge and trap system (Brooks Rand MERX).
202 Bromine monochloride was added to all Hg samples a minimum of 24 h before analysis. Hydroxylamine
203 and stannous chloride were added to the samples and the Hg^0 produced was purged from solution and
204 trapped onto gold-coated sand analytical traps²². The traps were subsequently heated to release the Hg

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3 205 which was detected by the Cold Vapor Atomic Fluorescence Spectrophotometer (CVAFS). Ambient
4 206 MMHg was analyzed using modifications of EPA method 1630²³ which involves the distillation of the
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6 207 water sample followed by ethylation, purge and trap onto Tenax traps, gas chromatographic separation of
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8 208 the Hg species and detection by ICP-MS. This analysis was performed with a Brooks Rand MERX
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10 209 MMHg instrument coupled with a Perkin Elmer Elan-DRC ICP-MS. An internal standard (MM²⁰⁰Hg)
11 210 was added to samples prior to distillation and this isotope was used to quantify MMHg concentrations²⁴.
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13 211 Additionally, we verified there was no artifact MMHg formation during processing and analysis of water
14 212 samples by adding ²⁰¹Hg to select samples and monitoring the absence of MM²⁰¹Hg. Particulate Hg and
15 213 MMHg (Hg_p and MMHg_p, respectively) were calculated as unfiltered concentration minus dissolved
16 214 concentration.
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21 216 DOC concentrations were measured using high-temperature platinum-catalyzed combustion followed by
22 217 infrared detection of CO₂ (Shimadzu TOC-5000A or Shimadzu TOC-L). UV-visible spectra were
23 218 collected at 1-nm interval and 0.5 second exposure time from 190-1100 nm wavelength with an HP 8453
24 219 spectrophotometer using a 1-cm path length quartz cuvette. Specific ultraviolet absorbance at 254 nm
25 220 (SUVA₂₅₄), a dissolved organic matter (DOM) composition indicator, was also computed. SUVA₂₅₄ was
26 221 calculated as the UV absorbance at 254 nm (m⁻¹) divided by the DOC concentration (mg·L⁻¹) and
27 222 reported in units of L·mg·C⁻¹·m⁻¹. Total iron concentrations were generally less than 0.1 mg·L⁻¹ and
28 223 were not high enough to cause significant interference²⁵.
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35 225 Dissolved anions (Cl⁻, NO₃⁻, SO₄²⁻, PO₄³⁻) and cations (Na, Mg, Ca, K, Al, Fe, Mn, Sr, Ba, U) were
36 226 analyzed by ion chromatography (IC) (Dionex DX-120, Sunnyville, CA, USA) and inductively coupled
37 227 plasma-mass spectrometer (ICP-MS, Perkin-Elmer, Waltham, MA, USA), respectively. POC was
38 228 determined by measuring the percent of carbon released from the total suspended solids sample by dry
39 229 combustion using an elemental analyzer (LECO-CNS-2000, St. Joseph, MI).
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44 231 Samples for DGM analysis were collected in 1-liter acid-washed glass bottles. Bottles were triple rinsed
45 232 with creek water then filled with no headspace, sealed and transported to the lab for immediate
46 233 processing. Once in the lab, approximately 750 mL of the sample was transferred into a borosilicate glass
47 234 washing bottle and the water was sparged with Hg-free N₂ for 2 hours at a flow rate of 0.75 mL·min⁻¹.
48 235 The off gas was directed into acidic potassium permanganate traps (0.06 M KMnO₄ in 10% H₂SO₄).
49 236 After sparging was completed, the sample was weighed to determine volume and the KMnO₄ traps
50 237 preserved for subsequent total Hg analysis by cold vapor atomic absorption spectroscopy (CVAAS).
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239 Details on Quality Assurance/Quality Control, calculation of solid-water partitioning coefficients (K_{sw}),
240 and potential outlier detection are provided in Supplemental Information.

241 2.4 STATISTICS AND COMPUTATIONS

242 Statistical analyses were conducted in R ²⁶. Potential outliers were identified using modified z-scores
243 (Supplemental Information, Section 4). For all statistical tests, results were considered significant by
244 adopting an *a priori* Type I error rate of 5%. Reported correlations are Spearman's Rank Correlation
245 coefficient (ρ), a nonparametric measure of monotonic correlation between two variables. The
246 magnitude of diel concentration variations, expressed as a percentage, was calculated by dividing the
247 range of observed values by the minimum value, excluding potential outliers. Determination of the
248 dependence of concentration on sampling time was made using the coefficient (ξ) introduced by
249 Chatterjee²⁷ which is very efficient at determining associations that are non-monotonic or oscillating
250 even with noisy data (Supporting Information; Figure S.1). Estimates of ξ were calculated using the R
251 package XICOR²⁸.

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3. RESULTS

254 3.1 GENERAL CREEK CONDITIONS DURING SAMPLING

255 Each sampling campaign occurred under generally clear skies five to seven days after the most recent
256 high flow event (Figure S.2). The combination of high and steep creek banks and summer canopy cover
257 limited the light intensity reaching the creek during all but the midday hours. Creek discharge during
258 each sampling campaign was typical of seasonal baseflow conditions (Table S.4; Figure S.3). Lower flow
259 at EFK 5.4 during the Su2015 campaign compared to Su2013 is attributed to drier antecedent conditions.
260 The antecedent precipitation index²⁹ based on daily rainfall totals was 83.6 mm and 55 mm in 2013 and
261 2015, respectively. These estimates include a decay function that assigns greater weight to more recent
262 precipitation (period of record begins 1 April 2001). Mean stream discharge during W2014 was higher
263 than during the summer campaigns due to lower evapotranspiration. Discharge from the ORWTF
264 constituted 12.7%, 10.5%, and 15.1% of the mean daily discharge at EFK 5.4 during the Su2013, W2014,
265 and Su2015 campaigns, respectively (Table S.4).

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267 Water temperature increased after sunrise, peaked in late afternoon, and decreased overnight (Figure 2A).
268 Dissolved oxygen (DO) concentration exhibited a diel pattern correlated with the daily photocycle
269 consistent with photosynthesis occurring in the creek (Figure 2B). pH followed a pattern similar to DO in
270 accordance with the canonical relationship between photosynthesis, O_2 production, and CO_2 consumption
271 driving increased pH (Figure 2C). The Su2013 sampling occurred during a trend of decreasing pH (mean

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3 272 daily pH from 26–31 August = 7.74, 7.68, 7.48, 7.26, 7.25, 7.22) during which the relationship between
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5 273 pH and DO was less clear. Temperature, pH, and DO show very strong dependence on sampling time (ξ
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7 274 > 0.89) across all sampling campaigns and sites with oscillations correlated to the photocycle (Table S.5).
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9 275 Across sampling campaigns and sites, there was no clear diel pattern in specific conductance.

10 276 **3.2 MAJOR CATIONS AND ANIONS**

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12 277 The concentrations of the major cations (Ca, Mg, Na, K) and anions (Cl^- , NO_3^- , SO_4^{2-}) generally had low
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14 278 variability and did not have strong temporal dependence in sync with the daily photocycle (Figures S.4
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16 279 and S.5). Concentration dependence on sampling time (Table S.5) was typically due to monotonic trends
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18 280 rather than cyclic variations. Exceptions to this are seen in the concentration histories for Na, K, Cl^- , and
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20 281 NO_3^- during the Su2015 campaign at EFK 5.4. The concentration of these four constituents declined then
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22 282 rebounded in sync and all these changes occurred during a ten-hour overnight period and thus were
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24 283 disconnected from the daily photocycle.

25 284 **3.3 PARTICULATE Hg AND MMHg AND ASSOCIATED PARAMETERS (TURBIDITY, TSS, POC)**

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28 286 In the Su2013 and Su2015 campaigns the concentration of particulate Hg and MMHg, expressed on a per
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30 287 volume basis ($\text{Hg}_{\text{P.ngL}}$ and $\text{MMHg}_{\text{P.ngL}}$, respectively) increased overnight at EFK 16.2 and EFK 5.4
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32 288 coincident with overnight increases in turbidity and TSS (Figure 3 and Figure 4) leading to significant ξ
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34 289 values and strong positive correlations between TSS and $\text{Hg}_{\text{P.ngL}}$ and $\text{MMHg}_{\text{P.ngL}}$ (Tables S.5 and S.6).
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36 290 Overnight increases in $\text{Hg}_{\text{P.ngL}}$ and $\text{MMHg}_{\text{P.ngL}}$ ranged between 73-200% and 105-500%, respectively,
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38 291 relative to daytime minima (potential outliers removed). Overnight increases in TSS and turbidity ranged
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40 292 between 119-194% and 122-495%, respectively. Turbidity was dependent on sampling time during the
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42 293 winter event, but the effect was weak, the estimate of ξ was imprecise (0.18 ± 0.06 ; Table S.5), and the
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44 294 dependence was a linear decrease rather than a diel oscillation. Similar diel oscillations in $\text{Hg}_{\text{P.ngL}}$ and
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46 295 $\text{MMHg}_{\text{P.ngL}}$ were observed in the Su2018 48-hour event although estimates of ξ were not significant
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48 296 (Figure S.6A and D; Table S.5). In contrast, there were no discernible diel patterns in $\text{Hg}_{\text{P.ngL}}$,
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50 297 $\text{MMHg}_{\text{P.ngL}}$, or TSS at EFK 23.4 or during the winter sampling at EFK 5.4.

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53 299 Because of mercury's strong affinity for particles (particulate Hg accounted for 70-88%, 72%, and 41 %
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55 300 of Hg_T at EFK 5.4, EFK 16.2, and EFK 23.4, respectively)^{30, 31}, we calculated the particulate Hg to TSS
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57 301 ratio expressed in units of $\text{mg} \cdot \text{kg}^{-1}$ ($\text{Hg}_{\text{P.ngL}} \cdot \text{TSS}$) at each sampling time for each campaign. This ratio
58
59 302 remained relatively constant over time in four of the five sets of observations. Sampling time dependence
60
303 with diel oscillation was seen in Su2013 at EFK 5.4 ($\xi = 0.41$, $p = 2.5 \times 10^{-3}$; Figure S.7A and C). The
304
sampling time dependence for $\text{Hg}_{\text{P.ngL}} \cdot \text{TSS}$ at EFK 5.4 and EFK 16.2 in Su2015 was due to monotonic

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3 305 trends rather than daily oscillations, the effects were weak to moderate and estimates of ξ were imprecise.
4
5 306 Compared to Hg_T , a smaller percentage of MMHg_T was particle associated (21-29%, 38%, and 25% of
6
7 307 MMHg_T at EFK 5.4, EFK 16.2, and EFK 23.4, respectively). The particulate MMHg , to TSS ratio
8
9 308 ($\text{MMHg}_{\text{p.ngL}}:\text{TSS}$) was independent of sampling time for all events (Figure S.7B and D; Table S.5).

10 309
11 310 More generally, $\text{Hg}_{\text{p.ngL}}$ was strongly positively correlated with TSS in four of the five cases (Figure S.8A
12
13 311 and C, Table S.6), the one exception being Su2015 at EFK 23.4 ($\rho = -0.0944$, $p = 0.719$). The slopes of
14
15 312 the $\text{Hg}_{\text{p.ngL}}$ versus TSS regression lines were the same for those cases with significant correlation, i.e.,
16
17 313 there was no evidence for a sampling event \times TSS interaction effect (analysis of covariance, $F_{3,75} = 1.87$,
18
19 314 $p = 0.14$). For each $1 \text{ mg}\cdot\text{L}^{-1}$ increase in TSS, $\text{Hg}_{\text{p.ngL}}$ increased by $9.3 \pm 1.2 \text{ ng}\cdot\text{L}^{-1}$. Similarly, there was
20
21 315 no evidence for a sampling event \times TSS interaction effect for the $\text{MMHg}_{\text{p.ngL}}$ data regardless of the
22
23 316 potential outlier value at EFK 16.2 ($F_{2,53} = 0.783$, $p = 0.46$). For each $1 \text{ mg}\cdot\text{L}^{-1}$ increase in TSS,
24
25 317 $\text{MMHg}_{\text{p.ngL}}$ increased by $0.017 \pm 0.004 \text{ ng}\cdot\text{L}^{-1}$ (0.0014 ± 0.003 excluding outlier).

26 318
27 319 Particulate organic carbon (POC, $\text{mg}\cdot\text{L}^{-1}$) was measured for the Su2013 and W2014 campaigns only and
28
29 320 for each the POC showed a diel oscillation pattern similar to that measured for TSS with daytime minima
30
31 321 and overnight maxima (Figure S.9A). The POC concentration dependence on sampling time was stronger
32
33 322 in summer than winter ($\xi = 0.59 \pm 0.13$ and 0.32 ± 0.13 in summer and winter, respectively; Table S.5).
34
35 323 In contrast, the percent organic C on suspended solids (percent OC per mg TSS) was independent of
36
37 324 sampling time (Figure S.9B) although values for winter samples (mean \pm sd = 6.4 ± 0.8) were
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39 325 significantly higher than those from summer (5.1 ± 0.5 ; $p = 3.32 \times 10^{-8}$; Welch's two-sample t-test;
40
41 326 including potential outlier).

39 327 **3.4 DISSOLVED Hg AND MMHg AND ASSOCIATED PARAMETERS (DISSOLVED** 40 328 **ORGANIC CARBON)**

42 329 Dissolved Hg (Hg_D) concentration showed diel oscillations at EFK 5.4 in W2014 and EFK 23.4 in
43
44 330 Su2015 ($\xi = 0.47 \pm 0.13$, $p = 4.5 \times 10^{-4}$ and 0.53 ± 0.15 , $p = 2 \times 10^{-3}$, respectively) in which concentration
45
46 331 increased during the day with mid to late afternoon maxima and overnight minima (Figure 5B and E).
47
48 332 Dissolved MMHg concentration showed diel oscillation in each campaign. During the summer
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50 333 campaigns at each location MMHg_D concentrations increased during the day with maxima in mid to late
51
52 334 afternoon and declined overnight. At EFK 23.4 in Su2015 MMHg_D concentrations were relatively
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54 335 constant during the day and greater than constant nighttime concentrations. Similar diel oscillations were
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56 336 seen in the surface water samples during the 48-hour campaign in Su2018. Dissolved MMHg
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58 337 concentrations in pore water were, on average, $7\times$ higher than in surface water with no apparent diel

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3 338 oscillations (Figures S.6E and F). In contrast, for the winter sampling the highest MMHg_D was measured
4
5 339 well after sunset, approximately six hours later than the daytime maxima measured in summer.

6 340
7
8 341 For most campaigns and locations, no diel oscillations were seen for DOC concentration or composition
9
10 342 as estimated from SUVA₂₅₄ (Figure S.10). However, at EFK 23.4 in Su2015 these two parameters show
11
12 343 opposing diel oscillation with DOC increasing during the day and declining overnight whereas SUVA₂₅₄
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14 344 values were lower during the day and higher at night. FDOM (measured in Su2013 and W2014 only)
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16 345 fluctuated over the sampling period in Su2013 with no apparent diel pattern. A strong diel pattern in
17
18 346 FDOM was seen in W2014 with values steadily decreasing during daylight hours and increasing
19
20 347 overnight (Figure S.11). The relationship between FDOM and DOC, SUVA₂₅₄, Hg_D, and MMHg_D varied
21
22 348 from summer to winter (Table S.8). For example, in Su2013 there were no significant correlations
23
24 349 between FDOM and the remaining four parameters regardless of outlier exclusion. In W2014, there was
25
26 350 a moderately strong positive correlation between FDOM and DOC and strong to very strong negative
27
28 351 correlations between FDOM and MMHg_D and Hg_D. A more extensive data set, spanning multiple
29
30 352 seasons and broader range of parameter values would be needed to evaluate more rigorously the potential
31
32 353 seasonal and extreme value effects on the use of FDOM as a reliable proxy variable for DOC, Hg_D, and
33
34 354 MMHg_D. No correlation was observed between FDOM and SUVA₂₅₄ in either Su2013 or W2014
35
36 355 regardless of outlier exclusion.

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38 356
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40 357 Because of the reported tight coupling of dissolved Hg and MMHg with DOC³²⁻³⁴ we examined the
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42 358 Hg_D:DOC and MMHg_D:DOC ratio over the course of each sampling campaign. DOC concentrations
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44 359 varied over time but generally did not show diel oscillations. Hg_D concentrations varied over time with
45
46 360 diel oscillations in two cases (W2014 at EFK 5.4 and Su2015 at EFK 23.4). MMHg_D concentration
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48 361 varied over time with diel oscillations for all events and locations. Nevertheless, the timing and
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50 362 magnitude of the DOC concentration variations overwhelmed the Hg_D and MMHg_D patterns and
51
52 363 significant diel oscillation in Hg_D:DOC was apparent only in W2014 at EFK 5.4 and for MMHg_D:DOC in
53
54 364 Su2015 at EFK 5.4 (Figure S.12; Table S.5). Additionally, significant positive correlations between DOC
55
56 365 and Hg_D were found for only two cases, Su2013 at EFK 5.4 and Su2015 at EFK 23.4 (Table S.7; Figure
57
58 366 S.13). Interestingly, these two parameters were strongly negatively correlated in Su2015 at EFK 16.2
59
60 367 ($\rho = -0.719$, $p = 7.67 \times 10^{-4}$). MMHg_D and DOC were positively correlated only for Su2015 at EFK 23.4
368 and uncorrelated for all other events and locations.

369 **3.5 DISSOLVED GASEOUS MERCURY (DGM)**

370 Dissolved gaseous mercury concentrations were lower in samples collected in early morning or during
371 night and higher in samples collected in mid-afternoon (Figure 6) with a moderately strong positive

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3 372 correlation between PAR and DGM ($\rho = 0.47$, $p = 0.05$) using average PAR value for the 2-hour period
4 373 preceding collection of the DGM sample. Samples for DGM analysis were collected less frequently than
5 374 for the other analytes precluding meaningful estimation of the ξ statistic.

9 375 **3.6 TEMPORAL MMHg_D CONCENTRATION PATTERNS IN MORNING VERSUS** 10 376 **AFTERNOON SAMPLES**

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12 377 Between the Su2013 and W2014 30-hour sampling campaigns, samples were collected once per month at
13 378 two different times during the day, once in the morning and again in the afternoon. During summer and
14 379 early autumn, the MMHg_D concentration was higher in the afternoon than the morning. By mid-autumn
15 380 there was no difference between samples at different times of day and this lack of difference persisted
16 381 into winter (Figure 7). These observations are consistent with the February 2014 diel sampling in which
17 382 there were no differences in MMHg_D concentration for samples collected during daylight.

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20 383
21 384 A summary of the primary results is provided in Table 1. We turn now to a discussion of the results and
22 385 the mechanisms proposed.

23 386 24 387 **4. DISCUSSION**

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27 388 Because of the estimated travel times between sites (Methods; EFK 23.4 to EFK 16.2 ≥ 28 hours; EFK
28 389 16.2 to EFK 5.4 ~ 24 hours; ORWTF to EFK 5.4 ~ 9 hours), diel concentration variations at EFK 16.2
29 390 likely reflect or integrate reach-scale processes independent of changes observed at EFK 23.4. Similarly,
30 391 concentration variations at EFK 5.4 were likely independent of changes at EFK 16.2 but were affected by
31 392 variable discharge from the ORWTF.

32 393 **4.1 MAJOR IONS**

33 394 Among the major cations and anions, sodium, potassium, chloride, and nitrate each had a synchronized
34 395 pattern of concentration dips and recovery during overnight hours at EFK 5.4 in Su2015. As there are no
35 396 geogenic sources of these solutes in East Fork valley, their presence in EFPC reflects anthropogenic
36 397 influences. The overnight concentration dynamics are likely the result of decreases in discharge of treated
37 398 effluent from the ORWTF 8.1 kilometers upstream of our sampling location at EFK 5.4. A similar
38 399 pattern was not observed for the other sampling campaigns at EFK 5.4 potentially because streamflow
39 400 was higher and therefore the influence of changes in upstream effluent flow would have a diminished
40 401 impact on stream concentrations. A long record of sampling this effluent has shown that it is not a source
41 402 of either Hg or MMHg to EFPC and daily fluctuations in effluent discharge are not expected to affect the
42 403 diel Hg and MMHg patterns. The dominant ions with geogenic sources in the watershed (calcium,
43 404 magnesium) showed no similar fluctuations over the sampling periods.

4.2 CONTROLS ON PARTICULATE Hg AND MMHg CONCENTRATIONS, TURBIDITY AND TSS

Over the course of several diel sampling campaigns conducted in summer months, $Hg_{p.ngL}$, $MMHg_{p.ngL}$, and POC concentrations oscillated in sync with corresponding changes in TSS and turbidity at two sites (forested and urban/suburban); these oscillations were counter to the daily photocycle. No such patterns were evident during winter or at the farthest upstream site. In contrast, those summertime patterns in particulate Hg, MMHg, and POC do not appear when normalized to TSS concentration. This suggests that overnight increases in TSS, turbidity, POC, $Hg_{p.ngL}$, and $MMHg_{p.ngL}$ were due to the (re)suspension of the same solids contributing to TSS load during the day and were most likely caused by bioturbation due to the activity of macrobiota in the creek³⁵⁻³⁸. Fine particulate matter export in streams depends on the activity of macroinvertebrates³⁶. Higher nocturnal activity has been documented for benthic macroinvertebrates in streams for decades³⁵. Studies that have evaluated diel suspended sediment dynamics, including a third-order Pennsylvania piedmont stream³⁷ and a cobble-bed Montana River³⁹ both found daytime minimums and nighttime maximums similar in magnitude to this study, from ~ 3 -9 $mg \cdot L^{-1}$ and ~ 0 -3 $mg \cdot L^{-1}$, respectively. Both studies concluded that bioturbation from benthic macroinvertebrates (as well as crayfish, amphibians, and eels in Pennsylvania) was the likely mechanism for increased nighttime concentrations. In the Pennsylvania stream, on average over eleven months, nighttime suspended solids and POC increased by 80% and 43%, respectively, over daytime minima. However, the diel variability was positively correlated with temperature and maximum variability was found in June where nighttime TSS and POC were 155% and 105% greater than daytime values, respectively. One other study assessing diel turbidity³⁸ found nighttime increases considerably larger than we observed, increasing from close to zero up to 20 $mg \cdot L^{-1}$ at night, however that site was located in an agricultural catchment that included cattle grazing as well as at least five species of bottom dwelling fish. The combination of greater sedimentation from animal induced bank erosion as well as larger and more mobile macrofauna, likely contributed to the relatively elevated nighttime TSS concentrations.

Similarly, in EFPC, the presence and magnitude of the diel turbidity pattern undergoes a seasonal cycle (Figure 8). There are no overnight increases in turbidity in late autumn and during winter. In spring the overnight increases in turbidity return, persist throughout summer, and dampen during autumn until they disappear again. The consistent relationship between turbidity and $Hg_{p.ngL}$ and $MMHg_{p.ngL}$ in seasons with unique diel dynamics indicate that this *in situ* measurement can be used to accurately estimate diel fluctuations in particulate concentration dynamics.

Other potential sources of diel variability in turbidity cannot account for the range of values or timing of our observations. The concordance between turbidity and TSS shows that other diel patterns did not

interfere with the turbidity instrument. For example, instrument manufacturers report a temperature dependence of up to -0.6% turbidity per degree Celsius increase. For the temperature range in Su2013 (1.9°C) this accounts for 1.1% variability in instrument reading compared to the 122% observed. Additionally, a similar temperature range occurred in the winter sampling, but no diel turbidity pattern was seen in that season.

Diel patterns in streamflow were not in phase with the turbidity and TSS patterns. For example, at EFK 5.4 during Su2015 a discharge-turbidity hysteresis plot shows turbidity remained constant while discharge varied and changed when discharge was constant (Figure 9). Finally, we cannot account for the turbidity patterns via pH-driven precipitation and dissolution reactions involving carbonate minerals (Figure S.14). The pH decreased overnight, favoring dissolution of carbonate minerals which would decrease turbidity and TSS, which is the opposite of the observed overnight increase. Lack of diel oscillations in calcium and magnesium concentrations also argue against cyclical precipitation-dissolution reactions affecting turbidity and TSS.

Benthic macroinvertebrate taxonomic richness in EFPC is similar to local reference streams for sampling locations near EFK 5.4 and EFK 16.2 whereas it is lower at EFK 23.4 and sites upstream of that location⁴⁰. Similarly, fish communities sampled near EFK 5.4 and EFK 16.2 are similar to local reference stream with respect to population density and the number of pollution-sensitive species⁴⁰ suggesting sufficient aquatic fauna in the lower sections of EFPC to cause the diel patterns in TSS and turbidity. The absence of diel patterns in TSS, $Hg_{P.ngL}$, and $MMHg_{P.ngL}$ at EFK 23.4 may be due to one or more reasons. Approximately 25% of the channel upstream of that sampling location is concrete lined with little sediment to be resuspended by fish or benthic invertebrates in addition to being poor habitat for the latter. The benthic macroinvertebrate community upstream of EFK 23.4 is lower density and less rich in pollutant tolerant species compared to downstream locations^{40, 41}. Finally, the EFK 23.4 site had the shortest upstream reach (2.6 km) of the three sites so integrates fewer cumulative upstream effects.

4.3 CONTROLS ON DISSOLVED Hg AND MMHg CONCENTRATIONS

The inconsistent relationship between Hg_D , $MMHg_D$, and DOC may be related to the high concentrations of Hg and MMHg in this contaminated stream coupled with the low and narrow range of DOC concentration across seasons and a broad land use gradient in the watershed (industrial, urban/suburban, forested). Dissolved Hg concentrations had diel oscillations at EFK 23.4 in Su2015 and EFK 5.4 in W2014. These correspond to sites and seasons with the least canopy cover. Much of the reach upstream of EFK 23.4 has little to no canopy cover and this site had the largest variation in other sunlight-driven variables (temperature, pH, dissolved O_2). While solar radiation was less intense and the daily

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3 474 photoperiod shorter in the winter sampling, there was no canopy, so sunlight affected a much longer
4 475 length of the creek upstream of our sampling location than during the summer samplings at EFK 5.4 or
5 476 EFK 16.2. Photochemical reactions may have promoted the release of particle-bound Hg leading to diel
6 477 oscillations in Hg_D in sync with the daily photocycle. This may be related to the inferred changes in
7 478 DOM composition as reflected in the oscillating values of $SUVA_{254}$ and FDOM (Figures S.10 and S.11).
8 479 Nevertheless, the different FDOM- Hg_D relationship between Su2013 and W2014 demonstrates that
9 480 additional research is needed before adopting FDOM as a reliable proxy measure for Hg_D . While each
10 481 parameter may be responsive to the daily photocycle the correlation between them remains unclear. The
11 482 changing DOC concentration at EFK 23.4 may also play a role in the oscillating Hg_D concentration but
12 483 the effect of DOC concentration alone is difficult to predict as numerous studies demonstrate that DOM
13 484 quantity and composition interact to affect Hg solid-water partitioning⁴²⁻⁴⁵.
14 485

15 486 In addition to the potential for photochemical reactions to directly affect Hg solid-water partitioning, there
16 487 are indirect light-driven reactions to consider. pH has long been acknowledged as a master variable
17 488 exerting control on sorption with cation sorption increasing with increasing pH and anion sorption
18 489 increasing with decreasing pH⁴⁶. In this regard, it is important to consider that we expect all the
19 490 dissolved Hg to be associated with DOM and that the Hg-DOM complex has a net negative charge at the
20 491 pH values in EFPC. Equilibrium aqueous speciation calculations indicate that >99% of the dissolved Hg
21 492 and MMHg would be present as DOM complexes⁴⁷ and laboratory experiments show the Hg-DOM
22 493 association to be kinetically fast⁴⁸. If these assumptions are correct then one would expect the light-
23 494 driven pH increase during the day to be accompanied by increasing Hg_D concentration, consistent with
24 495 our observations. In this case, the DOM would be a bridging ligand between surfaces and Hg and the
25 496 overall Hg partitioning behavior would be dominated by DOM sorption. Generalizations about the
26 497 sorption of heterogeneous mixtures like DOM are more problematic than for individual ions or molecules.
27 498 Nevertheless, DOM sorption onto a variety of minerals does seem to consistently decrease with
28 499 increasing pH and this pH dependence is particularly sensitive in the pH range 6.5-8.5 which
29 500 encompasses most of our observed range^{49,50}. Therefore, the diel pH oscillations likely contributed to the
30 501 oscillations in Hg_D concentration at EFK 23.4 and EFK 5.4.
31 502

32 503 Diel temperature changes can also affect sorption and this effect can be quantified by the reaction
33 504 enthalpy (ΔH°) using the van't Hoff equation. Still operating under the assumption that the
34 505 overwhelming majority of Hg_D is present as a Hg-DOM complex the apparent temperature effects on Hg
35 506 sorption will be dominated by the enthalpy of DOM sorption (also assuming temperature effects on Hg-
36 507 DOM complexation are negligible). Nguyen⁵¹ reported sorption enthalpies for Elliot Soil humic acid and
37 508 fulvic acid isolates onto hematite over the temperature range 15-35°C, showing the reactions to be

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3 509 exothermic for both isolates. Using their most extreme enthalpy value ($\Delta H^\circ = -60.7 \text{ kJ}\cdot\text{mol}^{-1}$) and our
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5 510 broadest diel temperature range (5.54°C , Su2015, EFK 23.4), diel temperature changes are predicted to
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7 511 change the K_{SW} for Hg by $0.2 \log_{10}$ units (i.e., if $\log_{10}(K_{\text{SW}, 25^\circ\text{C}}) = 6.0 \text{ L}\cdot\text{kg}^{-1}$, $\log_{10}(K_{\text{SW}, 19.5^\circ\text{C}}) = 6.2 \text{ L}\cdot\text{kg}^{-1}$).
8 512 This difference in $\log_{10}(K_{\text{SW}})$ is generally smaller than the observed variation (Figure S.15).
9
10 513 Therefore, while temperature changes could affect equilibrium Hg sorption, the predicted effect is small
11 514 and within the range of diel variability.

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13 515
14 516 MMHg_D concentrations showed diel oscillations during each campaign at all locations. Variability in diel
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16 517 MMHg_D concentrations in summer represented 26% - 38% of annual variability over the same period for
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18 518 monthly samples collected at a consistent time of day during baseflow conditions (Figure S.16). Other
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20 519 investigators have reported daytime MMHg_D maxima relative to nighttime minima in lakes, wetlands, and
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22 520 streams when sampled in spring and summer^{4, 9, 13, 52}. Dissolved phase MMHg concentrations are subject
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24 521 to some of the same controls that were discussed for Hg_D (photochemical release from surfaces, pH- and
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26 522 temperature-dependent sorption) with the same expected trends and magnitude of effects. Additionally,
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28 523 the measured MMHg concentrations are the net result of the opposing processes of Hg methylation and
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30 524 MMHg demethylation. Under the environmental conditions in this study, Hg methylation is an
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32 525 exclusively biotic process in which MMHg is generated by anaerobic bacteria and Archaea possessing the
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34 526 *hgcAB* two-gene cluster^{53, 54}. In contrast, MMHg demethylation can be mediated biotically by a broad
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36 527 array of aerobic and anaerobic microorganisms or by abiotic reactions^{12, 55, 56}.

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39 529 Hg methylation within algal biofilms (periphyton) in EFPC is an important process exerting control on
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41 530 the diel oscillations in MMHg_D concentration. Several investigators have reported net positive MMHg
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43 531 generation by algal biofilms⁵⁷⁻⁶⁵. Specifically, we have previously documented Hg methylation by EFPC
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45 532 periphyton at several sites in the creek and across seasons^{66, 67}. Importantly, MMHg generation by
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47 533 periphyton was significantly slower when biofilms were incubated in the dark or are grown under low
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49 534 light conditions with rates of net MMHg production between light and dark incubations varying by
50
51 535 1120%. In some experiments, biofilm samples incubated in the dark were net demethylating. Once
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53 536 formed inside bacterial cells, MMHg is rapidly exported to the surrounding media and can remain filter-
54
55 537 passing for up to 24 hours⁶⁸⁻⁷⁰. Rapid MMHg export from the cells coupled with turbulent advective flow
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57 538 in the stream mixes the periphyton-derived MMHg with the water column providing a link between
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59 539 production of MMHg in the periphyton and our observations of MMHg_D concentrations that are
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61 540 positively correlated with the daily photocycle.

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63 541
64 542 Another potentially important source of MMHg_D to surface water is the hyporheic zone. The Su2018
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66 543 campaign showed that in some locations interstitial porewater can have substantially higher MMHg_D

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3 544 concentration than the surface water (Figure S.6) and relatively small volumes of that water (<10%) could
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5 545 significantly impact surface water concentrations. But the porewater MMHg_D concentrations did not
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7 546 show diel oscillations suggesting either it is a small component of the surface water budget, or another,
8
9 547 oscillating, mechanism exerts control on the porewater-surface water exchange. Evapotranspiration (ET)
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11 548 driven oscillations in hyporheic-surface water exchange could have a diel signal in sync with the daily
12
13 549 photocycle but the direction of such a signal is the opposite of the observed effect. Increasing ET during
14
15 550 the day draws groundwater levels down creating a gradient from surface water to hyporheic and
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17 551 groundwater which is the opposite of that needed to deliver the higher MMHg_D concentration porewater
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19 552 to the surface water. Shorter flow paths disconnected from larger scale groundwater-surface water
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21 553 exchange may be involved but much more detailed hydrogeochemical characterization studies would be
22
23 554 needed to test these hypotheses. Additionally, these mechanisms may not transfer to other reaches of the
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25 555 creek where diel MMHg_D oscillations were seen. For example, the reach upstream of EFK 23.4 has few
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27 556 trees to drive the ET oscillations and has long channelized sections with no hyporheic zone.
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30
31 558 In addition to these potential sources of MMHg_D, an important demethylation mechanism to consider is
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33 559 photodemethylation which has been the subject of intense study in controlled lab settings and field studies
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35 560 ¹². Photodemethylation is a complex multivariate function of chemical (e.g., DOM composition and
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37 561 concentration) and physical (light duration and intensity, water depth) variables presenting significant
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39 562 challenges in generalizing results to our data ⁷¹. In controlled lab studies the rate and extent of MMHg_D
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41 563 photodemethylation were positively correlated to the concentration of an EFPC DOM isolate and was
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43 564 faster and more complete if the DOM was chemically reduced prior to use in experiments ⁷². However,
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45 565 photodemethylation would generate diel oscillations opposite to that which was observed. While
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47 566 photodemethylation was likely occurring in our sampling campaigns (particularly at EFK 23.4 in
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49 567 Su2015), the results suggest that the rate of instream Hg methylation was faster than photodemethylation
50
51 568 resulting in increasing concentrations during the day that peaked in mid to late afternoon. In W2014, the
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53 569 low water temperatures corresponded to an annual minimum in Hg methylation activity ^{66, 67, 73} and, given
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55 570 the lack of canopy cover over the creek, photodemethylation may have outpaced Hg methylation resulting
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57 571 in peak MMHg_D concentration after sunset.

572 **4.4 DISSOLVED GASEOUS MERCURY**

573 DGM, sampled only in Su2013 and W2014, showed similar patterns in both seasons and was positively
574 correlated with PAR and it is likely that the diel oscillations were attributable to photochemical reduction
575 of Hg(II) to generate Hg(0). Mid-day maxima and overnight minima in DGM concentration have been
576 reported in all seasons across a broad range of latitudes and correspondingly diverse range of water bodies
577 and chemistries ^{10, 74-79}. Siciliano et al. ⁷⁸ reported a diel cycle for DGM in two Canadian lakes with mid-

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3 578 day maxima and a corresponding cycle in microbial mercuric reductase activity that matched the DGM
4 579 cycle. Microbial mercury oxidase activity was offset from reductase activity and increased as DGM
5 580 decreased throughout the afternoon suggesting a role for microbial processes to contribute to abiotic
6 581 photochemical reactions in modulating DGM concentrations. Hg stable isotope patterns in EFPC suggest
7 582 that microbial Hg(II) reduction and, to a lesser extent, Hg(II) photoreduction contribute to Hg cycling in
8 583 the creek ⁸⁰. Notably, in the only other sunrise-to-sunset sampling study conducted in a Hg-contaminated
9 584 creek of which we are aware, DGM concentration peaked in mid-morning and declined throughout the
10 585 day ⁸¹.

16 586

17 587 5. CONCLUSIONS

19
20 588 Several diel sampling campaigns at multiple locations in EFPC confirm that portions of the Hg cycle in
21 589 freshwater streams undergo rapid fluctuations that are correlated, sometimes negatively, with the daily
22 590 photocycle. The oscillations for particulate phases (turbidity, TSS, Hg_{P.ngL}, MMHg_{P.ngL}) were consistent
23 591 within seasons, excluding the upstream industrial site, and were opposite in contrasting summer and
24 592 winter seasons for some dissolved phases (e.g., MMHg_D at the forested site). Both biotic (bioturbation,
25 593 Hg methylation) and abiotic (photochemical reactions, shading, pH variations) factors exerted control
26 594 over the observed patterns leading to loss of some diel oscillations in winter which had lower overall
27 595 biotic activity. In addition to providing insight into Hg transformations in freshwater stream ecosystems,
28 596 the diel oscillations have important implications with respect to sampling designs, site management, inter-
29 597 site comparisons, and uptake and toxicity to biotic receptors. For example, predictions of Hg
30 598 bioaccumulation factors are much better when high concentrations of MMHg_D are included in models ⁸².
31 599 Armed with that information, one might design sampling plans for MMHg_D to target mid to late afternoon
32 600 and keep sample timing consistent to ensure accurate assessments of long-term trends. Given the
33 601 potential importance of photochemical reactions on in-stream concentrations, changes in shading regime
34 602 (e.g., logging, canopy loss due to insect infestations) would be expected to have consequent effects on
35 603 MMHg concentration dynamics. At a minimum, stream shading regime is another factor to include in
36 604 inter-site comparisons of stream ecosystems. ⁸³Finally, from the perspective of ecosystem remediation,
37 605 given that much of the MMHg appears to originate from in-stream sources in this system, the relatively
38 606 rapid response of MMHg concentration to in-stream conditions gives hope that targeted actions could lead
39 607 to rapid improvements in water quality and, ultimately, in the health of biotic receptors in the
40 608 ecosystem⁸⁴.

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3 610 **6. DATA AVAILABILITY STATEMENT**
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6 611 The data presented in the paper and supplemental information are publicly available at
7 612 <https://msfa.ornl.gov/data/pages/MCI548.html> [doi.org/10.12769/1861075]. For additional information
8
9 613 please contact the corresponding author.
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12 615 **7. CONFLICTS OF INTEREST**
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14 616 There are no conflicts of interest to declare.
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16 617

17
18 618 **8. ACKNOWLEDGEMENTS**
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20
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30 625 contract no. DE-AC05-00OR22725.
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9. TABLES

Table 1. Summary of main diel patterns observed

Parameter	Season	Location	Peak Concentration	Mechanism
Hg _{P.ngL} , MMHg _{P.ngL} , TSS, Turbidity	Summer	Midstream urban/suburban, Downstream forested	Night	Sediment resuspension by benthic macrobiota
Hg _D	Summer, Winter	Upstream industrial, Downstream forested	Day	Photochemical release from particles; decreased sorption with pH change
MMHg _D	Summer	Upstream industrial, Midstream urban/suburban, Downstream forested	Day	MMHg production in periphyton biofilms
	Winter	Downstream forested	Night	Photodemethylation during day outpaced slower Hg methylation in colder temperatures

10. FIGURE CAPTIONS

Figure 1. Map of East Fork Poplar Creek and sampling sites in this study. Creek flow is from EFK 23.4 to EFK 5.4. SW = surface water, PW = pore water.

Figure 2. Temperature (A), dissolved oxygen concentration (B), and pH (C) for the diel sampling campaigns. Data were collected every 15 minutes, but symbols shown for every tenth data point for clarity. Shaded portions indicate the period from sunset to sunrise.

Figure 3. Turbidity and particulate Hg and MMHg ($\text{ng}\cdot\text{L}^{-1}$) concentration for the diel sampling campaigns. (A)-(C) Su2015 campaign at EFK 23.4 and EFK 16.2, (D)-(F) Su2013, W2014, and Su2015 campaigns at EFK 5.4. Shaded portions indicate the period from sunset to sunrise, the broader shaded portion in panels (D)-(F) and other figures represents nighttime for the W2014 campaign. Asterisks to the right of data points indicate potential outliers. Turbidity was not measured at EFK 23.4 in summer 2015.

Figure 4. Total suspended solids (symbols and lines) for each sampling campaign and site. The solid line in each plot represents the concurrently measured turbidity. Turbidity was not measured at EFK 23.4 in Su2015. Shaded portions indicate the period from sunset to sunrise. Asterisks to the right of data points indicate potential outliers.

Figure 5. DOC and dissolved total Hg and MMHg concentration for each diel sampling campaign. (A)-(C) Su2015 campaign at EFK 23.4 and EFK 16.2, (D)-(F) Su2013, W2014, and Su2015 campaigns at EFK 5.4. Shaded portions indicate the period from sunset to sunrise. Open symbols indicate potential outliers.

Figure 6. Dissolved gaseous mercury (DGM) at EFK 5.4 for the Su2013 and W2014 sampling campaigns. Shaded portions indicate the period from sunset to sunrise.

Figure 7. Water temperature (solid line) and dissolved MMHg concentration (symbols and dashed lines) in samples collected in the morning versus the afternoon from August 2013 through February 2014.

Figure 8. Turbidity versus discharge hysteresis curve for EFK 5.4 during the Su2015 campaign. The curve proceeds in a counterclockwise direction from the start of sampling where turbidity remained constant while discharge increased. After sunset, turbidity increased while discharge did not change. Turbidity then remained steady at the higher value while discharge decreased. Shortly after sunrise, turbidity decreased while discharge remained constant, then remained constant while discharge increased. Asterisk to the right of symbol indicates potential outlier.

Figure 9. Representative turbidity patterns at EFK 5.4 for each season. Shaded portions indicate the period from sunset to sunrise. The broad turbidity peaks in autumn and winter coincided with precipitation driven high flow events.

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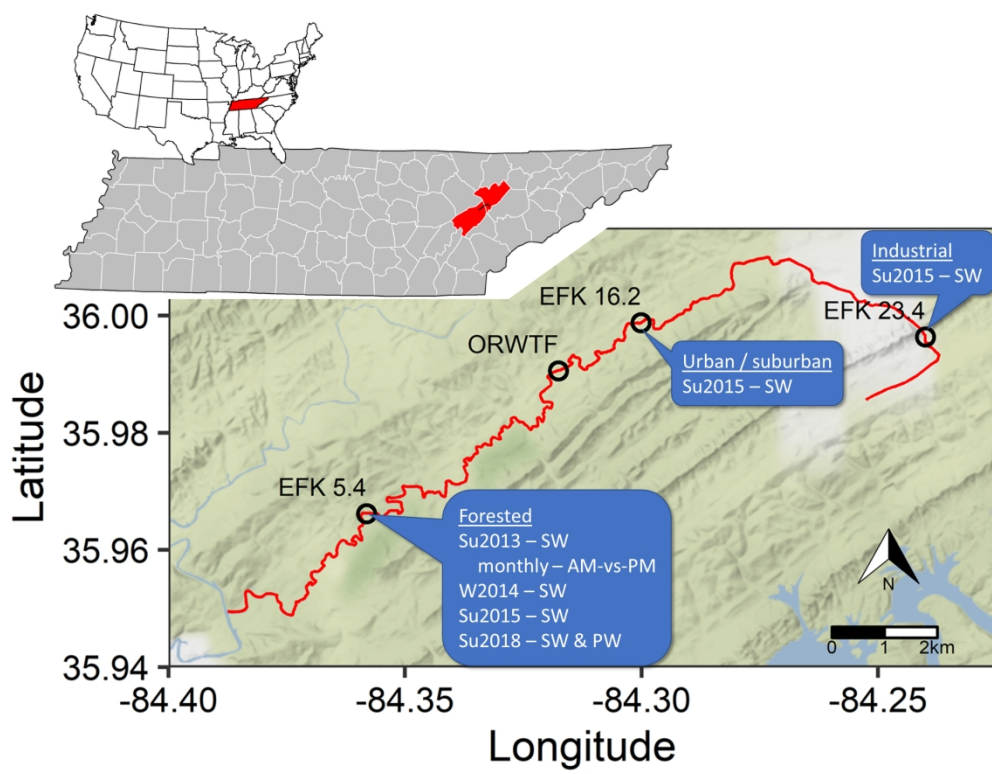


Figure 1. Map of East Fork Poplar Creek and sampling sites in this study. Creek flow is from EFK 23.4 to EFK 5.4. SW = surface water, PW = pore water.

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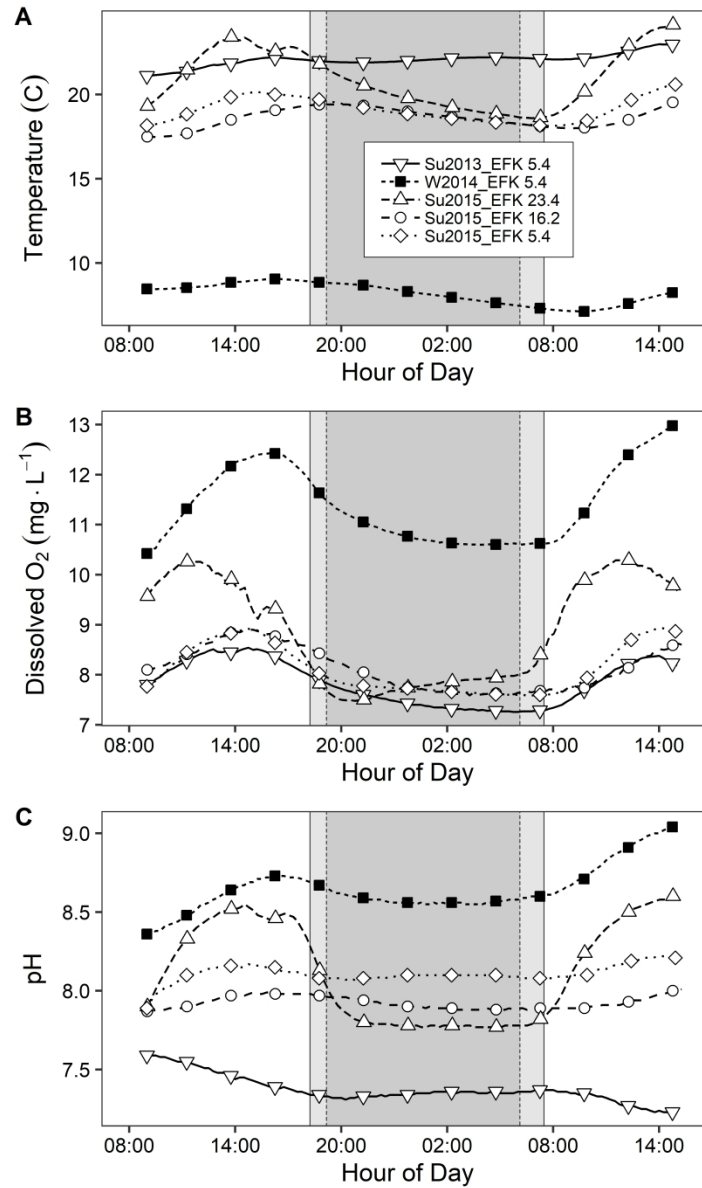


Figure 2. Temperature (A), dissolved oxygen concentration (B), and pH (C) for the diel sampling campaigns. Data were collected every 15 minutes, but symbols shown for every tenth data point for clarity. Shaded portions indicate the period from sunset to sunrise.

774x1290mm (126 x 126 DPI)

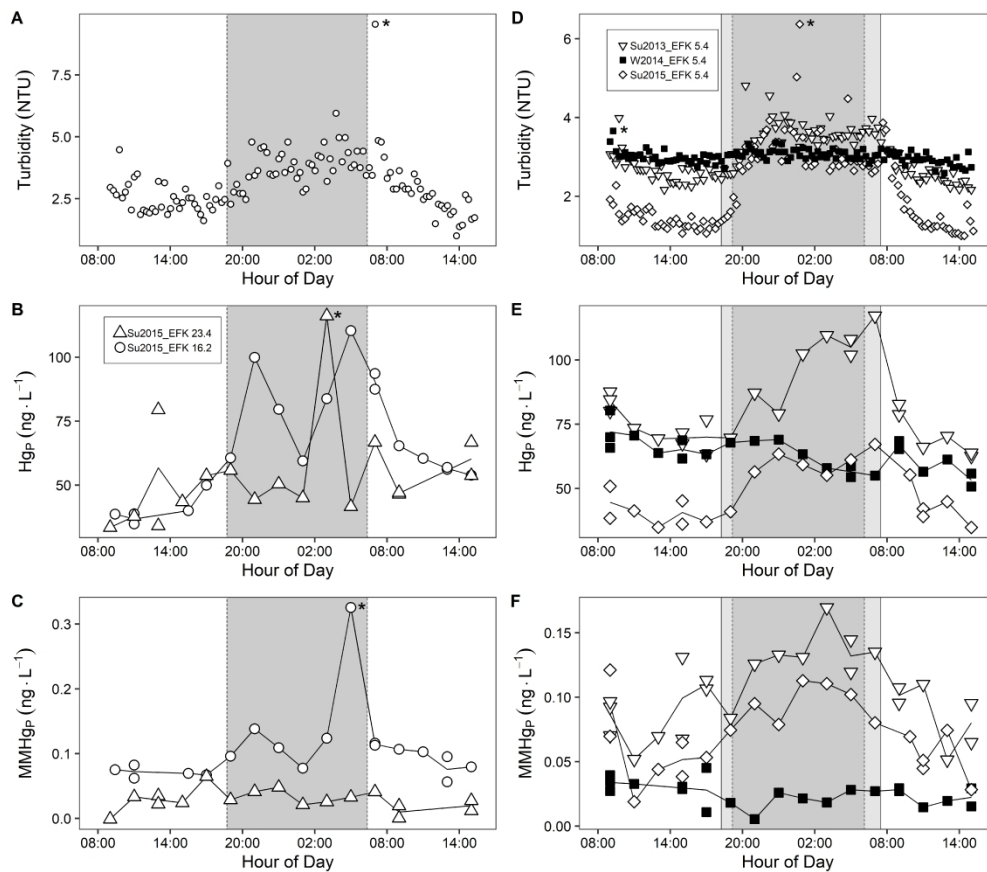


Figure 3. Turbidity and particulate Hg and MMHg ($\text{ng} \cdot \text{L}^{-1}$) concentration for the diel sampling campaigns. (A)-(C) Su2015 campaign at EFK 23.4 and EFK 16.2, (D)-(F) Su2013, W2014, and Su2015 campaigns at EFK 5.4. Shaded portions indicate the period from sunset to sunrise, the broader shaded portion in panels (D)-(F) and other figures represents nighttime for the W2014 campaign. Asterisks to the right of data points indicate potential outliers. Turbidity was not measured at EFK 23.4 in summer 2015.

1362x1192mm (126 x 126 DPI)

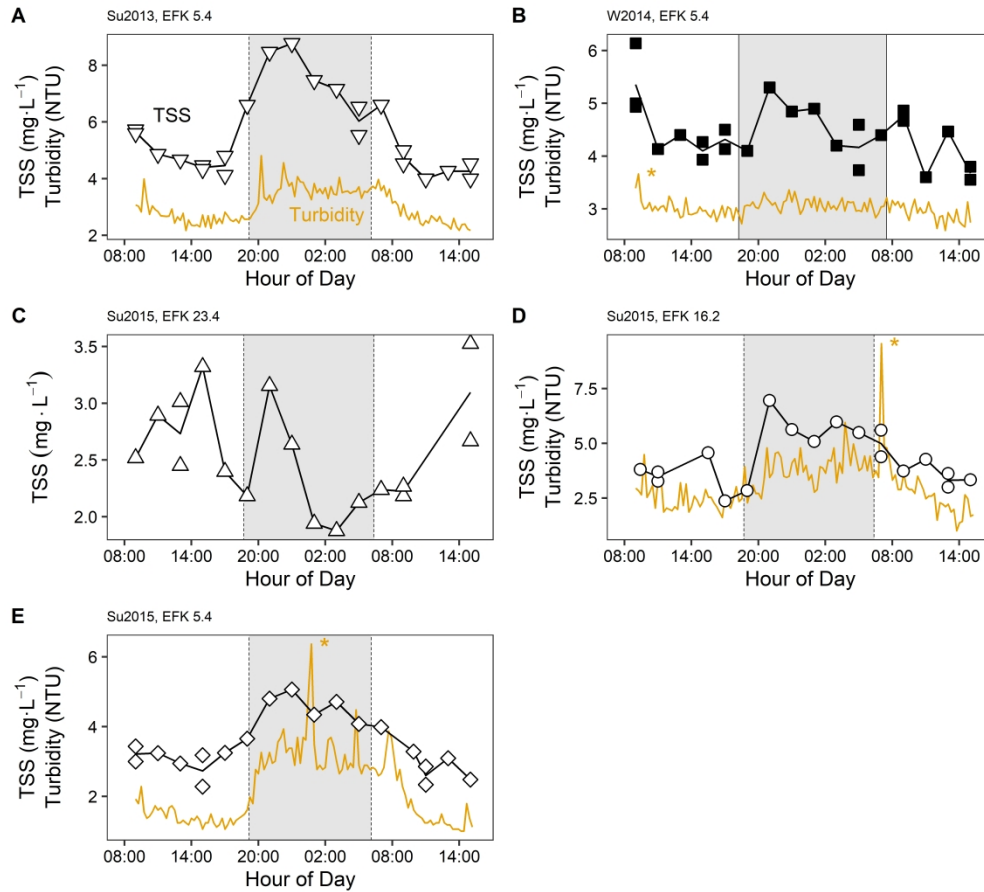


Figure 4. Total suspended solids (symbols and lines) for each sampling campaign and site. The solid line in each plot represents the concurrently measured turbidity. Turbidity was not measured at EFK 23.4 in Su2015. Shaded portions indicate the period from sunset to sunrise. Asterisks to the right of data points indicate potential outliers.

1290x1161mm (126 x 126 DPI)

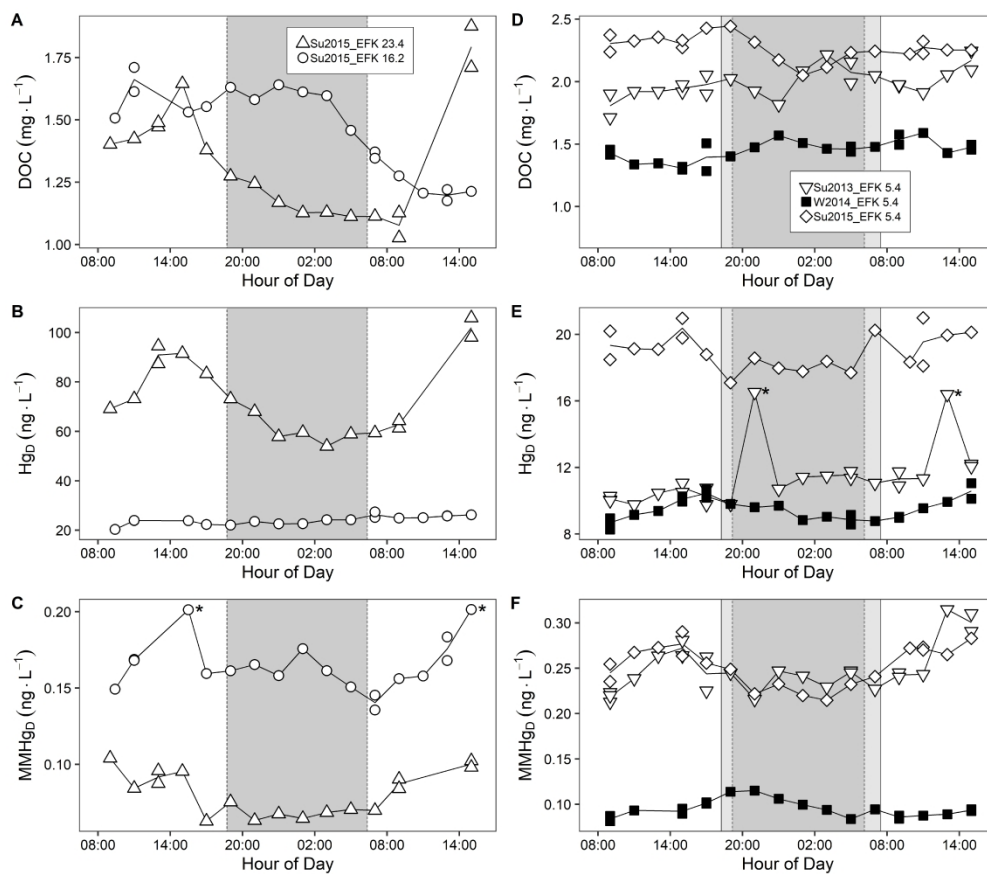


Figure 5. DOC and dissolved total Hg and MMHg concentration for each diel sampling campaign. (A)-(C) Su2015 campaign at EFK 23.4 and EFK 16.2, (D)-(F) Su2013, W2014, and Su2015 campaigns at EFK 5.4. Shaded portions indicate the period from sunset to sunrise. Open symbols indicate potential outliers.

1362x1192mm (126 x 126 DPI)

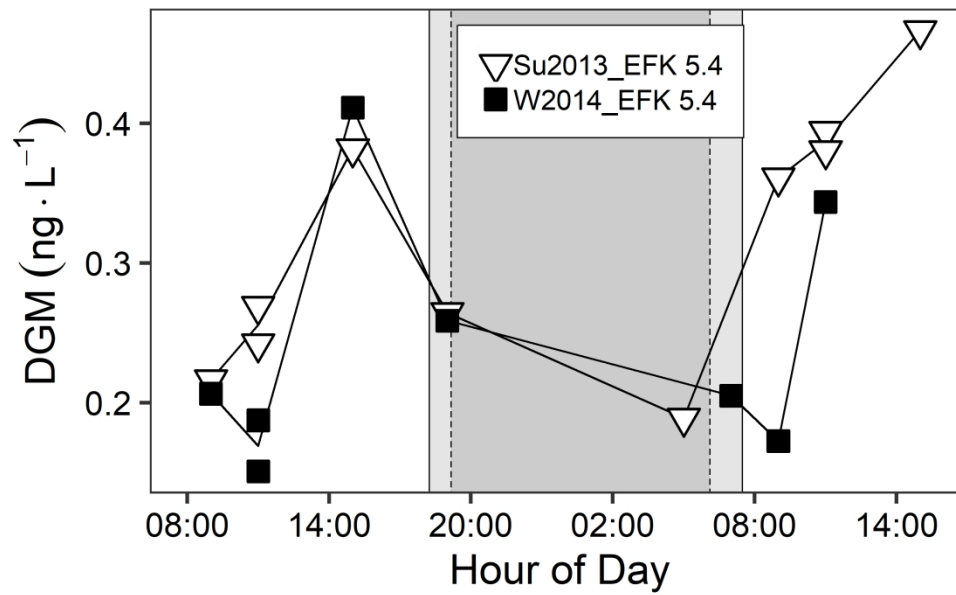


Figure 6. Dissolved gaseous mercury (DGM) at EFK 5.4 for the Su2013 and W2014 sampling campaigns. Shaded portions indicate the period from sunset to sunrise.

645x403mm (126 x 126 DPI)

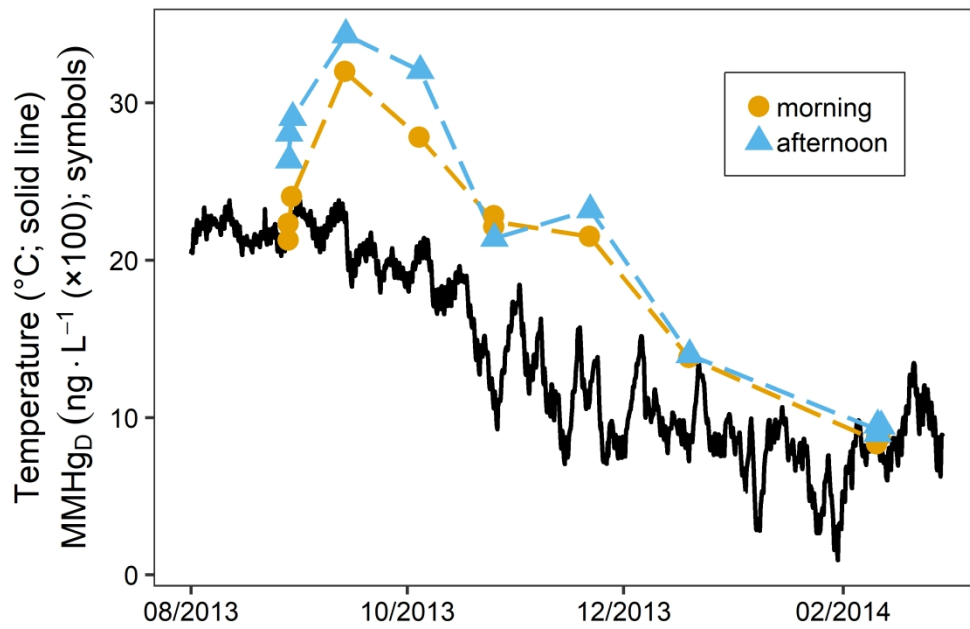


Figure 7. Water temperature (solid line) and dissolved MMHg concentration (symbols and dashed lines) in samples collected in the morning versus the afternoon from August 2013 through February 2014.

645x419mm (126 x 126 DPI)

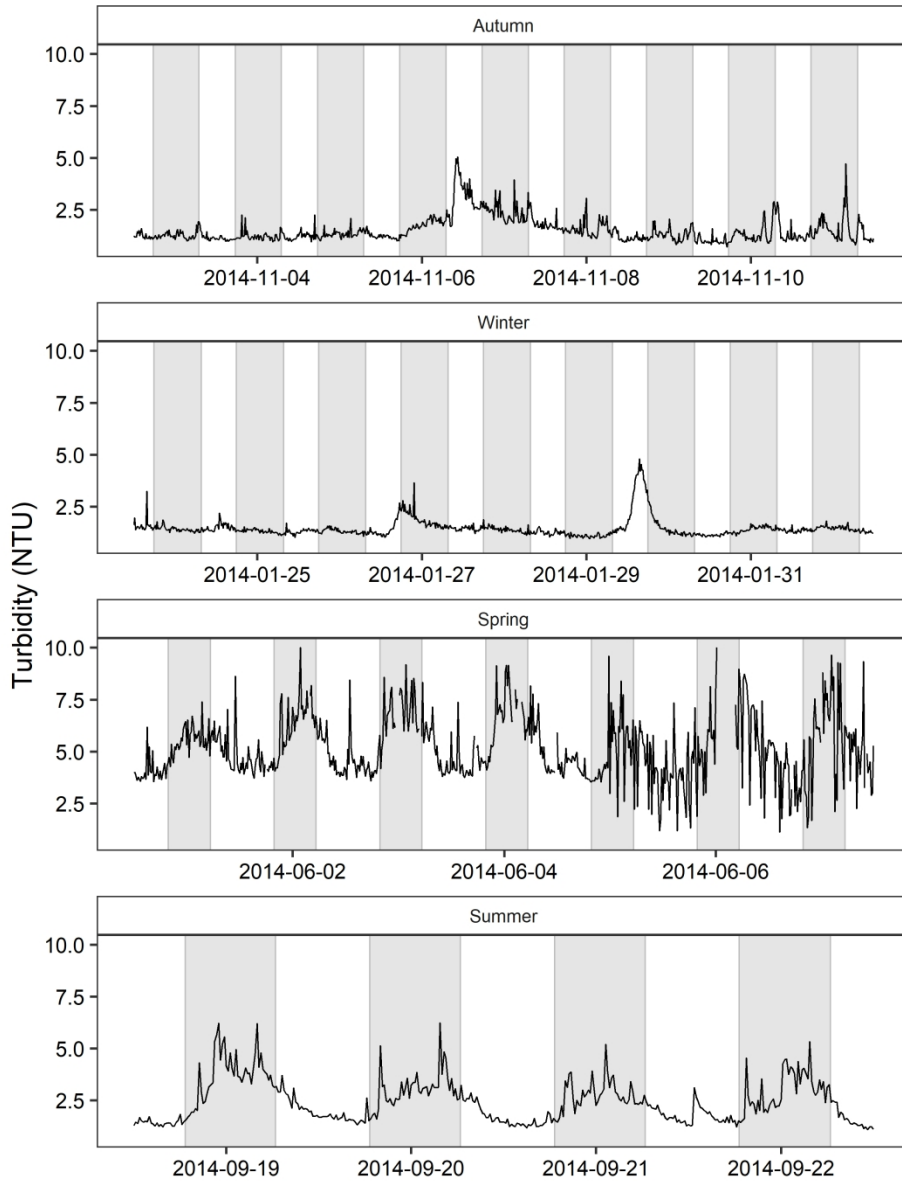


Figure 8. Representative turbidity patterns at EFK 5.4 for each season. Shaded portions indicate the period from sunset to sunrise. The broad turbidity peaks in autumn and winter coincided with precipitation driven high flow events.

548x709mm (126 x 126 DPI)

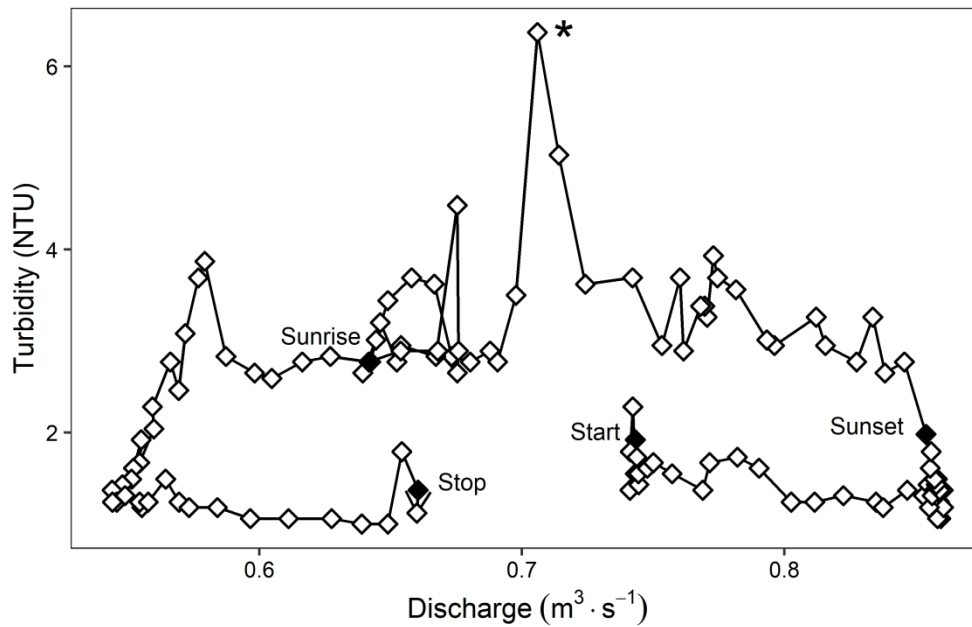


Figure 9. Turbidity versus discharge hysteresis curve for EFK 5.4 during the Su2015 campaign. The curve proceeds in a counterclockwise direction from the start of sampling where turbidity remained constant while discharge increased. After sunset, turbidity increased while discharge did not change. Turbidity then remained steady at the higher value while discharge decreased. Shortly after sunrise, turbidity decreased while discharge remained constant, then remained constant while discharge increased. Asterisk to the right of symbol indicates potential outlier.

645x419mm (126 x 126 DPI)